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The Integrated WRF/Urban Modeling System: Development, Evaluation, and Applications to Urban Environmental Problems

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39 **Abstract**

40 To bridge the gaps between traditional mesoscale modeling and microscale modeling, the
41 National Center for Atmospheric Research (NCAR), in collaboration with other agencies and
42 research groups, has developed an integrated urban modeling system coupled to the Weather
43 Research and Forecasting (WRF) model as a community tool to address urban environmental
44 issues. The core of this WRF/urban modeling system consists of: 1) three methods with
45 different degrees of freedom to parameterize urban surface processes, ranging from a simple
46 bulk parameterization to a sophisticated multi-layer urban canopy model with an indoor-
47 outdoor exchange sub-model that directly interacts with the atmospheric boundary layer, 2)
48 coupling to fine-scale Computational Fluid Dynamic (CFD) Reynolds-averaged Navier–Stokes
49 (RANS) and Large-Eddy Simulation (LES) models for Transport and Dispersion (T&D)
50 applications, 3) procedures to incorporate high-resolution urban land-use, building
51 morphology, and anthropogenic heating data using the National Urban Database and Access
52 Portal Tool (NUDAPT), and 4) an urbanized high-resolution land-data assimilation system (u-
53 HRLDAS). This paper provides an overview of this modeling system; addresses the daunting
54 challenges of initializing the coupled WRF/urban model and of specifying the potentially vast
55 number of parameters required to execute the WRF/urban model; explores the model
56 sensitivity to these urban parameters; and evaluates the ability of WRF/urban to capture urban
57 heat islands, complex boundary layer structures aloft, and urban plume T&D for several major
58 metropolitan regions. Recent applications of this modeling system illustrate its promising
59 utility, as a regional climate-modeling tool, to investigate impacts of future urbanization on
60 regional meteorological conditions and on air quality under future climate change scenarios.

61

62 **1 Introduction**

63 We describe in this paper an international collaborative research and development effort
64 between the National Center for Atmospheric Research (NCAR) and partners with regards to a
65 coupled land surface and urban modeling system for the community Weather Research and
66 Forecasting (WRF) model. The goal of this collaboration is to develop a cross-scale modeling
67 capability that can be used to address a number of emerging environmental issues in urban
68 areas.

69 Today's changing climate poses two formidable challenges. On one hand, the projected
70 climate change by IPCC (Fourth Assessment Report, 2007) may lead to more frequent
71 occurrences of heat waves, severe weather, and floods. On the other hand, the current trend of
72 population increase and urban expansion is expected to continue. For instance, in 2007 half of
73 the world's population lived in cities, and that proportion is projected to be 60% in 2030
74 (United Nations, 2007). The combined effect of global climate change and rapid urban growth,
75 accompanied with economic and industrial development, will likely make people living in
76 cities more vulnerable to a number of urban environmental problems, including: extreme
77 weather and climate conditions, sea-level rise, poor public health and air quality, atmospheric
78 transport of accidental or intentional releases of toxic material, and limited water resources.
79 For instance, Nicholls et al. (2007) suggested that by the 2070s, total world population exposed
80 to coastal flooding could grow more than threefold to around 150 million people due to the
81 combined effects of climate change (sea-level rise and increased storminess), atmospheric
82 subsidence, population growth, and urbanization. The total asset exposure could grow even
83 more dramatically, reaching US \$35,000 billion by the 2070s. Zhang et al. (2009)

84 demonstrated that urbanization contributes to a reduction of summer precipitation in Beijing,
85 and that augmenting city green-vegetation coverage would enhance summer rainfall and
86 mitigate the increasing threat of water shortage in Beijing.

87 It is therefore imperative to understand and project effects of future climate change and
88 urban growth on the above environmental problems and to develop mitigation and adaptation
89 strategies. One valuable tool for this purpose is a cross-scale atmospheric modeling system,
90 which is able to predict/simulate meteorological conditions from regional to building scales
91 and which can be coupled to human-response models. The community WRF model, often
92 executed with a grid spacing of 0.5-1 km, is in a unique position to bridge gaps in traditional
93 mesoscale numerical weather prediction ($\sim 10^5$ m) and microscale T&D modeling ($\sim 10^0$ m).
94 One key requirement for urban applications is for WRF to accurately capture influences of
95 cities on wind, temperature, and humidity in the atmospheric boundary layer and their
96 collective influences on the atmospheric mesoscale motions.

97 Remarkable progress has been made in the last decade to introduce a new generation of
98 urbanization schemes into atmospheric models such as the Fifth-generation Pennsylvania State
99 University (PSU)–NCAR Mesoscale Model (MM5) (Taha, 1999, Taha and Bornstein, 1999,
100 Dupont et al., 2004, Liu et al., 2006, Otte et al., 2004, Taha 2008a,b), WRF model (Chen et al.,
101 2004), UK Met Office operational mesoscale model (Best, 2005), French Meso-NH (Lemonsu
102 and Masson 2002) model, and NCAR global climate model (Oleson et al., 2008). Moreover,
103 fine-scale models, such as computational fluid dynamics models (Coirier et al., 2005) and fast-
104 response urban T&D models (Brown 2004), can explicitly resolve airflows around city
105 buildings. However, these parameterization schemes vary considerably in their degrees of
106 freedom to treat urban processes. An international effort is thus underway to compare these

107 urban models and to evaluate them against site observations (Grimmond et al., 2010). It is,
108 nonetheless, not clear at this stage which degree of complexity of urban modeling should be
109 incorporated in atmospheric models, given that the spatial distribution of urban land-use and
110 building morphology is highly heterogeneous even at urban scales and given the wide range of
111 applications such a model may be used for.

112 WRF is used for both operations and research in the fields of numerical weather prediction,
113 regional climate, emergency response, air quality (through its companion online chemistry
114 model WRF-Chem, Grell et al., 2005), and regional hydrology and water resources. In WRF-
115 Chem, the computations of meteorology and atmospheric chemistry share the same vertical
116 and horizontal coordinates, surface parameterizations (and hence same urban models), physics
117 parameterization for subgrid-scale transport, vertical mixing schemes, and time steps for
118 transport and vertical mixing. Therefore, our goal is to develop an integrated WRF/urban
119 modeling system to satisfy this wide range of WRF applications. As shown in Fig. 1, the core
120 of this system consists of: 1) a suite of urban parameterization schemes with varying degrees of
121 complexities, 2) the capability of incorporating in-situ and remotely-sensed data of urban land-
122 use, building characteristics, anthropogenic heating, and moisture sources, 3) companion fine-
123 scale atmospheric and urbanized land data assimilation systems, and 4) the ability to couple
124 WRF/urban with fine-scale urban T&D models and chemistry models. It is anticipated that in
125 the future, this modeling system will interact with human response models and be linked to
126 urban decision systems.

127 In the next section we describe the integrated WRF/urban modeling system. We address the
128 issue of initializing the state variables required to run WRF/urban in Section 3 and the issue of
129 specifying urban parameters and model sensitivity to these parameters in Section 4. Section 5

130 gives examples of model evaluation and of applying the WRF/urban model to various
131 urbanization problems, and it is followed by a summary in Section 6.

132 **2 Description of the integrated WRF/urban modeling system**

133 **2.1 Modeling system overview**

134 The WRF model (Skamarock et al., 2005) is a non-hydrostatic, compressible model with a
135 mass coordinate system. It was designed as a numerical weather prediction model, but can also
136 be applied as a regional climate model. It has a number of options for various physical
137 processes. For example, WRF has a non-local closure planetary boundary layer (PBL) scheme
138 and a 2.5 level PBL scheme based on the Mellor and Yamada scheme (Janjic, 1994). Among
139 its options for land surface models (LSMs), the community Noah LSM has been widely used
140 (e.g., Chen et al., 1996, Chen and Dudhia, 2001, Ek et al., 2003; Leung et al., 2006, Jiang et al.,
141 2008) in weather prediction models; in land data assimilation systems, such as the North
142 America Land Data Assimilation System (Mitchell et al., 2004); and in the community
143 mesoscale MM5 and WRF models.

144 One basic function of the Noah LSM is to provide surface sensible and latent heat fluxes
145 and surface skin temperature as lower boundary conditions for coupled atmospheric models. It
146 is based on a diurnally-varying Penman potential evaporation approach, a multi-layer soil
147 model, a modestly complex canopy resistance parameterization, surface hydrology, and frozen
148 ground physics (Chen et al. 1996; Chen et al., 1997; Chen and Dudhia 2001; Ek et al. 2003).
149 Prognostic variables in Noah include: liquid water, ice, and temperature in the soil layers;
150 water stored in the vegetation canopy; and snow water equivalent stored on the ground.

151 Here, we mainly focus the urban modeling efforts on coupling different urban canopy
152 models (UCMs) with Noah in WRF. Such coupling is through the parameter urban percentage

153 (or urban fraction, F_{urb}) that represents the proportion of impervious surfaces in the WRF sub-
154 grid scale. For a given WRF grid cell, the Noah model calculates surface fluxes and
155 temperature for vegetated urban areas (trees, parks, etc.) and the UCM provides the fluxes for
156 anthropogenic surfaces. The total grid-scale sensible heat flux, for example, can be estimated
157 as follows:

$$158 \quad Q_H = F_{veg} \times Q_{Hveg} + F_{urb} \times Q_{Hurb}$$

159 where Q_H is the total sensible heat flux from the surface to the WRF model lowest
160 atmospheric layer, F_{veg} the fractional coverage of natural surfaces, such as grassland, shrubs,
161 crops, and trees in cities, F_{urb} the fractional coverage of impervious surfaces, such as buildings,
162 roads, and railways. Q_{Hveg} the sensible heat flux from Noah for natural surfaces, and Q_{Hurb} the
163 sensible heat flux from the UCM for artificial surfaces. Grid-integrated latent heat flux, upward
164 long wave radiation flux, albedo, and emissivity are estimated in the same way. Surface skin
165 temperature is calculated as the averaged value of the artificial and natural surface temperature
166 values, and is subsequently weighted by their areal coverage.

167 **2.2 Bulk urban parameterization**

168 The WRF V2.0 release in 2003 included a bulk urban parameterization in Noah using the
169 following parameter values to represent zero-order effects of urban surfaces (Liu et al., 2006):
170 1) roughness length of 0.8 m to represent turbulence generated by roughness elements and drag
171 due to buildings; 2) surface albedo of 0.15 to represent shortwave radiation trapping in urban
172 canyons; 3) volumetric heat capacity of $3.0 \text{ J m}^{-3} \text{ K}^{-1}$ for urban surfaces (walls, roofs, and
173 roads), assumed as concrete or asphalt; 4) soil thermal conductivity of $3.24 \text{ W m}^{-1} \text{ K}^{-1}$ to
174 represent the large heat storage in urban buildings and roads; and 5) reduced green-vegetation
175 fraction over urban areas to decrease evaporation. This approach has been successfully

176 employed in real-time weather forecasts (Liu et al., 2006) and to study the impact of
177 urbanization on land-sea breeze circulations (Lo et al., 2007).

178 **2.3 Single-layer urban canopy model**

179 The next level of complexity incorporated uses the single-layer UCM (SLUCM) developed
180 by Kusaka et al. (2001) and Kusaka and Kimura (2004). It assumes infinitely-long street
181 canyons parameterized to represent urban geometry, but recognizes the three dimensional
182 nature of urban surfaces. In a street canyon, shadowing, reflections, and trapping of radiation
183 are considered, and an exponential wind profile is prescribed. Prognostic variables include:
184 surface skin temperatures at the roof, wall, and road (calculated from the surface energy
185 budget) and temperature profiles within roof, wall, and road layers (calculated from the
186 thermal conduction equation). Surface sensible heat fluxes from each facet are calculated using
187 Monin-Obukhov similarity theory and the Jurges formula (Fig. 2). The total sensible heat flux
188 from roof, wall, roads, and the urban canyon is passed to the WRF-Noah model as $Q_{H_{urb}}$
189 (Section 2.1). The total momentum flux is passed back in a similar way. SLUCM calculates
190 canyon drag coefficient and friction velocity using a similarity stability function for
191 momentum. Total friction velocity is then aggregated from urban and non-urban surfaces and
192 passed to WRF boundary layer schemes. Anthropogenic heating and its diurnal variation are
193 considered by adding them to the sensible heat flux from the urban canopy layer. SLUCM has
194 about 20 parameters, as listed in Table 1.

195 **2.4 Multi-layer urban canopy (BEP) and indoor-outdoor exchange (BEM) models**

196 Unlike the SLUCM (embedded within the first model layer), the multi-layer UCM
197 developed by Martilli et al. (2002), called BEP for Building Effect Parameterization, represents
198 the most sophisticated urban modeling in WRF, and it allows a direct interaction with the PBL

199 (Fig. 2). BEP recognizes the three-dimensional nature of urban surfaces and the fact that
200 buildings vertically distributes sources and sinks of heat, moisture, and momentum through the
201 whole urban canopy layer, which substantially impacts the thermodynamic structure of the
202 urban roughness sub-layer and hence the lower part of the urban boundary layer. It takes into
203 account effects of vertical (walls) and horizontal (streets and roofs) surfaces on momentum
204 (drag force approach), turbulent kinetic energy, and potential temperature (Fig. 2). The
205 radiation at walls and roads considers shadowing, reflections, and trapping of shortwave and
206 longwave radiation in street canyons. The Noah-BEP model has been coupled with two
207 turbulence schemes: Bougeault and Lacarrere (1989) and Mellor-Yamada-Janjic (Janjic, 1994)
208 in WRF by introducing a source term in the TKE equation within the urban canopy and by
209 modifying turbulent length scales to account for the presence of buildings. As illustrated in Fig.
210 3, BEP is able to simulate some of the most observed features of the urban atmosphere, such as
211 the nocturnal Urban Heat Island (UHI) and the elevated inversion layer above the city.

212 To take full advantage of BEP, it is necessary to have high vertical resolution close to the
213 ground (to have more than one model level within the urban canopy). Consequently, this
214 approach is more appropriate for research (when computational demands are not a constraint)
215 than for real-time weather forecasts.

216 In the standard version of BEP (Martilli et al., 2002), the internal temperature of the
217 buildings is kept constant. To improve estimation of exchanges of energy between the interior
218 of buildings and the outdoor atmosphere, which can be an important component of the urban
219 energy budget, a simple Building Energy Model (BEM, Salamanca and Martilli, 2009) has
220 been developed and linked to BEP. BEM accounts for the: 1) diffusion of heat through the
221 walls, roofs, and floors; 2) radiation exchanged through windows; 3) longwave radiation

222 exchanged between indoor surfaces; 4) generation of heat due to occupants and equipment; and
223 5) air conditioning, ventilation, and heating. Buildings of several floors can be considered, and
224 the evolution of indoor air temperature and moisture can be estimated for each floor. This
225 allows the impact of energy consumption due to air conditioning to be estimated. The coupled
226 BEP+BEM has been tested offline using the BUBBLE (Basel UrBan Boundary Layer
227 Experiment, Rotach et al., 2005) data. Incorporating building energy in BEP+BEM
228 significantly improves sensible heat-flux calculations over using BEP alone (Fig. 4). The
229 combined BEP+BEM has been recently implemented in WRF, and is currently being tested
230 before its public release in WRF V3.2 in Spring 2010.

231 **2.5 Coupling to fine-scale Transport and Dispersion (T&D) models**

232 Because WRF can parameterize only aggregated effects of urban processes, it is necessary
233 to couple it with finer-scale models for applications down to building-scale problems. One key
234 requirement for fine-scale T&D modeling is to obtain accurate, high-resolution meteorological
235 conditions to drive T&D models. These are often incomplete and inconsistent, due to limited
236 and irregular coverage of meteorological stations within urban areas. To address this limitation,
237 fine-scale building-resolving models, e.g., Eulerian/semi-Lagrangian fluid solver (EULAG)
238 and CFD-Urban, are coupled to WRF to investigate the degree to which the: 1) use of WRF
239 forecasts for initial and boundary conditions can improve T&D simulations through
240 downscaling and 2) feedback, through upscaling, of explicitly resolved turbulence and wind
241 fields from T&D models can improve WRF forecasts in complex urban environments.

242 In the coupled WRF-EULAG/CFD-Urban models (Fig. 5), WRF generates mesoscale (~1-
243 10 km) atmospheric conditions to provide initial and boundary conditions, through
244 downscaling, for microscale (~1-10 m) EULAG/CFD-Urban simulations. WRF meso-scale

245 simulations are performed usually at 500 m grid spacing. Data from WRF model (i.e., grid
246 structure information, horizontal and vertical velocity components, and thermodynamic fields,
247 such as pressure, temperature, water vapor, as well as turbulence) are saved at appropriate time
248 intervals (usually each 5-15 min) required by CFD simulations. WRF model grid structure and
249 coordinates are transformed to the CFD model grid before use in the simulations.

250 The CFD-Urban model resolve building structures explicitly by considering different urban
251 aerodynamic features, such as channeling, enhanced vertical mixing, downwash, and street-
252 level flow. These microscale flow features can be aggregated and transferred back, through
253 upscaling, to WRF to increase the accuracy of mesoscale forecasts for urban and downstream
254 regions. The models can be coupled in real time; and data transfer is realized through the
255 Model Coupling Environmental Library (MCEL).

256 As an example, Tewari et al. (2010) ran the WRF model at a sub-kilometer resolution (0.5
257 km), and its temporal and spatial meteorological fields were downscaled and used in the
258 unsteady coupling mode to supply initial and time-varying boundary conditions to the CFD-
259 Urban model developed by Coirier et al. (2005). Traditionally, most CFD models used for
260 T&D studies are initialized with a single profile of atmospheric sounding data, which does not
261 represent the variability of weather elements within urban areas. This often results in errors in
262 predicting urban plumes. The CFD-Urban T&D predictions using the above two methods of
263 initialization were evaluated against the URBAN 2000 field experiment data for Salt Lake City
264 (Allwine et al., 2002). For concentrations of a passive tracer, the WRF-CFD-Urban
265 downscaling better produced the observed high-concentration tracer in the northwestern part of
266 the downtown area, largely due to the fact that the turning of lower boundary layer wind to
267 NNW from N is well represented in WRF and the imposed WRF simulated pressure gradient is

268 felt by the CFD-Urban calculations (Fig. 6). These improved steady-state flow fields result in
269 significantly improved plume transport behavior and statistics.

270 The NCAR LES model EULAG has been coupled to WRF. EULAG is a multi-scale, multi-
271 physics computational model for simulating urban canyon thermodynamic and transport fields
272 across a wide range of scales and physical scenarios (see Prusa et al., 2008 for a review). Since
273 turbulence in the mesoscale model (WRF in our case) is parameterized, there is no direct
274 downscaling of the turbulent quantities (e.g., TKE) from WRF to the LES model. The LES
275 model assumes the flow at the boundaries to be laminar (with small scale random noise added
276 to the mean flow), and the transition zone is preserved between the model boundary and
277 regions where the turbulence develops internally within the LES model domain. Contaminant
278 transport in urban areas is simulated with a passive tracer in time-dependent adaptive mesh
279 geometries (Wyszogrodzki and Smolarkiewicz, 2009). Building structures are explicitly
280 resolved using the immersed boundary (IMB) approach, where fictitious body forces in the
281 equations of motion represent internal boundaries, effectively imposing no-slip boundary
282 conditions at building walls (Smolarkiewicz et al., 2007). The WRF/EULAG coupling with a
283 downscaling data transfer capability was applied for the daytime Intensive Observation Period
284 (IOP)-6 case during the Joint Urban Oklahoma City 2003 experiment (JU2003, Allwine et al.,
285 2004). With five two-way nested domains, with grid spacing ranging from 0.5 to 40 km, the
286 coupled model was integrated from 1200UTC 16 July 2003 (0700CDT) for a 12-h simulation.
287 WRF was able to reproduce the observed horizontal wind and temperature fields near the
288 surface and in the boundary layer reasonably well. The macroscopic features of EULAG-
289 simulated flow compare well with measurements. Figure 7 shows EULAG-generated near-

290 surface wind and dispersion of the passive scalar from the first release of IOP-6, starting at
291 0900 CDT.

292 **3 Challenges in initializing the WRF/urban model system**

293
294 Executing the coupled WRF/urban modeling system raises two challenges: 1) initialization
295 of the detailed spatial distribution of UCM state variables, such as temperature profiles within
296 wall, roofs, and roads and 2) specification of a potentially vast number of parameters related to
297 building characteristics, thermal properties, emissivity, albedo, anthropogenic heating, etc. The
298 former issue is discussed in this section and the latter in Section 4.

299 High-resolution routine observations of wall/roof/road temperature are rarely available to
300 initialize the WRF/urban model, which usually covers a large domain (e.g., $\sim 10^6$ km²) and may
301 include urban areas with a typical size of $\sim 10^2$ km². Nevertheless, to a large extent, this
302 initialization problem is analogous to that of initializing soil moisture and temperature in a
303 coupled atmospheric-land surface model. One approach is to use observed rainfall, satellite-
304 derived surface solar insolation, and meteorological analyses to drive an uncoupled (off-line)
305 integration of an LSM, so that the evolution of the modeled soil state can be constrained by
306 observed forcing conditions. The North-American Land Data Assimilation System (NLDAS,
307 Mitchell et al., 2004) and the NCAR High-Resolution Land Data Assimilation System
308 (HRLDAS, Chen et al., 2007) are two examples that employ this method. In particular,
309 HRLDAS was designed to provide consistent land-surface input fields for WRF nested
310 domains and is flexible enough to use a wide variety of satellite, radar, model, and in-situ data
311 to develop an equilibrium soil state. The soil state spin-up may take up to several years and
312 thus cannot be reasonably handled within the computationally-expensive WRF framework
313 (Chen et al., 2007).

314 Therefore, the approach adopted is to urbanize HRLDAS (u-HRLDAS) by running the
315 coupled Noah/urban model in an offline mode to provide initial soil moisture, soil temperature,
316 snow, vegetation, and wall/road/roof temperature profiles. As an example, a set of experiments
317 with the u-HRLDAS using Noah/SLUCM was performed for the Houston region. Similar to
318 Chen et al. (2007), an 18-month u-HRLDAS simulation was considered long enough for the
319 modeling system to reach an equilibrium state, and the temperature difference ΔT between this
320 18-month simulation and other simulations with shorter simulation period (e.g., 6 months, 2
321 months, etc.) is used to investigate the spin-up of SLUCM. The time required for SLUCM state
322 variables to reach a quasi-equilibrium state ($\Delta T < 1$ K) is short (less than a week) for roof and
323 wall temperature (Fig. 8), but longer (approximately two months) for road temperature, due to
324 the larger thickness and thermal capacity of roads. However, this spin-up is considerably
325 shorter than that for natural surfaces (up to several years, Chen et al., 2007). Results also show
326 that the spun-up temperatures of roofs, walls, and roads are different (by ~ 1 -2 K) and exhibit
327 strong horizontal heterogeneity in different urban land-use and buildings. Using a uniform
328 temperature to initialize WRF/urban will not capture such urban variability.

329 **4 Challenges in specifying parameters for urban models**

330 **4.1 Land-use based approach, gridded data set, and NUDAPT**

331 Using UCMs in WRF requires users to specify at least 20 urban canopy parameters (UCPs)
332 (Table 1). A combination of remote-sensing and in-situ data can be used for this purpose
333 thanks to recent progress in developing UCP data sets (Burian et al., 2004, Feddema et al.,
334 2006, Taha, 2008b, Ching et al., 2009). While the availability of these data is growing, data
335 sets are currently limited to a few geographical locations. High-resolution data sets on global
336 bases comprising the full suite of UCPs simply do not exist. In anticipation of increased

337 database coverage, we employ three methods to specify UCPs in WRF/urban: 1) urban land-
338 use maps and urban-parameter tables, 2) gridded high-resolution UCP data sets, and 3) a
339 mixture of the above.

340 For many urban regions, high-resolution urban land-use maps, derived from in-situ
341 surveying (e.g., urban planning data) and remote-sensing data (e.g., Landsat 30-m images) are
342 readily available. We currently use the USGS National Land Cover Data (NLCD)
343 classification with three urban land-use categories: 1) low-intensity residential, with a mixture
344 of constructed materials and vegetation (30-80 % covered with constructed materials), 2) high-
345 intensity residential, with highly-developed areas such as apartment complexes and row houses
346 (usually with 80-100 % covered with constructed materials), and 3) commercial/industrial/
347 transportation including infrastructure (e.g., roads, railroads, etc.). An example of the spatial
348 distribution of urban land-use for Houston is given in Fig. 9. Once the type of urban land-use is
349 defined for each WRF model grid, urban morphological and thermal parameters can be
350 assigned using the urban-parameters in Table 1. Although this approach may not provide the
351 most accurate UCP values, it captures some degree of their spatial heterogeneity, given the
352 limited input land-use-type data.

353 The second approach, to directly incorporate gridded UCPs into WRF, was tested in the
354 context of the National Urban Database and Access Portal Tool (NUDAPT) project (Ching et
355 al., 2009). NUDAPT was developed to provide the requisite gridded sets of UCPs for
356 urbanized WRF and other advanced urban meteorological, air quality, and climate modeling
357 systems. These UCPs account for the aggregated effect of sub-grid building and vegetation
358 morphology on grid-scale properties of the thermodynamics and flow fields in the layer
359 between the surface and the top of the urban canopy. High definition (1 to 5 m) three-

360 dimensional data sets of individual buildings, conglomerates of buildings, and vegetation in
361 urban areas are now available, based on airborne lidar systems or photogrammetric techniques,
362 to provide the basis for these UCPs (Burian et al., 2004, 2006, 2007). Each cell can have a
363 unique combination of UCPs. Currently, NUDAPT hosts datasets (originally acquired by the
364 National Geospatial Agency, NGA) for more than 40 cities in the United States, with different
365 degrees of coverage and completeness for each city. In the future, it is anticipated that high-
366 resolution building data will become available for other cities. With this important core-design
367 feature, and by using web portal technology, NUDAPT can serve as the database infrastructure
368 for the modeling community to facilitate customizing of data handling and retrievals
369 (<http://www.nudapt.org>) for such future datasets and applications in WRF and other models.

370

371 **4.2 Incorporating anthropogenic heat sources**

372 The scope of NUDAPT is to provide ancillary information, including gridded albedo,
373 vegetation coverage, population data, and anthropogenic heating (AH) for various urban
374 applications ranging from climate to human exposure modeling studies. Taha (1999), Taha and
375 Ching (2007), and Miao et al. (2009a) demonstrated that the intensity of the UHI is greatly
376 influenced by the introduction of AH, probably the most difficult data to obtain. If AH is not
377 treated as a dynamic variable (section 2.4), then it is better to treat it as a parameter rather than
378 to ignore it.

379 Anthropogenic emissions of sensible heat arise from buildings, industry/manufacturing,
380 and vehicles, and can be estimated either through inventory approaches or through direct
381 modeling. In the former approach (e.g., Sailor and Lu, 2004), aggregated consumption data are
382 typically gathered for an entire city or utility service territory, often at monthly or annual

383 resolution, and then must be mapped onto suitable spatial and temporal profiles. Waste heat
384 emissions from industrial sectors can be obtained at the state or regional level (from sources
385 such as the Federal Energy Regulatory Commission, FERC 2006), but it is difficult to assess
386 the characteristics of these facilities that would enable estimation of diurnal (sensible and
387 latent) anthropogenic flux emission profiles.

388 Regarding the transportation sector, the combustion of gasoline and diesel fuel
389 produces sensible waste heat and water vapor. Since the network of roadways is well
390 established, the transportation sector lends itself to geospatial modeling that can estimate
391 diurnal profiles of sensible and latent heating from vehicles, as illustrated by Sailor and Lu
392 (2004). A more sophisticated method incorporating mobile source emissions modeling
393 techniques is from the air quality research community.

394 Existing whole-building-energy models can estimate both the magnitude and timing of
395 energy consumption (Section 2.4). The physical characteristics of buildings, with details of the
396 mechanical equipment and building internal loads (lighting, plug loads, and occupancy), can be
397 used to estimate hourly energy usage, and hence to produce estimates of sensible and latent
398 heat emissions from the building envelope and from the mechanical heating, cooling, and
399 ventilation equipment. Correctly estimating AH relies on building size and type data spatially
400 explicit for a city. Such geospatial data are commonly available for most large cities and can
401 readily be combined with output from simulations of representative prototypical buildings
402 (Heiple and Sailor, 2008). Recently the US Department of Energy and the National Renewable
403 Energy Research Laboratory created a database of prototypical commercial buildings
404 representing the entire building stock across the US (Torcellini et al., 2008). This database
405 provides a unique opportunity to combine detailed building energy simulation with

406 Geographical Information System (GIS) data to create a US-wide resource to estimate
407 anthropogenic heat emissions from the building sector at high spatial and temporal resolutions.

408 Gridded fields of AH from NUDAPT (Ching et al. 2009), based on methodologies
409 described in Sailor and Lu (2004) and Sailor and Hart (2006), provide a good example of a
410 single product, combining waste heat from all sectors, that can be ingested into WRF/urban.
411 Inclusion of hourly gridded values of AH, along with the BEM indoor-outdoor model in
412 WRF/urban, should provide an improved base to conduct UHI mitigation studies and
413 simulations for urban planning.

414 **4.3 Model sensitivity to uncertainty in UCPs**

415 A high level of uncertainty in the specification of UCP values is inherent to the
416 methodology of aggregating fine-scale heterogeneous UCPs to the WRF modeling grid,
417 particularly to the table-based approach. It is critical to understand impacts from such
418 uncertainty on model behavior. Loridan *et al.* (2010) developed a systematic and objective
419 model response analysis procedure by coupling the offline version of SLUCM with the Multi-
420 objective Shuffled Complex Evolution Metropolis (MOSCEM) optimization algorithm of
421 Vrugt *et al.* (2003). This enables direct assessment of how a change in a parameter value
422 impacts the modeling of the surface energy balance (SEB).

423 For each UCPs in Table 1, upper and lower limits are specified. MOSCEM is set to
424 randomly sample the entire parameter space, iteratively run SLUCM, and identify values that
425 minimize the Root Mean Square Error (RMSE) of SEB fluxes relative to observations. The
426 algorithm stops when it identifies parameter values leading to an optimum compromise in the
427 performance of modeled fluxes. As an example, Fig. 10 presents the optimum values selected
428 by MOSCEM for roof albedo (α_r) when using forcing and evaluation data from a measurement

429 campaign in Marseille (Grimmond *et al.*, 2004; Lemonsu *et al.*, 2004). The algorithm is set to
430 minimize the RMSE for net all-wave radiation (Q^*) and turbulent sensible heat flux (Q_H) (two
431 objectives) using 100 samples. The optimum state identified represents a clear trade-off
432 between the two fluxes, as decreasing the value of α_r improves modeled Q^* (lower RMSE) but
433 downgrades modeled Q_H (higher RMSE). Identification of all parameters leading to such
434 trade-offs is of primary importance to understand how the model simulates the SEB, and
435 consequently how default table parameter values should be set.

436 This model-response-analysis procedure also provides a powerful tool to identify the most
437 influential UCPs, i.e., by linking the best possible improvement in RMSE for each flux to
438 corresponding parameter value changes, all inputs can be ranked in terms of their impact on the
439 modeled SEB. A complete analysis of the model response for the site of Marseille is presented
440 in Loridan *et al.* (2009). Results show that for a dense European city like Marseille, the correct
441 estimation of roof-related parameters is of critical importance, with albedo and conductivity
442 values as particularly influential. On the other hand, the impact of road characteristics appears
443 to be limited, suggesting that a higher degree of uncertainty in their estimation would not
444 significantly degrade the modeling of the SEB. This procedure, repeated for a variety of sites
445 with distinct urban characteristics (i.e., with contrasting levels of urbanization, urban
446 morphology, and climatic conditions) can provide useful guidelines for prioritizing efforts to
447 obtain urban land use characteristics for WRF.

448 **5 Evaluation of the WRF/Urban model and its recent applications**

449 The coupled WRF/Urban model has been applied to major metropolitan regions (e.g.,
450 Beijing, Guangzhou/Hong Kong, Houston, New York City, Salt Lake City, Taipei, and
451 Tokyo), and its performance was evaluated against surface observations, atmospheric

452 soundings, wind profiler data, and precipitation data (Chen et al., 2004, Holt and Pullen, 2007,
453 Miao and Chen, 2008, Lin et al., 2008, Jiang et al., 2008, Miao et al., 2009a, Miao et al.,
454 2009b, Wang et al., 2009, Kusaka et al., 2009; Tewari et al., 2010).

455 For instance, Fig. 11 shows a comparison of observed and WRF/SLUCM simulated diurnal
456 variation of 2-m temperature, surface temperatures, 10-m wind speed, and 2-m specific
457 humidity averaged over high-density urban stations in Beijing. Among the urban surface
458 temperatures, urban ground surface temperature has the largest diurnal amplitude, while wall
459 surface temperature has the smallest diurnal range, reflecting the differences in their thermal
460 conductivities and heat capacities. Results show the coupled WRF/Noah/SLUCM modeling
461 system able to reproduce the following observed features reasonably well (Miao and Chen,
462 2008, Miao et al., 2009a): 1) diurnal variation of UHI intensity; 2) spatial distribution of the
463 UHI in Beijing; 3) diurnal variation of wind speed and direction, and interactions between
464 mountain-valley circulations and the UHI; 4) small-scale boundary layer horizontal convective
465 rolls and cells; and 5) nocturnal boundary layer low-level jet.

466 Similarly, Lin et al. (2008) showed that using the WRF/Noah/SLUCM model significantly
467 improved the simulation of the UHI, boundary-layer development, and land-sea breeze in
468 northern Taiwan, when compared to observations obtained from weather stations and lidar.
469 Their sensitivity tests indicate that anthropogenic heat (AH) plays an important role in
470 boundary layer development and UHI intensity in the Taipei area, especially during nighttime
471 and early morning. For example, when AH was increased by 100 Wm^{-2} , the average surface
472 temperature increased nearly $0.3\text{-}1 \text{ }^\circ\text{C}$ in Taipei. Moreover, the intensification of the UHI
473 associated with recent urban expansion enhances the daytime sea breeze and weakens the
474 nighttime land breeze, substantially modifying the air pollution transport in northern Taiwan.

475 The WRF/urban model was used as a high-resolution regional climate model to assess the
476 uncertainty in the simulated summer UHI of Tokyo for four consecutive years (Fig. 12). When
477 the simple slab model is used in WRF, the heat island of Tokyo and of the urban area in the
478 inland northwestern part of the plain is not reproduced at all. When the WRF/Noah/SLUCM is
479 used, however, a strong nocturnal UHI is seen and warm areas are well reproduced.

480 One important goal for developing the integrated WRF/urban modeling system is to apply it
481 to understand the effects of urban expansion, so we can use such knowledge to predict and
482 assess impacts of urbanization and future climate change on our living environments and risks.
483 For instance, the Pearl River Delta (PRD) and Yangtze River Delta (YRD) regions, China,
484 have experienced a rapid, if not the most rapid in the world, economic development and
485 urbanization in the past two decades. These city clusters, centered around mega cities such as
486 Hong Kong, Guangzhou, and Shanghai (Fig. 13), have resulted in a deterioration of air quality
487 for these regions (e.g., Wang et al., 2007).

488 In a recent study by Wang et al. (2009), the online WRF Chemistry (WRF-Chem) model,
489 coupled with Noah/SLUCM and biogenic-emission models, was used to explore the influence
490 of such urban expansion. Month-long (March 2001) simulations using two land-use scenarios
491 (pre-urbanization and current) indicate that urbanization: 1) increases daily mean 2-m air
492 temperature by about 1 °C, 2) decreases 10-m wind speeds for both daytime (by 3.0 m s⁻¹) and
493 nighttime (by 0.5 to 2 m s⁻¹), and 3) increases boundary-layer depths for daytime (more than
494 200 m) and nighttime (50-100 m) periods. Changes in meteorological conditions result in an
495 increase of surface ozone concentrations by about 4.7-8.5% for nighttime and about 2.9-4.2%
496 for daytime (Fig. 14). Furthermore, despite the fact that both the PRD and the YRD have
497 similar degrees of urbanization in the last decade, and that both are located in coastal zones,

498 urbanization has different effects on the surface ozone for the PRD and the YRD, presumably
499 due to their differences in urbanization characteristics, topography, and emission source
500 strength and distribution.

501 The WRF-Chem model coupled with UCMs is equally useful to project, for instance, air
502 quality change in cities under future climate change scenarios. For example, the impact of
503 future urbanization on surface ozone in Houston under the future IPCC A1B scenario for
504 2051–2053 (Jiang et al. 2008) shows generally a 2⁰C increase in surface air temperature due to
505 the combined change in climate and urbanization. In this example, the projected 62% increase
506 of urban areas exerted more influence than attributable to climate change alone. The combined
507 effect of the two factors on O₃ concentrations can be up to 6.2 ppbv. The Jang et al. (2008)
508 sensitivity experiments revealed that future change in anthropogenic emissions produces the
509 same order of O₃ change as those induced by climate and urbanization.

510 **6 Summary and conclusions**

511 An international collaborative effort has been underway since 2003 to develop an
512 integrated, cross-scale urban modeling capability for the community WRF model. The goal is
513 not only to improve WRF weather forecasts for cities, and thereby to improve air quality
514 prediction, but also to establish a modeling tool for assessing the impacts of urbanization on
515 environmental problems by providing accurate meteorological information for planning
516 mitigation and adaptation strategies in a changing climate. The central distinction between our
517 efforts and other atmosphere-urban coupling work is the availability of multiple choices of
518 models to represent the effects of urban environments on local and regional weather and the
519 cross-scale modeling ability (ranging from continental, to city, and to building scales) in the
520 WRF/urban model. These currently include: 1) a suite of urban parameterization schemes with

521 varying degrees of complexities, 2) a capability of incorporating in-situ and remote-sensing
522 data of urban land use, building characteristics, and anthropogenic heat and moisture sources,
523 3) companion fine-scale atmospheric and urbanized land data assimilation systems, and 4)
524 ability to couple WRF/urban to fine-scale urban T&D models and with chemistry models.

525 Inclusion of three urban parameterization schemes (i.e., bulk parameterization, SLUCM,
526 and BEP) provides users with options for treating urban surface processes. Parallel to an
527 international effort to evaluate 30 urban models, executed in offline 1-D mode, against site
528 observations (Grimmond et al., 2010), work is underway within our group to evaluate three
529 WRF urban models in coupled mode against surface and boundary layer observations from the
530 Texas Air Quality Study 2000 (TexAQS2000) field program in the greater Houston area,
531 Central California Ozone Study (CCOS2000), and Southern California Ozone Study
532 (SCOS1997). Choice of specific applications will dictate careful selection of different sets of
533 science options and available databases. For instance, the bulk parameterization and SLUCM
534 may be more suitable for real-time weather and air quality forecasts than the resource-
535 demanding BEP. On the other hand, studying, for instance, the impact of air conditioning on
536 the atmosphere and in developing an adaptation strategy for planning the use of air
537 conditioning in less-developed countries in the context of intensified heat waves projected by
538 IPCC, will need to invoke the more sophisticated BEP coupled with the BEM indoor-outdoor
539 exchange model.

540 Initializing UCM state variables is a difficult problem, which has not yet received much
541 attention in the urban modeling community. Although in its early stage of development
542 (largely due to lack of appropriate data for its evaluation), u-HRLDAS may provide better
543 initial conditions for the state variables required by UCMs than the current solution that assigns

544 a uniform temperature profile for model grid points cross a city. Similarly, specification of
545 twenty-some UCPs will remain a challenge, due to the large disparity in data availability and
546 methodology for mapping fine-scale, highly variable data for the WRF modeling grid.
547 Currently the WRF pre-processor (WPS) is able to ingest: 1) high-resolution urban land-use
548 maps and to then assign UCPs based on a parameter table and 2) gridded UCPs, such as those
549 from NUDAPT (Ching et al., 2009). It would be useful to blend these two methods whenever
550 gridded UCPs are available. Bringing optimization algorithms together with UCMs and
551 observations, as recently demonstrated by Loridan *et al.* (2010), is a useful methodology to
552 identify a set of UCPs to which the performance of the UCM is most sensitive, and to
553 eventually define optimized values for those UCPs for a specific city.

554 Among these UCPs, anthropogenic heating (AH) has emerged as the most difficult
555 parameter to obtain. Methods to estimate AH from buildings, industry/manufacturing, and
556 transportation sectors have been developed (e.g., Sailor and Lu, 2004, Sailor and Hart, 2006,
557 Torcellini et al., 2008). Although data regarding the temporal and spatial distribution of waste
558 heat emissions from industry, buildings, and vehicle combustion do exist for most cities,
559 obtaining and processing these data are far from automated tasks. Nevertheless, the data
560 currently available for major US cities in NUDAPT provide examples of combining all AH
561 sources to create a single, hourly input for the WRF/urban model.

562 Evaluations and applications of this newly developed WRF/urban modeling system have
563 demonstrated its utility in studying air quality and regional climate. Preliminary results that
564 verify the performance of WRF/UCM for several major cities are encouraging (e.g., Chen et al.,
565 2004, Holt and Pullen, 2007, Miao and Chen, 2008, Lin et al., 2008, Miao et al., 2009a, Miao
566 et al., 2009b, Wang et al., 2009, Tewari et al., 2010, Kusaka et al., 2009). They show that the

567 model is generally able to capture influences of urban processes on near-surface
568 meteorological conditions and on the evolution of atmospheric boundary-layer structures in
569 cities. More importantly, recent studies (Jiang et al., 2008, Wang et al., 2009, Tewari et al.,
570 2010) have demonstrated the promising value of employing this model to investigate urban and
571 street-level plume T&D and air quality, and to predict impacts of urbanization on our living
572 environments and for risks in the context of global climate change.

573 While this WRF/urban model has been released (WRF V3.1, April 2009), except for the
574 BEM model that is in the final stages of testing, much work still remains to be done. We
575 continue to: further improve the UCMs, explore new methods of blending various data sources
576 to enhance the specification UCPs, increase the coverage of high resolution data sets,
577 particularly enhancing anthropogenic heating and moisture inputs, and link this physical
578 modeling system with, for instance, human-response models and decision support systems.

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785 **Table 1.** Urban canopy parameters currently in WRF for three urban land-use categories:
 786 low-intensity residential, high-intensity residential, and industrial and commercial. The last
 787 two columns indicate if a specific parameter is used in SLUCM and BEP, and the last three
 788 parameters are exclusively used in BEP.

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Parameter	Unit	Specific Values for			SLUCM	BEP
		Low intensity residential	High intensity residential	Industrial, commercial		
h (Building Height)	m	5	7.5	10	Yes	No
l_{roof} (Roof Width)	m	8.3	9.4	10	Yes	No
l_{road} (Road Width)	m	8.3	9.4	10	Yes	No
AH (Anthropogenic Heat)	W m^{-2}	20	50	90	Yes	No
F_{urb} (Urban fraction)	Fraction	0.5	0.9	0.95	Yes	Yes
C_{R} (Heat capacity of roof)	$\text{J m}^{-3} \text{K}^{-1}$	1.0E6	1.0E6	1.0E6	Yes	Yes
C_{W} (Heat capacity of building wall)	$\text{J m}^{-3} \text{K}^{-1}$	1.0E6	1.0E6	1.0E6	Yes	Yes
C_{G} (Heat capacity of road)	$\text{J m}^{-3} \text{K}^{-1}$	1.4E6	1.4E6	1.4E6	Yes	Yes
λ_{R} (Thermal Conductivity of roof)	$\text{J m}^{-1} \text{s}^{-1} \text{K}^{-1}$	0.67	0.67	0.67	Yes	Yes
λ_{W} (Thermal Conductivity of building wall)	$\text{J m}^{-1} \text{s}^{-1} \text{K}^{-1}$	0.67	0.67	0.67	Yes	Yes
λ_{G} (Thermal Conductivity of road)	$\text{J m}^{-1} \text{s}^{-1} \text{K}^{-1}$	0.4004	0.4004	0.4004	Yes	Yes
α_{R} (Surface Albedo of roof)	Fraction	0.20	0.20	0.20	Yes	Yes
α_{W} (Surface Albedo of building wall)	Fraction	0.20	0.20	0.20	Yes	Yes
α_{G} (Surface Albedo of road)	Fraction	0.20	0.20	0.20	Yes	Yes
ε_{R} (Surface emissivity of roof)	-	0.90	0.90	0.90	Yes	Yes

ϵ_w (Surface emissivity of building wall)	-	0.90	0.90	0.90	Yes	Yes			
ϵ_G (Surface emissivity of road)	-	0.95	0.95	0.95	Yes	Yes			
Z_{0R} (Roughness length for momentum over roof)	m	0.01	0.01	0.01	Yes*	Yes			
Z_{0W} (Roughness length for momentum over building wall)	m	0.0001	0.0001	0.0001	No*	No			
Z_{0G} (Roughness length for momentum over road)	m	0.01	0.01	0.01	No*	Yes			
b) Parameters used only in BEP									
Street Parameters		Directions from North (degrees)		Directions from North (degrees)		Directions from north (degrees)		No	Yes
		0	90	0	90	0	90		
W (Street Width)	m	15	15	15	15	15	15		
B (Building Width)	m	15	15	15	15	15	15		
h (Building Heights)	m	Height	%	Height	%	Height	%		
		5	50	10	3	5	30		
		10	50	15	7	10	40		
				20	12	15	50		
				25	18				
				30	20				
				35	18				
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				45	7				
				50	3				

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*Note: For SLUCM, if the Jurges' formulation is selected instead of Monin-Obukhov formulation (a default option in WRF V3.1), Z_{0W} and Z_{0G} are not used.

Figure Captions

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Figure 1. Overview of the integrated WRF/urban modeling system, which includes urban-modeling data-ingestion enhancements in the WRF Preprocessor System (WPS), a suite of urban modeling tools in the core physics of WRF V 3.1, and its potential applications.

Figure 2. A schematic of the single-layer UCM (SLUCM, on the left-hand side) and the multi-layer BEP models (on the right-hand side).

Figure 3. Simulated vertical profiles of nighttime temperature above a city and a rural site upwind of the city. Results obtained with WRF/BEP for a 2-D simulation (from Martilli and Schmitz, 2007).

Figure 4. Kinematic sensible heat fluxes: measured (solid line); computed offline with BEP+BEM and air conditioning working 24-h a day (ucp-bemac); with BEP+BEM and air conditioning working only from 0800 to 2000 LST (ucp-bemac*); with BEP+BEM, but without air conditioning (ucp-bem); and with the old version BEP. Results are at 18 m for a three-day period during the BUBBLE campaign (from Salamanca and Martilli, 2009).

Figure 5. Schematic representation of the coupling between the mesoscale WRF and the fine-scale urban T&D EULAG model.

Figure 6: Contours are the density of SF6 tracer gas (in parts per thousand) 60 minutes after the third release, simulated by CFD-urban using: a) single sounding observed at the Raging Waters site and (b) WRF 12-h forecast. Dots represent observed density (in same scale as in scale bar) at sites throughout the downtown area of Salt Lake City (from Tewari et al. 2010).

Figure 7. Dispersion footprint for IOP6 0900 CDT release from source located at Botanical Gardens (near Sheridan & Robinson avenues, Oklahoma City, Oklahoma) calculated with WRF/EULAG.

826 Figure 8. Noah/SLUCM simulated differences in 4th-layer road temperature (K), valid at 1200
827 UTC 23 August 2006 for Houston, Texas, between the control simulation with 20-month spin-
828 up time and a sensitivity simulation with: a) six-month, b) two-month, c) one-month, and d)
829 14-day spin-up times.

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831 Figure 9. Land use and land cover in the Greater Houston area, Texas, based on 30-m Landsat
832 from the NLCD 1992 data.

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834 Figure 10: Optimum roof albedo values (α_r) identified by MOSCEM, when considering the
835 RMSE ($W m^{-2}$) for Q^* and Q_H with forcing and evaluation data from Marseille.

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837 Figure 11. The diurnal variation of: (a) temperature ($^{\circ}C$), as observed (obs), modeled 2-m air
838 temperature (t2) and within the canyon (T2C), modeled aggregated land surface (TSK), and
839 facet temperatures for roof (TR), wall (TB) and ground (TG); (b) observed (obs) and modeled
840 10-m wind speed (wsp) and simulated wind-speed within the urban canyon in $m s^{-1}$; and (c)
841 observed (obs) and modeled (q2) 2-m specific humidity ($g kg^{-1}$). Variables were averaged over
842 high-density urban area stations for Beijing (from Miao et al., 2009a).

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844 Figure 12. Monthly mean surface air temperature at 2 m in Tokyo area at 0500 JST in August
845 averaged for 2004-2007: (a) AMeDAS observations, (b) from WRF/Slab model, and (c) from
846 WRF/SLUCM (from Kusaka et al., 2009).

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848 Figure 13. Urban land-use change in the PRD and YRD regions, China, marked in red from
849 pre-urbanization (1992-93) and current (2004): a) WRF-Chem domain with 12-km grid
850 spacing; b) 1992-1993 USGS data for PRD, c) 2004 MODIS data for YRD, d) 1992-1993
851 USGS data for PRD, and e) 2004 MODIS data for YRD (from Wang et al., 2009).

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853 Figure 14. Difference of surface ozone (in ppbv) and relative 10-m wind vectors: (a) daytime,
854 (b) nighttime (from Wang et al., 2009).