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# Evaluation of Nitrous Acid Sources and Sinks in Urban Outflow

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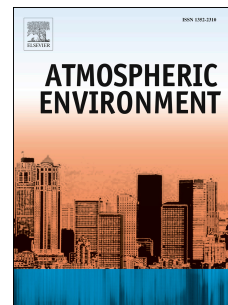
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## 1 Evaluation of nitrous acid sources and sinks in urban outflow

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## 20 ABSTRACT

21 Intensive air quality measurements made from June 22-25, 2011 in the outflow of the Dallas-Fort  
22 Worth (DFW) metropolitan area are used to evaluate nitrous acid (HONO) sources and sinks. A  
23 two-layer box model was developed to assess the ability of established and recently identified  
24 HONO sources and sinks to reproduce observations of HONO mixing ratios. A baseline model  
25 scenario includes sources and sinks established in the literature and is compared to scenarios  
26 including three recently identified sources: volatile organic compound-mediated conversion of  
27 nitric acid to HONO (S1), biotic emission from the ground (S2), and re-emission from a surface  
28 nitrite reservoir (S3). For all mechanisms, ranges of parametric values span lower- and upper-  
29 limit values. Model outcomes for ‘likely’ estimates of sources and sinks generally show under-  
30 prediction of HONO observations, implying the need to evaluate additional sources and  
31 variability in estimates of parameterizations, particularly during daylight hours. Monte Carlo  
32 simulation is applied to model scenarios constructed with sources S1-S3 added independently  
33 and in combination, generally showing improved model outcomes. Adding sources S2 and S3  
34 (scenario S2/S3) appears to best replicate observed HONO, as determined by the model  
35 coefficient of determination and residual sum of squared errors ( $r^2 = 0.55 \pm 0.03$ ,  $SSE = 4.6 \times 10^6$   
36  $\pm 7.6 \times 10^5$  ppt<sup>2</sup>). In scenario S2/S3, source S2 is shown to account for 25% and 6.7% of the  
37 nighttime and daytime budget, respectively, while source S3 accounts for 19% and 11% of the  
38 nighttime and daytime budget, respectively. However, despite improved model fit, there remains  
39 significant underestimation of daytime HONO; on average, a 0.15 ppt/s unknown daytime  
40 HONO source, or 67% of the total daytime source, is needed to bring scenario S2/S3 into  
41 agreement with observation. Estimates of ‘best fit’ parameterizations across lower to upper-limit  
42 values results in a moderate reduction of the unknown daytime source, from 0.15 to 0.10 ppt/s.

43 Keywords: air quality; unknown HONO source; Monte Carlo simulation; evolutionary solver

## 44 1. INTRODUCTION

45 Atmospheric nitrous acid (HONO) is important due to the role of HONO in generation of  
46 the hydroxyl radical (OH). There are a number of known sources of OH in the troposphere;  
47 however, OH production from HONO is of interest because the sources, fate, and diurnal cycling  
48 of HONO in the atmosphere have only recently begun to be elucidated. Models of atmospheric  
49 HONO generally employ a mass balance approach that allows evaluation of the HONO budget,  
50 often with a potentially limiting photostationary state assumption. As summarized by Spataro  
51 and Ianniello (2014) models generally include sources, sinks, and transport, the last relevant as  
52 formation processes hypothesized to occur at the ground result in vertical gradients of HONO.

53 Homogeneous and heterogeneous reactions, as well as direct emission of HONO from  
54 combustion sources, contribute to the presence of HONO in the troposphere (Finlayson-Pitts and  
55 Pitts, 1999). Nitrous acid strongly absorbs sunlight at wavelengths shorter than 390 nm resulting  
56 in photolytic degradation to OH and nitric oxide (NO). This results in suppressed, but non-zero,  
57 mixing ratios of daytime HONO due to the presence of daytime sources (Kleffmann, 2007). At  
58 night, the absence of this photolytic loss mechanism results in HONO accumulation, generally  
59 on the order of 0.1 ppb to 10 ppb (Kleffmann et al., 2003; Su et al., 2008; Young et al., 2012).  
60 The resumption of HONO photolysis after sunrise can lead to substantial formation of OH in the  
61 early morning. Alicke et al. (2003) report that during the BERLIOZ investigation at a rural,  
62 lightly trafficked site with low anthropogenic emissions during the summer months, photolysis  
63 of HONO was the dominant source of OH in the morning, and contributed as much as 20% of  
64 24-h integrated OH production.

65 Modeling studies generally show the need for an unknown daytime source to close the  
66 HONO budget (Staffelbach et al., 1997; Lee et al., 2015). A number of photochemically driven

67 homogeneous reactions have been identified or considered: e.g., the known reaction of OH and  
68 NO and the hypothesized reaction of photolytically excited nitrogen dioxide ( $\text{NO}_2$ ) and water (Li  
69 et al., 2008). The latter, however, may not proceed sufficiently rapidly or at adequate yields to  
70 affect HONO mixing ratios in the atmosphere (Carr et al., 2009). Other potential homogeneous  
71 sources are under discussion and review. For example, Li et al. (2014) proposed an internal  
72 source of HONO that consumed nitrogen oxides, although follow up discussion and further  
73 experiments indicate the source was likely strongly overestimated (Li et al., 2015; Ye et al.,  
74 2015).

75 Nitrous acid formation mediated by aerosol surface area (SA) is a topic of ongoing  
76 research, largely because the complexity of aerosols results in substantial uncertainty regarding  
77 their ultimate role in HONO formation. Static surfaces such as the ground (Stemmler et al.,  
78 2006) also may enhance HONO formation. Other hypothesized daytime sources include  
79 emissions resulting from acid/base chemistry in soils (Su et al., 2011) and photolysis of nitric  
80 acid ( $\text{HNO}_3$ ) on forest canopy surfaces (Zhou et al., 2011). Photoenhanced conversion of  $\text{NO}_2$  on  
81 organic surfaces, including the ground and aerosols, are also thought to contribute to the daytime  
82 HONO budget (George et al., 2005; Stemmler et al., 2006, 2007).

83 Given the many identified and proposed HONO source and sink mechanisms, single  
84 value estimates of parameterizations of HONO sources and sinks limit the ability to understand  
85 the impact of variability in multiple input parameters on models of HONO dynamics in the  
86 atmosphere. Monte Carlo simulation (MCS) provides a tool to observe the combined effects of  
87 ranges of input parameters and the resulting impact on the agreement between model output and  
88 measurements. In this work, we identify fourteen HONO sources or sinks established in the  
89 literature, including three sources that have recently (2013-2014) been identified. We evaluate

90 these recently identified sources through incorporation into a baseline model with a full-factorial,  
91 deterministic screening analysis. We then identify scenarios for which we stochastically  
92 parameterize source and sink mechanisms with MCS to determine probability distributions of  
93 modeled HONO mixing ratios.

## 94 **2. METHODS**

### 95 *2.1 Measurements*

96 Measurements of gas- and particle-phase constituents were made from May 30 to July 1,  
97 2011 in a semi-urban area approximately 68 km northwest of the Dallas-Fort Worth (DFW)  
98 metropolitan area. The monitoring site was co-located with the Texas Commission on  
99 Environmental Quality Eagle Mountain Lake (EML) continuous ambient monitoring station  
100 (CAMS 75). Further details regarding the geography, surrounding industrial and biogenic  
101 activities, and site conditions have been outlined previously (Rutter et al., 2015)

102 Temperature, humidity (Vaisala, HMP-45C in a RM Young 10-plate solar radiation  
103 shield), and planetary boundary layer (PBL) height (Vaisala, CL31) were measured throughout  
104 the duration of the campaign. Mixing ratios of HONO and HNO<sub>3</sub> were measured every five  
105 minutes using a method that coupled a mist chamber with ion chromatography (Dionex, CD20-  
106 1), described in greater detail elsewhere (Dibb et al., 2004). First-order photolysis rate constants  
107 (*j*-values) were determined with radiometric measurements of actinic flux determined with a 2- $\pi$   
108 double monochromator with photomultiplier and subsequent calculations following IUPAC  
109 recommendations. Nitrogen oxides were recorded every minute using a chemiluminescence trace  
110 level NO-NO<sub>2</sub>-NO<sub>x</sub> analyzer (Thermo Electron Corp., Model 42C) equipped with a Blue Light  
111 Converter (Air Quality Design, Inc.) for NO<sub>2</sub> quantification. Hydroxyl radical was observed  
112 using atmospheric pressure chemical ionization mass spectrometry (Kim et al., 2013). One-hour

113 averaged mixing ratios of volatile organic compounds (VOCs) were measured using a thermal  
 114 desorption gas chromatograph with flame ionization detection (Perkin-Elmer O<sub>3</sub> Precursor  
 115 Analyzer System). Continuous measurements of number-based particle size distributions  
 116 (diameter range of 20 nm to 500 nm) were made every ten minutes with a scanning electrical  
 117 mobility sizer (SEMS, Brechtel Inc. Model 2002) and were converted to SA distributions  
 118 assuming spherical particles. Concentrations of particulate phase nitrate were determined with an  
 119 Aerodyne high-resolution time-of-flight aerosol mass spectrometer, as described by Rutter et al.  
 120 (2015). Black carbon concentrations were measured using an aethalometer.

## 121 2.2 Baseline model

122 A two-layer box model describing HONO mixing ratios was developed, with the height  
 123 of the first layer set to 36 m to represent a surface layer and the height of layer 2 set to 72 m to  
 124 facilitate use of HONO observations above the surface layer that are available in the literature.  
 125 Established source (labeled as ‘B1-B8’ in Table 1) and sink mechanisms (labeled ‘L1-L3’ in  
 126 Table 1) are described in full in the Supporting Information (SI) (including Figures S1-S5 and  
 127 equations S1-S20). The timeframe selected for continuous modeling was 22 June 01:00 to 25  
 128 June 14:00 (all times local) based on the longest uninterrupted period during the campaign with  
 129 observations of HNO<sub>3</sub>, HONO, aerosol SA, NO<sub>2</sub>, NO, gas-phase chloride (assumed to be  
 130 hydrochloric acid, HCl), and  $j_{\text{HONO}}$ . Mixing ratios of constituents during this period were  
 131 generally typical of the broader study period. Equation 1 describes baseline sources and sinks  
 132 modeled with a transient approach:

$$\frac{d[\text{HONO}]_{\text{trans}}}{dt} = F_{B1} + F_{B2} + F_{B3} + F_{B5} + F_{B6} + F_{B7} + F_{B8} - (F_{L1} + F_{L2} + F_{L3}) - \Psi_{\text{trans}} \quad (1)$$



133 where  $[\text{HONO}]_{\text{trans}}$  is the mixing ratio of HONO from modeled transient sources and sinks (ppt),  
 134  $dt$  is the time step (s) between measurements for which observations of all constituents present in  
 135 Equation 1 were made,  $F$  represents the source or sink strength of the indicated mechanism  
 136 (ppt/s), and  $\Psi_{\text{trans}}$  is the loss (or source) of HONO from layer 1 to (or from) layer 2 due to vertical  
 137 transport (ppt/s).

138 Equation 1 describes the transient processes occurring in the model; source B4 was  
 139 incorporated into the model after accounting for transient processes as shown in Equation 2:

$$[\text{HONO}]_{\text{total}} = [\text{HONO}]_{\text{trans}} + f_{\text{emiss}} \Delta[\text{NO}_x] \quad (2)$$

140 where  $[\text{HONO}]_{\text{total}}$  is the mixing ratio of HONO at a time step resulting from transient and  
 141 instantaneous processes (ppt) and  $f_{\text{emiss}}$  is the direct HONO emission factor described in Table 1.  
 142 Equation 2 may overestimate the contribution of B4 in a box-model, as during the daytime,  
 143 HONO will rapidly photolyze prior to the measurement of emitted  $\text{NO}_x$ .

144 Vertical transport,  $\Psi_{\text{trans}}$  (ppt/s), is calculated using a first-order flux-gradient relationship  
 145 simulated with the 1D CACHE model (Bryan et al., 2012) where mass is transported by eddy  
 146 diffusion at a magnitude proportional to the eddy diffusivity for heat ( $K_h$ ), shown in equation 3:

$$\Psi_{\text{trans}} = -K_h(z, t) \frac{\partial C(z, t)}{\partial z} \frac{1}{h} \quad (3)$$

147 where  $K_h(z, t)$  is the eddy diffusivity ( $\text{m}^2/\text{s}$ ) at height  $z$  (m) and time  $t$ . As shown in equation 3,  
 148 estimates of flux are divided by  $h$ , the height of the second layer in the model (m), prior to  
 149 inclusion in equation 1.

150 Two 1D simulations during the campaign were used to derive  $K_h$ , including one  
 151 simulation for 7-9 June and one for 10-12 June. For the layers corresponding to the upper  
 152 boundary that are used in the results here,  $K_h$  is derived based on a length scale, vertical wind

153 shear, and a stability parameter (Forkel et al., 1990). It is calculated at each time step within the  
154 model, providing a diurnal cycle that is based on meteorological conditions during the campaign.

155 Observations of HONO were made at one elevation, approximately 10 m above surface,  
156 and were used to represent the HONO mixing ratio in layer 1 of the model. Equation 3 requires  
157 an estimate of the HONO mixing ratio in layer 2 to estimate the HONO gradient. Three scenarios  
158 were considered: 1) no gradient (i.e., [HONO] in layer 1 equals that in layer 2 at all times); 2) a  
159 gradient created using fractions of [HONO] presented in Vandenkoer et al. (2013), representative  
160 of a stronger nighttime gradient and a weaker daytime gradient (GrN); and 3) a gradient created  
161 from fractions of [HONO] presented in Villena et al. (2011) that is representative of a stronger  
162 daytime gradient and weaker nighttime gradient (GrD). Diurnal profiles of the three gradient  
163 conditions are shown in Figure S6 of the SI and implications of this limitation are discussed in  
164 Section 3.2.

### 165 *2.3 Parameterization and evaluation of newly identified HONO sources*

166 Three recently identified HONO source mechanisms were parameterized to assess the  
167 potential of these mechanisms (in conjunction with B1-B8 and L1-L3) to independently or  
168 jointly account for HONO mixing ratios observed in DFW. The three mechanisms, listed in  
169 Table 1 as S1, S2, S3 are incorporated into Equation 1 as additional sources of HONO.

170 Source S1 is the formation of HONO from the reduction of  $\text{HNO}_3$  to HONO mediated by  
171 VOCs emitted from motor vehicles (Rutter et al., 2014). The source strength ( $F_{S1}$ , ppt/s) was  
172 parameterized using HONO source strength and reactant mixing ratios presented in Table 1 of  
173 Rutter et al. (2014) and is shown in equation 4:

$$F_{S1} = f_{HNO_3, VOC} \left[ \frac{\left( \frac{[\text{Propylene}]}{[\text{Benzene}]} \right)_{EML}}{\left( \frac{[\text{Propylene}]}{[\text{Benzene}]} \right)_{Max, DFW}} \right] \left[ \frac{[HNO_3]_{EML}}{[HNO_3]_{Rutter}} \right] \quad (4)$$

174 where  $f_{HNO_3, VOC}$  is the observed HONO formation rate ( $\text{ppt s}^{-1}$ ) in Rutter et al. (2014), and  
 175 normalizing ratios are further described in the SI. Estimates of ‘likely’  $f_{HNO_3, VOC}$  were taken for  
 176 experiments conducted at 50% RH while ‘lower-limit’ and ‘upper-limit’ estimates were taken as  
 177 the minimum and average across experiments shown in Table 1 of Rutter et al. (2014).  
 178 Normalizing assumptions shown in equation 4 resulted in, on average, ~95% reduction of  
 179  $f_{HNO_3, VOC}$  when calculating  $F_{S1}$ . The form of the parameterization in equation 4 is speculative;  
 180 propylene is chosen as a proxy for reactive VOCs while benzene is chosen to account for dilution  
 181 that may occur as air masses move from DFW to EML (see Figure S7 in the SI for a diurnal  
 182 profile of propylene/benzene). Identification of specific reactive species participating in the  
 183 HONO formation process identified in Rutter et al. (2014) would enable improvements in  
 184 development and assessment of parameterizations of VOC-mediated conversion of  $HNO_3$  to  
 185 HONO.

186 Source S2 is HONO emissions from soil bacteria as described by Oswald et al. (2013).  
 187 Emission from the soil ( $F_{S2}$ ,  $\text{ppt/s}$ ) was assumed to mix instantaneously through the first model  
 188 layer as shown in equation 5:

$$F_{S2} = \frac{f_{soil}}{h} \Gamma_{S2} \quad (5)$$

189 where  $f_{soil}$  is the ‘‘optimum’’ HONO flux from a soil type ( $\text{molec cm}^{-2} \text{s}^{-1}$ ),  $h$  is the height of the  
 190 model layer, and  $\Gamma_{S2}$  represents the conversion factor to  $\text{ppt/s}$  prior to inclusion in equation 1 (see  
 191 the SI equations S21-S24 for an example calculation). The ‘lower-limit’ value of  $f_{soil}$  was taken

192 as the value of HONO flux for pasture, and the ‘upper-limit’ value was taken as that for  
193 grassland. No ‘likely’ value of  $f_{soil}$  was selected, as pasture and grassland were the only two  
194 relevant soil types for the DFW region. Despite specifying a ‘lower-limit’ value, this  
195 investigation may be effectively considering the high end of contribution of soil bacteria to  
196 HONO because “optimum” values of flux are used for both soil types.

197 Source S3 is the re-emission of HONO from a surface nitrite reservoir by displacement  
198 from  $\text{HNO}_3$  and  $\text{HCl}$ , as in Vandenboer et al. (2014, 2015) and shown in equation 6:

$$F_{S3} = \frac{[\text{HNO}_3] + [\text{HCl}]}{h} v_d \eta \quad (6)$$

199 where  $F_{S3}$  is the source strength of S3 ( $\text{ppt s}^{-1}$ ),  $v_d$  is the deposition velocity of  $\text{HNO}_3$  and  $\text{HCl}$ ,  
200 taken as  $1 \text{ cm s}^{-1}$ , and  $\eta$  is the displacement efficiency, ranging from 1% to 9% to 20% for  
201 ‘lower-limit’, ‘likely’, and ‘upper-limit’ values, respectively (VandenBoer et al., 2014). This  
202 parameterization was constrained by the calculation of a ‘reservoir’ of nitrite from deposited  
203 HONO, approximated from a material balance on the ground where the source of nitrite is  
204 mechanism L1 and loss is due to displacement from mechanism S3. Mechanism S3 was set to 0  
205 when the reservoir was equal to 0. As there may be additional sources of surface nitrite other  
206 than gas-phase HONO and surface nitrite accumulation over greater than diurnal time-scales,  
207 equation 6 likely represents a conservative estimate of the source strength of S3. Further  
208 description of the constraints on source S3 is given in the SI and dynamics are depicted in Figure  
209 S8, also in the SI.

#### 210 *2.4 Model calculation and assessment*

211 Nitrous acid mixing ratios were first modeled with the baseline scenario using the B and  
212 L parameterizations summarized in Table 1. The ‘likely’ parameterization incorporates HONO

213 source and sink estimations thought most representative of each mechanism, while ‘upper-limit’  
214 and ‘lower-limit’ are values that result in maximum or minimum HONO production,  
215 respectively, e.g. in the ‘upper-limit’, parameterizations of sources result in greater formation  
216 while those of sinks result in lower loss rates. Predictions of HONO mixing ratios were assessed  
217 through the residual sum of squared errors (SSE) and the coefficient of determination ( $r^2$ ), both  
218 determined from differences between modeled and measured HONO mixing ratios.

219 Model scenarios were constructed to assess the three new mechanisms (mechanism ID =  
220 S1, S2, and S3 shown in Table 1) and gradient conditions (GrN or GrD); scenarios are named  
221 according to the gradient used and sources added, e.g., GrN S2/S3 refers to a model scenario  
222 with the stronger nighttime gradient as described previously and with sources S2 and S3 added to  
223 baseline sources B1-B8 and sinks L1-L3. Sources S1-S3 were added to the baseline model in a  
224 full-factorial deterministic screening analysis (using ‘likely’ estimates of parameterizations) to  
225 identify scenarios for further analysis. Monte Carlo simulation (Crystal Ball v. 11.1.2.3, Oracle)  
226 was used to evaluate the probability of model scenarios to account for observed HONO mixing  
227 ratios. Input distributions of source and sink parameterizations were assumed to be triangular  
228 probability distributions, bounded by ‘lower-limit’ and ‘upper-limit’ values with the ‘likely’  
229 value as the most frequently occurring. Model sensitivity to the number of trial simulations was  
230 performed to ensure a trial-independent solution was achieved; all MCS were conducted with  
231 5,000 iterations. A bounded evolutionary solver was applied to the baseline model scenario and  
232 to the model scenario with the highest  $r^2$  and lowest residual SSE in the deterministic screening  
233 analysis. The evolutionary solver used a genetic algorithm to estimate source and sink  
234 parameterizations with a minimum SSE across the range of ‘lower-limit’ to ‘upper-limit’ values  
235 for each source or sink mechanism.

### 236 3. RESULTS AND DISCUSSION

#### 237 3.1 Ambient air monitoring in the outflow of DFW

238 Experimental observations of mixing ratios of ambient gases and particles input to the  
239 model are shown in Figure 1; diurnal profiles of selected constituents across the full monitoring  
240 campaign are shown in Figure S9 of the SI. Values of HONO/NO<sub>2</sub> are variable and elevated  
241 during the daytime, possibly indicative of a secondary daytime source of HONO. Mixing ratios  
242 of HNO<sub>3</sub> are suppressed in the morning and evenings and elevated during daytime hours, likely a  
243 result of strong daytime HNO<sub>3</sub> production from the reaction of NO<sub>2</sub> and OH (Aneja et al., 1994).  
244 The highest observed mixing ratios of HNO<sub>3</sub> across the full monitoring campaign are included in  
245 the model period shown in Figure 1, exceeding 5000 ppt in the early evening of June 22, 2011.  
246 Mixing ratios of HCl exhibit similar trends to those observed for HNO<sub>3</sub>. Mixing ratios of HONO  
247 show accumulation over the nighttime and suppression during the daytime, a result of the strong  
248 loss due to photolysis and convective dilution during the daytime hours. Aerosols and aerosol-  
249 phase constituents appear elevated during the nighttime hours of 6/23 and 6/24 compared to  
250 daytime concentrations, but are suppressed during the nighttime of 6/25. Across the model  
251 period, the SA of particulate matter averages 125  $\mu\text{m}^2 \text{cm}^{-3}$ , consistent with typical values across  
252 the month-long monitoring campaign (Figure S1), and ranges 22  $\mu\text{m}^2 \text{cm}^{-3}$  - 392  $\mu\text{m}^2 \text{cm}^{-3}$ .

#### 253 3.2 Baseline model

254 Mixing ratios of HONO are first calculated with the model under the baseline scenario  
255 for 'likely' estimates of parameterizations. Predicted and measured mixing ratios of HONO for  
256 the baseline scenario with three HONO gradient conditions described in Section 2.2 are shown in  
257 Figure 2. The "no gradient" condition results in substantial over-estimation of nighttime HONO  
258 mixing ratios, logical given the role of the ground surface in HONO formation processes

259 included in the baseline scenario and the first layer height of 36 m. Conversely, the GrN and GrD  
260 conditions both result in underestimation of nighttime HONO, with relatively small differences  
261 between the two conditions. A strong daytime sink, due to photolysis, results in suppression of  
262 modeled daytime mixing ratios below observation for all three gradient conditions, implying the  
263 need for daytime sources beyond those considered in the baseline scenario. The underestimation  
264 may also result from the limited vertical resolution in the two-layer box model used here and the  
265 measurement height in the lower portion of the first layer (10 m); it is likely that a continuous  
266 HONO gradient is present in the 36 m of the model first layer resulting in a lower modeled  
267 mixing ratio across the first model layer than the 10 m observation.

268 While relatively few studies report measurements of vertical gradients of HONO,  
269 available profiles generally show higher HONO mixing ratios in surface layers than aloft,  
270 indicative of ground surface HONO formation. Michoud et al. (2014) summarize several studies  
271 reporting vertical gradients, four of which show the presence of a vertical gradient (Veitel, 2002;  
272 Zhang et al., 2009; Villena et al., 2011; Wong et al., 2012) and one that does not (Häseler et al.,  
273 2009). Vandenboer et al. (2013) report high-resolution vertical profiles measured from a tower in  
274 Boulder, CO, and show the presence of both daytime and nighttime HONO gradients. Veitel et  
275 al. (2002) report that over 13 months of measurements, HONO mixing ratios were observed to  
276 decrease with height under nearly all atmospheric conditions. For the present investigation, we  
277 interpret the over-prediction of HONO mixing ratios in the nighttime for the “no gradient”  
278 condition, when convective mixing is most likely to be diminished, to indicate a HONO vertical  
279 gradient. Thus, conditions GrN or GrD better represent the vertical structure of HONO mixing  
280 ratios in the outflow of DFW. While this appears to be in agreement with the preponderance of  
281 available HONO vertical gradient measurements, a site-specific HONO gradient would clearly

282 improve the present study. Nevertheless, parameterizations here allow an estimation of the  
283 source and sink processes in the outflow of DFW and exploration of two estimates of gradients  
284 to assess model sensitivity to the HONO vertical profile. The impact of the vertical gradient and  
285 of parameterizations of established and recently identified HONO sources and sinks are further  
286 explored in Sections 3.3-3.5.

### 287 *3.3 Deterministic screening analysis*

288 A deterministic screening analysis was employed to evaluate model outcomes when  
289 sources S1-S3, acting independently or in any combination, are incorporated into the model. This  
290 full-factorial analysis, consisting of 24 possible scenarios, is conducted for only the ‘likely’  
291 parameterizations of the mechanisms, as shown in Table S1 of the SI. Full output of model runs  
292 across all gradient conditions and scenarios of parameterizations are provided in Figures S10-  
293 S12.

294 Generally, ‘likely’ estimates of parameterizations showed improved model fit compared  
295 to ‘upper-limit’ estimates, implying additional sources of HONO, rather than increased  
296 production from baseline sources result in improved model outcomes. Subsequent discussion in  
297 this section reflects ‘likely’ parameterizations. Scenarios identified for further investigation are  
298 those with a combination of low SSE and high  $r^2$ . The baseline model generally is characterized  
299 by the highest model SSE, and the addition of source mechanisms S1-S3 generally lowers SSE  
300 and increases  $r^2$ . In cases, however, the SSE is lowered while the  $r^2$  decreases (for example, from  
301 GrN Baseline to GrN S1). This is a result of improvement in model prediction for only a subset  
302 of times in the modeling period. The screening analysis identified scenario S2/S3 and scenario  
303 S1/S2/S3 as having the lowest SSE and highest  $r^2$  (SSE range:  $4.3 \times 10^6$ – $6.7 \times 10^6$ ;  $r^2$  range: 0.42-



304 0.58). These scenarios, along with baseline scenarios for comparison, are further explored with  
305 MCS and an evolutionary solver.

### 306 *3.4 Monte Carlo simulation*

307 Six model scenarios that vary the new sources and vertical gradient conditions were  
308 evaluated with MCS to incorporate uncertainty and variability in each mechanism into the  
309 model; model estimates of HONO are determined as probabilistic distributions at each model  
310 time step. Summarized output of MCS are shown in Figure 3 as hourly-averaged diurnal profiles  
311 of measured and modeled distributions of HONO mixing ratios across the model period. The  
312 MCS reinforces the conclusions that ‘baseline’ source mechanisms cannot explain observed  
313 HONO mixing ratios; in the GrN Baseline condition, 90<sup>th</sup> percentile values of model output  
314 underestimate observed HONO mixing ratios in 23 of 24 reported hours, and 75<sup>th</sup> percentile  
315 values underestimate observed HONO mixing ratios all 24 reported hours.

316 The addition of source mechanisms S2 and S3 to the model (Figure 3) results in  
317 improved agreement between the model and observations for nighttime mixing ratios of HONO  
318 for both GrN and GrD conditions. GrN S2/S3 shows 9 of the 10 hours in the 21:00-07:00  
319 nighttime period are between the 10<sup>th</sup> and 90<sup>th</sup> percentile values determined in the model. GrD  
320 S2/S3 shows improvement over the GrD Baseline condition; however, metrics of goodness of fit  
321 are lower than GrN S2/S3, and there is less improvement over baseline. This appears to be a  
322 result of sustained accumulation over the nighttime period, due to the smaller HONO nighttime  
323 vertical gradient in the GrD condition. Under both GrN and GrD conditions for scenario S2/S3,  
324 daytime mixing ratios of HONO remain substantially underpredicted as in the baseline condition.

325 The addition of all three sources (S1, S2, and S3) does not appear to resolve  
326 underprediction of the daytime HONO mixing ratio. In the GrN condition, the addition of source  
327 S1 results in a small increase in over-estimation of nighttime HONO mixing ratios and metrics of  
328 model fit worsen. In the GrD condition, there is a limited impact from the combined effect of  
329 sources S1, S2 and S3, with a modest reduction in both SSE and correlation coefficient when  
330 comparing GrD S1/S2 to GrD S1/S2/S3. Figure 3 shows GrN S2/S3 results in improved model  
331 fit compared to other scenarios, although daytime HONO remains substantially underestimated.

332 An estimation of average total and relative source and sink strength across both nighttime  
333 (21:00 – 07:00) and daytime (07:00 – 21:00) is shown in Figure 4 for GrN S2/S3. Estimates of  
334 sources and sinks are reported for ‘likely’ values of parameterizations for the indicated time  
335 period. Considerable temporal differences in the contributions of various source and sinks to the  
336 HONO budget exist. At night, HONO from NO<sub>2</sub> conversion at the ground (B7) is the major  
337 source, contributing 53% of the HONO budget. Biotic release from the ground (S2) and re-  
338 emission from the nitrite reservoir (S3) are the next two largest contributors at 25% and 19%,  
339 respectively. Nighttime HONO is slightly over-estimated; an ‘unknown’ nighttime sink of  
340 0.0016 ppt/s, or 3% of the total, is required to bring the model into agreement with observations.  
341 Major nighttime sinks are vertical transport and deposition of HONO at the ground surface,  
342 contributing 73% and 21%, respectively. These nighttime sources and sinks are in general  
343 agreement with relative estimates of mechanisms reported by Czader et al. (2012), who report  
344 71% of HONO production due to heterogeneous surface chemistry and losses due to transport  
345 and deposition of 77% and 23%, respectively, during the nighttime and pre-sunrise morning.

346 During the daytime, a missing HONO source dominates; however there are meaningful  
347 contributions to the daytime HONO budget from S3, S2, B8, B7 and B5. A missing daytime

348 source of  $0.15 \text{ ppt s}^{-1}$ , or 67% of the total HONO source budget shown in Figure 4, is needed to  
349 bring modeled and measured results into full agreement. This “missing” source is in the range of  
350 magnitudes identified in other investigations, ranging from  $0.03 - 0.3 \text{ ppt s}^{-1}$  (Su et al., 2008;  
351 Elshorbany et al., 2009; Sörgel et al., 2011; VandenBoer et al., 2013; Lee et al., 2015). Unless  
352 there is a positive artifact that depends on sunlight, a strong daytime source is needed to balance  
353 the substantial sink of HONO due to photolysis (89% of the total sink). In section 3.5, we  
354 explore the potential for ‘best fit’ estimates of parameterizations in GrN S2/S3 to close some  
355 portion of the HONO budget through optimization of parameterizations across the range of  
356 values presented in Table 1.

### 357 *3.5 Evolutionary solver and sensitivity analysis*

358 An evolutionary solver was employed to estimate the optimal combination of input  
359 values within ‘lower-limit’ to ‘upper-limit’ ranges of parameterizations and the resulting impact  
360 on the estimate of the “missing” HONO source or sink. The evolutionary solver was applied to  
361 the GrN baseline scenario and GrN S2/S3. Model outcomes with optimal estimates for GrN  
362 baseline and GrN S2/S3 are shown in Figure 5 and parameterizations are reported in Table 2.

363 Across optimization of both GrN Baseline and GrN S2/S3, the largest changes to the  
364 parameterizations relate to heterogeneous conversion of  $\text{NO}_2$  on aerosol (B1 and B2) and on the  
365 ground (B7, B8), and HONO uptake to the ground (L1). Aerosol processes increase substantially  
366 as a result of a speculative upper-limit as described in the SI; B1 was allowed to vary over 1.5  
367 orders of magnitude and B2 over 2.5 orders of magnitude based on prior modeling studies, rather  
368 than experimental estimates. However, contributions from B1 and B2 remain limited ( $< 1\%$  as  
369 can be determined from absence of B1 and B2 in Figure 4), in part a result of the two layer box-  
370 model used here that emphasizes ground-level phenomena. In both GrN Baseline and GrN

371 S2/S3, the optimization resulted in B8 at the upper-limit of the parameterization. Source B7  
372 increased by  $\sim 2\times$  in GrN Baseline, but more moderately in GrN S2/S3, a result of the  
373 contribution of sources S2 and S3 in GrN S2/S3. In GrN S2/S3, deposition loss (L1) increased, a  
374 result of the need to balance increases in parameterizations of sources that act over both daytime  
375 and nighttime periods (e.g., S3) and contribute to reductions in the daytime “unknown” source  
376 but also nighttime accumulation.

377 Figure 5 shows greater improvements in metrics of model goodness of fit for the optimal  
378 solution of GrN S2/S3 compared to the optimal solutions of the GrN Baseline. This indicates that  
379 baseline mechanisms are not able to similarly explain HONO observations under any  
380 combination of input parameters compared to the scenario with S2/S3 present. This appears to  
381 largely result from stronger parameterizations of S2/S3 resulting in improved estimates of  
382 daytime HONO mixing ratio, although levels are still lower than observed. Best-fit  
383 parameterizations of GrN S2/S3 result in a missing daytime source of 0.10 ppt/s, reduced from  
384 0.15 ppt/s (Figure 4), implying that a substantial missing HONO source remains even across a  
385 statistically optimized range of parameterizations.

386 The “best-fit” estimates of GrN S2/S3 reflect an improved statistical outcome for the  
387 model when parameterizations are allowed to vary across a range of values. Parameterizations in  
388 Table 2 with larger percentage changes imply a combination of model sensitivity to the  
389 parameter as well as uncertainty in the value of the parameterization. We conducted a sensitivity  
390 analysis to identify the most important parametrizations impacting the estimates of goodness-of-  
391 fit, the model  $r^2$  and SSE. The sensitivity analysis for GrN S2/S3 is summarized in Table S2 of  
392 the SI, reported as the Spearman’s rank correlation coefficient ( $\rho$ ) between each mechanism’s  
393 input parameter and the model output  $r^2$  or SSE. Uptake of  $\text{NO}_2$  at the ground (B7) is the

394 parameter with the largest impact on both the model SSE and  $r^2$ , by a comparatively large  
395 margin. Given that there is a wide range of estimates of the uptake coefficient parameterizing B7  
396 in the literature, this source represents a large source of uncertainty in the model. Sources S3, B8,  
397 and S2 are the next three strongest correlations with model SSE; interestingly, all four sources  
398 with highest sensitivity (B7, B8, S2, and S3) are ground-level phenomena. Source B7 was  
399 strongest correlated with night-time (21:00-07:00) HONO mixing ratios while source S3 was  
400 strongest correlated with daytime HONO. This underscores the importance of characterizing the  
401 role of the ground surface mechanisms, including biotic release and ground-level chemical  
402 transformations.

403         The presence of a substantial missing daytime source is further explored via estimation of  
404 correlation coefficients between measured constituents and products of constituents with the  
405 missing HONO source, similar to the analysis presented by Lee et al. (2015). This analysis  
406 employed time-series measurements for constituents and the estimate of missing HONO at each  
407 time step required for model agreement with observation. Outcomes are shown in Table S3 for  
408 ‘likely’ and ‘best-fit’ estimates of GrN S2/S3. Relatively strong correlation coefficients ( $r^2 > 0.5$ )  
409 were observed for  $j_{\text{NO}_2}$  and  $j_{\text{NO}_2} \times \text{temperature}$  with the missing HONO source, the latter in close  
410 agreement to the results of Lee et al (2015). However, the correlation of  $j_{\text{NO}_2} \times \text{NO}_2$  with the  
411 missing HONO source is weak ( $r^2 = 0.09 - 0.17$ ), as is the correlation of  $j_{\text{NO}_2} \times \text{SEMS SA} \times \text{NO}_2$   
412 ( $r^2 = 0.08 - 0.16$ ) and with  $\text{NO}_2$  alone ( $r^2 = 0.21-0.25$ ). The stronger correlation with  $j_{\text{NO}_2}$  and  
413  $j_{\text{NO}_2} \times \text{temperature}$  may imply photosensitized conversion on organics, including humic acids,  
414 which are mainly ground surface sources (Stemmler et al., 2006, 2007), are underestimated. The  
415 weak correlation of the missing HONO source with  $\text{NO}_2$  and products containing  $\text{NO}_2$  mixing  
416 ratios appears aligned with a recent analysis of weekday-weekend HONO and  $\text{NO}_2$  relationships

417 that shows HONO production rates do not increase with increases in  $\text{NO}_2$ , implying daytime  
418 HONO production may not be rate-limited by  $\text{NO}_2$  (Pusede et al., 2015). Weakening correlations  
419 for products of gas- and particle-phase constituents and  $j_{\text{NO}_2}$  also may result from the two-layer  
420 model that lends greater emphasis to interactions at the ground level, consistent with the results  
421 of the sensitivity analysis in Table S2 and discussed previously.

### 422 3.6 Model limitations

423 The model described in this work is subject to a number of important limitations. Source  
424 S1 assumes the source strength determined in the laboratory is possible in the ambient  
425 environment, with several normalizing assumptions. However, as we did not observe meaningful  
426 formation of HONO from source S1, the impact of the speculative parameterization is therefore  
427 limited in this investigation. Future field efforts should further investigate the potential for VOC-  
428 mediated reduction of  $\text{HNO}_3$  to HONO in near-source environments. Source S2 was  
429 parameterized using a single value for a model simulation; there are likely to be diurnal  
430 variations in biological activity and soil water content that would impact the parameterization of  
431 source S2. Source S3 considered only gas-phase HONO as an input to the surface nitrite  
432 reservoir and that the reservoir was empty at the beginning of the model period. This may result  
433 in a conservative estimate of the contribution of source S3.

434 Input distributions in MCS were assumed to be triangular. This assumption may over-  
435 weight estimates of parameterizations at the ‘upper-limit’ and ‘lower-limit’ extents of the  
436 distribution as compared to a normal distribution. A triangular distribution was chosen, in part, to  
437 ensure parameterizations did not exceed upper or lower-limit estimates in MCS. The two-layer  
438 box model uses instantaneous and *in-situ* mixing ratios to constrain the model, with the  
439 assumption of instantaneous mixing up to the first layer height. Transport between layers was

440 estimated using an approximation of HONO vertical gradients at similar heights taken from  
441 literature. We assume transport time for NO<sub>x</sub> sources that exceeds the atmospheric age of HONO  
442 (Lee et al. 2013). During the daytime periods (07:00-21:00), the atmospheric age of HONO  
443 across the modeling period in this work averaged 19.4 min and ranged from 8.9 to 128 min. We  
444 assume NO<sub>x</sub> sources input to the model originate from the metropolitan DFW area (~70 km  
445 away), while the wind speed averaged 19 km/h, resulting in a transport time of 220 min.

#### 446 **4. CONCLUSIONS**

447 Model predictions of HONO that account for ranges in parameterizations of HONO  
448 source and sink mechanisms enable a statistical assessment of the likelihood of the model to  
449 match observation. Observations of HONO appear most accurately simulated when emission  
450 from soil biota (S2) and re-emission from a ground level nitrite source (S3) are included in the  
451 model. Model output for GrN S2/S3 accounted for, on average, 33% of the daytime HONO  
452 budget and 103% of the nighttime HONO budget. Major nighttime sources included (in order)  
453 NO<sub>2</sub> conversion at the ground (B7), biotic release from soil (S2), and re-emission from the nitrite  
454 reservoir (S3). Major daytime sources include S3, S2, photoenhanced NO<sub>2</sub> conversion at the  
455 ground (B8), B7, and the reaction of OH with NO (B5). Model fit improved after application of  
456 an evolutionary solver, resulting in a reduction of the estimate of the unknown daytime source  
457 for GrN S2/S3. However, the presence of a substantial unknown daytime source (on average  
458 0.10 ppt/s) even with a statistically optimal fit for GrN S2/S3 implies additional sources of  
459 HONO than those evaluated here must be included to reproduce accurately daytime HONO  
460 mixing ratios. Analyses of model sensitivity and correlations between the missing HONO source  
461 and constituents imply the presence of additional, or underestimation of considered, ground-level  
462 HONO sources in this investigation.

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467 **REFERENCES**

- 468 Aliche, B., Geyer, A., Hofzumahaus, A., Holland, F., Konrad, S., Pätz, H-W. et al., OH formation by  
469 HONO photolysis during the BERLIOZ experiment. *J. Geophys. Res. Atmos.*, 108 (D4), 8247.
- 470 Aneja, V.P., Claiborn, C.S., Li, Z., Murthy, A., 1994. Trends, seasonal variations, and analysis of high-  
471 elevation surface nitric acid, ozone, and hydrogen peroxide. *Atmospheric Environment* 28, 1781–  
472 1790.
- 473 Aumont, B., Chervier, F., Laval, S., 2003. Contribution of HONO sources to the NO<sub>x</sub>/HO<sub>x</sub>/O<sub>3</sub> chemistry  
474 in the polluted boundary layer. *Atmospheric Environment* 37, 487–498.
- 475 Bryan, A.M., Bertman, S.B., Carroll, M.A., Dusanter, S., Edwards, G.D., Forkel, R. et al., 2012. In-  
476 canopy gas-phase chemistry during CABINEX 2009: sensitivity of a 1-D canopy model to  
477 vertical mixing and isoprene chemistry. *Atmos. Chem. Phys.* 12, 8829–8849.
- 478 Carr, S., Heard, D.E., Blitz, M.A., 2009. Comment on “Atmospheric Hydroxyl Radical Production from  
479 Electronically Excited NO<sub>2</sub> and H<sub>2</sub>O.” *Science* 324, 336b.
- 480 Dibb, J.E., Scheuer, E., Whitlow, S.I., Vozella, M., Williams, E., Lerner, B.M., 2004. Ship-based nitric  
481 acid measurements in the Gulf of Maine during New England Air Quality Study 2002. *J.*  
482 *Geophys. Res. Atmos.* 109, D20303.
- 483 Elshorbany, Y.F., Kurtenbach, R., Wiesen, P., Lissi, E., Rubio, M., Villena, G. et al., 2009. Oxidation  
484 capacity of the city air of Santiago, Chile. *Atmospheric Chemistry and Physics* 9, 2257–2273.
- 485 Finlayson-Pitts, B.J., Pitts, J., 1999. *Chemistry of the upper and lower atmosphere: Theory, experiments,*  
486 *and applications.* Academic Press.
- 487 Forkel, R., Seidl, W., Dlugi, R., Deigele, E., 1990. A one-dimensional numerical model to simulate  
488 formation and balance of sulfate during radiation fog events. *J. Geophys. Res.* 95, 18501–18515.
- 489 George, C., Strekowski, R.S., Kleffmann, J., Stemmler, K., Ammann, M., 2005. Photoenhanced uptake of  
490 gaseous NO<sub>2</sub> on solid organic compounds: a photochemical source of HONO? *Faraday Discuss.*  
491 130, 195–210; discussion 241–264, 519–524.
- 492 Häsel, R., Brauers, T., Holland, F., Wahner, A., 2009. Development and application of a new mobile  
493 LOPAP instrument for the measurement of HONO altitude profiles in the planetary boundary  
494 layer. *Atmos. Meas. Tech. Discuss.* 2, 2027–2054.
- 495 Kim, S., Wolfe, G.M., Mauldin, L., Cantrell, C., Guenther, A., Karl, T. et al., 2013. Evaluation of HO<sub>x</sub>  
496 sources and cycling using measurement-constrained model calculations in a 2-methyl-3-butene-2-  
497 ol (MBO) and monoterpene (MT) dominated ecosystem. *Atmos. Chem. Phys.* 13, 2031–2044.
- 498 Kirchstetter, T.W., Harley, R.A., Littlejohn, D., 1996. Measurement of nitrous acid in motor vehicle  
499 exhaust. *Environ. Sci. Technol.* 30, 2843–2849.
- 500 Kleffmann, J., 2007. Daytime sources of nitrous acid (HONO) in the atmospheric boundary layer.  
501 *Chemphyschem* 8, 1137–1144.
- 502 Kleffmann, J., Becker, K.H., Wiesen, P., 1998. Heterogeneous NO<sub>2</sub> conversion processes on acid  
503 surfaces: possible atmospheric implications. *Atmospheric Environment* 32, 2721–2729.
- 504 Kleffmann, J., Kurtenbach, R., Lörzer, J., Wiesen, P., Kalthoff, N. et al., 2003. Measured and simulated  
505 vertical profiles of nitrous acid—Part I: Field measurements. *Atmospheric Environment* 37,  
506 2949–2955.



- 507 Kurtenbach, R., Becker, K.H., Gomes, J.A.G., Kleffmann, J., Lörzer, J.C., Spittler, M. et al., 2001.  
508 Investigations of emissions and heterogeneous formation of HONO in a road traffic tunnel.  
509 Atmospheric Environment 35, 3385–3394.
- 510 Lee, J.D., Whalley, L.K., Heard, D.E., Stone, D., Dunmore, R.E., Hamilton, J.F., Young, D.E., Allan,  
511 J.D., Laufs, S., Kleffmann, J., 2015. Detailed budget analysis of HONO in central London reveals  
512 a missing daytime source. Atmos. Chem. Phys. Discuss. 15, 22097–22139.
- 513 Lee, B.H., Wood, E.C., Herndon, S.C., Lefer, B.L., Luke, W.T., Brune, W.H. et al., 2013. Urban  
514 measurements of atmospheric nitrous acid: A caveat on the interpretation of the HONO  
515 photostationary state. J. Geophys. Res. Atmos. 118, 12274–12281.
- 516 Li, S., Matthews, J., Sinha, A., 2008. Atmospheric hydroxyl radical production from electronically  
517 excited NO<sub>2</sub> and H<sub>2</sub>O. Science 319, 1657–1660.
- 518 Li, X., Rohrer, F., Hofzumahaus, A., Brauers, T., Häsel, R., Bohn, B. et al., 2014. Missing Gas-Phase  
519 Source of HONO Inferred from Zeppelin Measurements in the Troposphere. Science 344, 292–  
520 296.
- 521 Li, X., Rohrer, F., Hofzumahaus, A., Brauers, T., Häsel, R., Bohn, B. et al., 2015. Response to  
522 Comment on “Missing gas-phase source of HONO inferred from Zeppelin measurements in the  
523 troposphere.” Science 348, 1326e.
- 524 Michoud, V., Colomb, A., Borbon, A., Miet, K., Beekmann, M., Camredon, M., et al. 2014. Study of the  
525 unknown HONO daytime source at a European suburban site during the MEGAPOLI summer  
526 and winter field campaigns. Atmos. Chem. Phys. 14, 2805–2822.
- 527 Monge, M.E., D’Anna, B., Mazri, L., Giroir-Fendler, A., Ammann, M., Donaldson, D.J., George, C.,  
528 2010. Light changes the atmospheric reactivity of soot. PNAS 107, 6605–6609.
- 529 NASA, 2011. Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies (No. JPL 10-  
530 6). California Institute of Technology, Pasadena, CA, Jet Propulsion Laboratory.
- 531 Oswald, R., Behrendt, T., Ermel, M., Wu, D., Su, H., Cheng, Y. et al., 2013. HONO emissions from soil  
532 bacteria as a major source of atmospheric reactive nitrogen. Science 341, 1233–1235.
- 533 Pusede, S.E., VandenBoer, T.C., Murphy, J.G., Marković, M.Z., Young, C.J., Veres, P.R. et al., 2015. An  
534 Atmospheric Constraint on the NO<sub>2</sub> Dependence of Daytime Near-Surface Nitrous Acid  
535 (HONO). Environ. Sci. Technol. doi:10.1021/acs.est.5b02511
- 536 Rutter, A.P., Griffin, R.J., Cevik, B.K., Shakya, K.M., Gong, L., Kim, S., Flynn, J.H., Lefer, B.L., 2015.  
537 Sources of air pollution in a region of oil and gas exploration downwind of a large city.  
538 Atmospheric Environment. 120, 89–99.
- 539 Rutter, A.P., Malloy, Q.G.J., Leong, Y.J., Gutierrez, C.V., Calzada, M., Scheuer, E., Dibb, J.E., Griffin,  
540 R.J., 2014. The reduction of HNO<sub>3</sub> by volatile organic compounds emitted by motor vehicles.  
541 Atmospheric Environment 87, 200–206.
- 542 Sörgel, M., Regelin, E., Bozem, H., Diesch, J.-M., Drewnick, F., Fischer, H. et al., 2011. Quantification  
543 of the unknown HONO daytime source and its relation to NO<sub>2</sub>. Atmospheric Chem. Phys. 11,  
544 10433–10447.
- 545 Spataro, F., Ianniello, A., 2014. Sources of atmospheric nitrous acid: State of the science, current research  
546 needs, and future prospects. J. Air Waste Manage. Assoc. 64, 1232–1250.
- 547 Staffelbach, T., Neftel, A., Horowitz, L.W., 1997. Photochemical oxidant formation over southern  
548 Switzerland, 2. Model results. Journal of Geophysical Research 102, 23363–23373.
- 549 Stemmler, K., Ammann, M., Donders, C., Kleffmann, J., George, C., 2006. Photosensitized reduction of  
550 nitrogen dioxide on humic acid as a source of nitrous acid. Nature 440, 195–198.
- 551 Stemmler, K., Ndour, M., Elshorbany, Y., Kleffmann, J., D’Anna, B., George, C., et al., 2007. Light  
552 induced conversion of nitrogen dioxide into nitrous acid on submicron humic acid aerosol.  
553 Atmospheric Chem. Phys. 7, 4237–4248.
- 554 Su, H., Cheng, Y.F., Cheng, P., Zhang, Y.H., Dong, S., Zeng, L.M. et al, 2008. Observation of nighttime  
555 nitrous acid (HONO) formation at a non-urban site during PRIDE-PRD2004 in China.  
556 Atmospheric Environment 42, 6219–6232.

- 557 Su, H., Cheng, Y., Oswald, R., Behrendt, T., Trebs, I., Meixner, F.X., et al., 2011. Soil nitrite as a source  
558 of atmospheric HONO and OH radicals. *Science* 333, 1616–1618.
- 559 Trick, S., 2004. Formation of Nitrous Acid on Urban Surfaces - A physical-chemical Perspective.  
560 Dissertation. Universität Heidelberg, 2004, <http://www.ub.uni-heidelberg.de/archiv/4814>.
- 561 VandenBoer, T.C., Brown, S.S., Murphy, J.G., Keene, W.C., Young, C.J., Pszenny, A. A. P. et al.,  
562 Roberts, J.M., 2013. Understanding the role of the ground surface in HONO vertical structure:  
563 High resolution vertical profiles during NACHTT-11. *J. Geophys. Res. Atmos.* 118, 10155–  
564 10171.
- 565 VandenBoer, T.C., Markovic, M.Z., Sanders, J.E., Ren, X., Pusede, S.E., Browne, E.C. et al., 2014.  
566 Evidence for a nitrous acid (HONO) reservoir at the ground surface in Bakersfield, CA, during  
567 CalNex 2010. *J. Geophys. Res. Atmos.* 119, 9093–9106.
- 568 VandenBoer, T.C., Young, C.J., Talukdar, R.K., Markovic, M.Z., Brown, S.S., Roberts, J.M., Murphy,  
569 J.G., 2015. Nocturnal loss and daytime source of nitrous acid through reactive uptake and  
570 displacement. *Nature Geosci.* 8, 55–60.
- 571 Veitel, H., 2002. Vertical Profiles of NO<sub>2</sub> and HONO in the Planetary Boundary Layer [WWW  
572 Document]. URL <http://archiv.ub.uni-heidelberg.de/volltextserver/2490/> (accessed 10.6.15).
- 573 Villena, G., Kleffmann, J., Kurtenbach, R., Wiesen, P., Lissi, E., Rubio, M.A. et al., 2011. Vertical  
574 gradients of HONO, NO<sub>x</sub> and O<sub>3</sub> in Santiago de Chile. *Atmospheric Environment* 45, 3867–  
575 3873.
- 576 Walcek, C.J., Brost, R.A., Chang, J.S., Wesely, M.L., 1986. SO<sub>2</sub>, sulfate and HNO<sub>3</sub> deposition velocities  
577 computed using regional landuse and meteorological data. *Atmospheric Environment* 20, 949–  
578 964.
- 579 Wong, K.W., Tsai, C., Lefer, B., Grossberg, N., Stutz, J., 2013. Modeling of daytime HONO vertical  
580 gradients during SHARP 2009. *Atmos. Chem. Phys.* 13, 3587–3601.
- 581 Wong, K.W., Tsai, C., Lefer, B., Haman, C., Grossberg, N., Brune, W.H. et al., 2012. Daytime HONO  
582 vertical gradients during SHARP 2009 in Houston, TX. *Atmos. Chem. Phys.* 12, 635–652.
- 583 Ye, C., Zhou, X., Pu, D., Stutz, J., Festa, J., Spolaor, M. et al., 2015. Comment on “Missing gas-phase  
584 source of HONO inferred from Zeppelin measurements in the troposphere.” *Science* 348, 1326-d.
- 585 Young, C.J., Washenfelder, R.A., Roberts, J.M., Mielke, L.H., Osthoff, H.D., Tsai, C. et al., 2012.  
586 Vertically resolved measurements of nighttime radical reservoirs in Los Angeles and their  
587 contribution to the urban radical budget. *Environ. Sci. Technol.* 46, 10965–10973.
- 588 Zhang, N., Zhou, X., Shepson, P.B., Gao, H., Alaghmand, M., Stirn, B., 2009. Aircraft measurement of  
589 HONO vertical profiles over a forested region. *Geophys. Res. Lett.* 36, L15820.
- 590 Zhou, X., Gao, H., He, Y., Huang, G., Bertman, S.B., Civerolo, K., Schwab, J., 2003. Nitric acid  
591 photolysis on surfaces in low-NO<sub>x</sub> environments: Significant atmospheric implications.  
592 *Geophysical Research Letters* 30(23), 2217.
- 593 Zhou, X., Zhang, N., TerAvest, M., Tang, D., Hou, J., Bertman, S. et al., 2011. Nitric acid photolysis on  
594 forest canopy surface as a source for tropospheric nitrous acid. *Nature Geosci.* 4, 440–443.

595 Table 1. HONO source and sink mechanisms considered for modeling HONO in the outflow of the DFW metropolitan area.

Mechanism	ID	Parameter	Lower-limit	Likely	Upper-limit	Reference
Aerosol uptake of NO <sub>2</sub>	B1	$\gamma_{\text{NO}_2}$ (-)	$2.0 \times 10^{-7}$	$1.0 \times 10^{-6}$	$5.0 \times 10^{-6}$	Kleffmann et al. (1998); Aumont et al.(2003)
Photoenhanced aerosol uptake of NO <sub>2</sub>	B2	$\gamma_{\text{NO}_2,\text{hv}}$ (-)	$4.0 \times 10^{-6}$	$1.0 \times 10^{-5}$	$1.0 \times 10^{-3}$	Stemmler et al. (2007); Wong et al. (2013)
Photoenhanced conversion of NO <sub>2</sub> soot	B3	$\gamma_{\text{soot,BET}}$ (-)	$4.0 \times 10^{-7}$	$5.0 \times 10^{-7}$	$6.0 \times 10^{-7}$	Monge et al. (2010)
		BET surface area (cm <sup>2</sup> /g)	$9.7 \times 10^5$	$1.2 \times 10^6$	$1.3 \times 10^6$	
Direct HONO emission	B4	$f_{\text{emiss}}$ (% V, $\Delta\text{HONO}/\Delta\text{NO}_x$ )	0.0029	0.0055	0.0080	Kirchstetter et al. (1996); Kurtenbach et al. (2001)
OH + NO	B5	$k_{\infty}(\text{T})$ (cm <sup>3</sup> molec <sup>-1</sup> s <sup>-1</sup> )	$3.0 \times 10^{-11}$	$3.6 \times 10^{-11}$	$4.3 \times 10^{-11}$	NASA (2011)
		$k_o(\text{T})$ (cm <sup>6</sup> molec <sup>-2</sup> s <sup>-1</sup> )	$5.8 \times 10^{-31}$	$7.0 \times 10^{-31}$	$8.4 \times 10^{-31}$	NASA (2011)
HONO from surface HNO <sub>3</sub> photolysis	B6	$j_{\text{HNO}_3\text{-HONO}}$ (s <sup>-1</sup> )	$1.0 \times 10^{-5}$	$1.2 \times 10^{-5}$	$1.4 \times 10^{-5}$	Zhou et al. (2003)
		$v_{d,\text{HNO}_3}$ (cm s <sup>-1</sup> )	1.50	1.75	2.25	Walcek et al. (1986)
HONO from NO <sub>2</sub> conversion at ground	B7	$\gamma_{\text{NO}_2,\text{gr}}$ (-)	$1.0 \times 10^{-6}$	$5.0 \times 10^{-6}$	$1.0 \times 10^{-5}$	Kleffmann et al. (1998); Kurtenbach et al. (2001)
Photoenhanced NO <sub>2</sub> conversion, ground	B8	$\gamma_{\text{NO}_2,\text{gr,hv}}$ (-)	$1.7 \times 10^{-5}$	$2.0 \times 10^{-5}$	$6.0 \times 10^{-5}$	Stemmler et al. (2006); Wong et al. (2013)
HNO <sub>3</sub> → HONO, VOC	S1	$f_{\text{HNO}_3,\text{VOC}}$ (ppt s <sup>-1</sup> )	$3.6 \times 10^{-2}$	$5.8 \times 10^{-2}$	$8.3 \times 10^{-2}$	Rutter et al. (2014)
Biotic release, ground	S2	$f_{\text{soil}}$ (molec cm <sup>-2</sup> s <sup>-1</sup> )	-	$1.7 \times 10^9$	$4.0 \times 10^9$	Oswald et al. (2013)
Re-emission from NO <sub>2</sub> -(p) reservoir	S3	$v_d \times \eta$ (cm s <sup>-1</sup> )	$1.0 \times 10^{-2}$	$9.0 \times 10^{-2}$	$2.0 \times 10^{-1}$	Vandenboer et al. (2014)
HONO uptake at ground	L1	$\gamma_{\text{HONO,gr}}$ (-)	$1.0 \times 10^{-4}$	$2.0 \times 10^{-5}$	$1.8 \times 10^{-5}$	Vandenboer et al. (2013); Wong et al. (2013); Trick (2004)
HONO + OH	L2	$k_{\text{HONO+OH}}$ (cm <sup>3</sup> molec <sup>-1</sup> s <sup>-1</sup> )	$6.75 \times 10^{-12}$	$4.5 \times 10^{-12}$	$3.0 \times 10^{-12}$	NASA (2011)
HONO photolysis	L3	$j_{\text{HONO}}$ (s <sup>-1</sup> )	$1.8 \times 10^{-3} - 3.9 \times 10^{-5}^{\text{a}}$			This investigation

596 <sup>a</sup>Maximum-minimum range of the experimentally determined time-series values of  $j_{\text{HONO}}$  input to the model (not varied).

597 Table 2. Best estimates of parameterizations of sources and sinks of HONO in the outflow of  
 598 DFW for baseline and scenario GrN S2/S3.

ID	Parameter	Best-fit estimate (% difference from 'likely')	
		GrN S2, S3	GrN Baseline
B1	$\gamma_{\text{NO}_2}$ (-)	$3.9 \times 10^{-6}$ (294%)	$2.5 \times 10^{-6}$ (152%)
B2	$\gamma_{\text{NO}_2, \text{hv}}$ (-)	$8.5 \times 10^{-4}$ (8500%)	$1.0 \times 10^{-3}$ (9900%)
B3	$\gamma_{\text{soot, BET}}$ (-)	$5.3 \times 10^{-7}$ (6%)	$5.3 \times 10^{-7}$ (7.1%)
	BET surface area ( $\text{cm}^2/\text{g}$ )	$1.1 \times 10^2$ (-6.5%)	$1.2 \times 10^2$ (-3%)
B4	$f_{\text{emiss}}$ (%v, $\Delta\text{HONO}/\Delta\text{NO}_2$ )	0.0043 (-22%)	0.0049 (-10%)
B5	$k_{\infty}(\text{T})$ ( $\text{cm}^3 \text{ molec}^{-1} \text{ s}^{-1}$ )	$3.7 \times 10^{-11}$ (4.4%)	$3.8 \times 10^{-11}$ (4.8%)
	$k_{\text{o}}(\text{T})$ ( $\text{cm}^6 \text{ molec}^{-2} \text{ s}^{-1}$ )	$7.6 \times 10^{-31}$ (9%)	$7.3 \times 10^{-31}$ (4.8%)
B6	$j_{\text{HNO}_3\text{-HONO}}$ ( $\text{s}^{-1}$ )	$1.2 \times 10^{-5}$ (-3%)	$1.3 \times 10^{-5}$ (7.7%)
	$v_{\text{d, HNO}_3}$ ( $\text{cm s}^{-1}$ )	1.8 (4.6%)	2.0 (17%)
B7	$\gamma_{\text{NO}_2, \text{gr}}$ (-)	$6.1 \times 10^{-6}$ (22%)	$9.9 \times 10^{-6}$ (97%)
B8	$\gamma_{\text{NO}_2, \text{gr, hv}}$ (-)	$6 \times 10^{-5}$ (200%)	$6 \times 10^{-5}$ (200%)
S1	$f_{\text{HNO}_3, \text{VOC}}$ ( $\text{ppt s}^{-1}$ )	n/a	n/a
S2	$f_{\text{soil}}$ ( $\text{molec cm}^{-2} \text{ s}^{-1}$ )	$2.8 \times 10^9$ (66%)	n/a
S3	$v_{\text{d}} \times \eta$ ( $\text{cm s}^{-1}$ )	0.18 (105%)	n/a
L1	$\gamma_{\text{HONO, gr}}$ (-)	$5.7 \times 10^{-5}$ (185%)	$2.0 \times 10^{-5}$ (-1.1%)
L2	$k_{\text{HONO+OH}}$ ( $\text{cm}^3 \text{ molec}^{-1} \text{ s}^{-1}$ )	$5.7 \times 10^{-12}$ (28%)	$4.6 \times 10^{-12}$ (2.1%)
L3	$j_{\text{HONO}}$ ( $\text{s}^{-1}$ )	unchanged	unchanged
	Missing source or sink: daytime, nighttime ( $\text{ppt s}^{-1}$ )	0.10, -0.0112	0.15, -0.006

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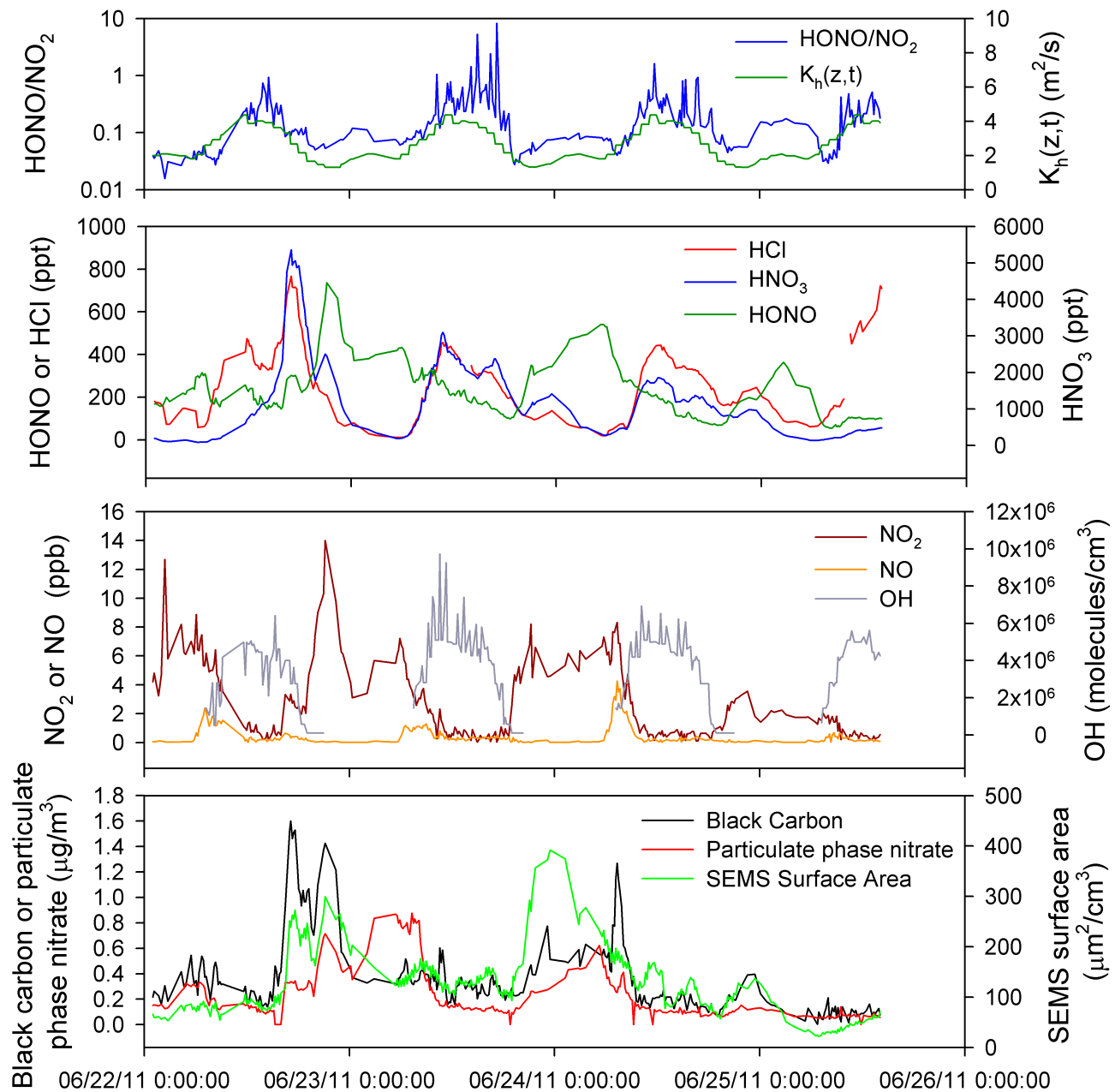


Figure 1. Time series inputs to the two-layer box model of HONO mixing ratios in the outflow of DFW.

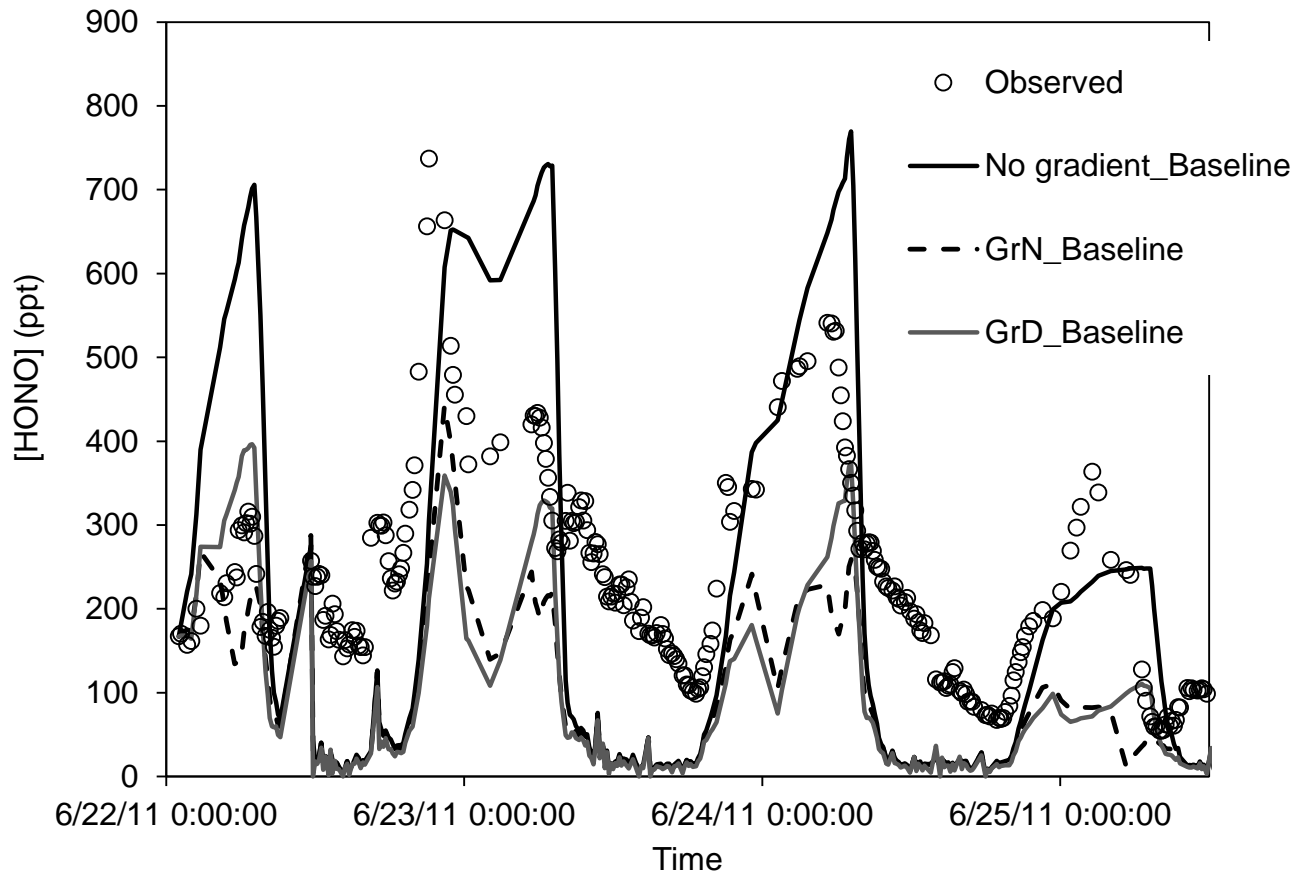


Figure 2. Model output for 'likely' estimates of parameterizations under conditions of no gradient, stronger nighttime gradient (GrN), and stronger daytime gradient (GrD).

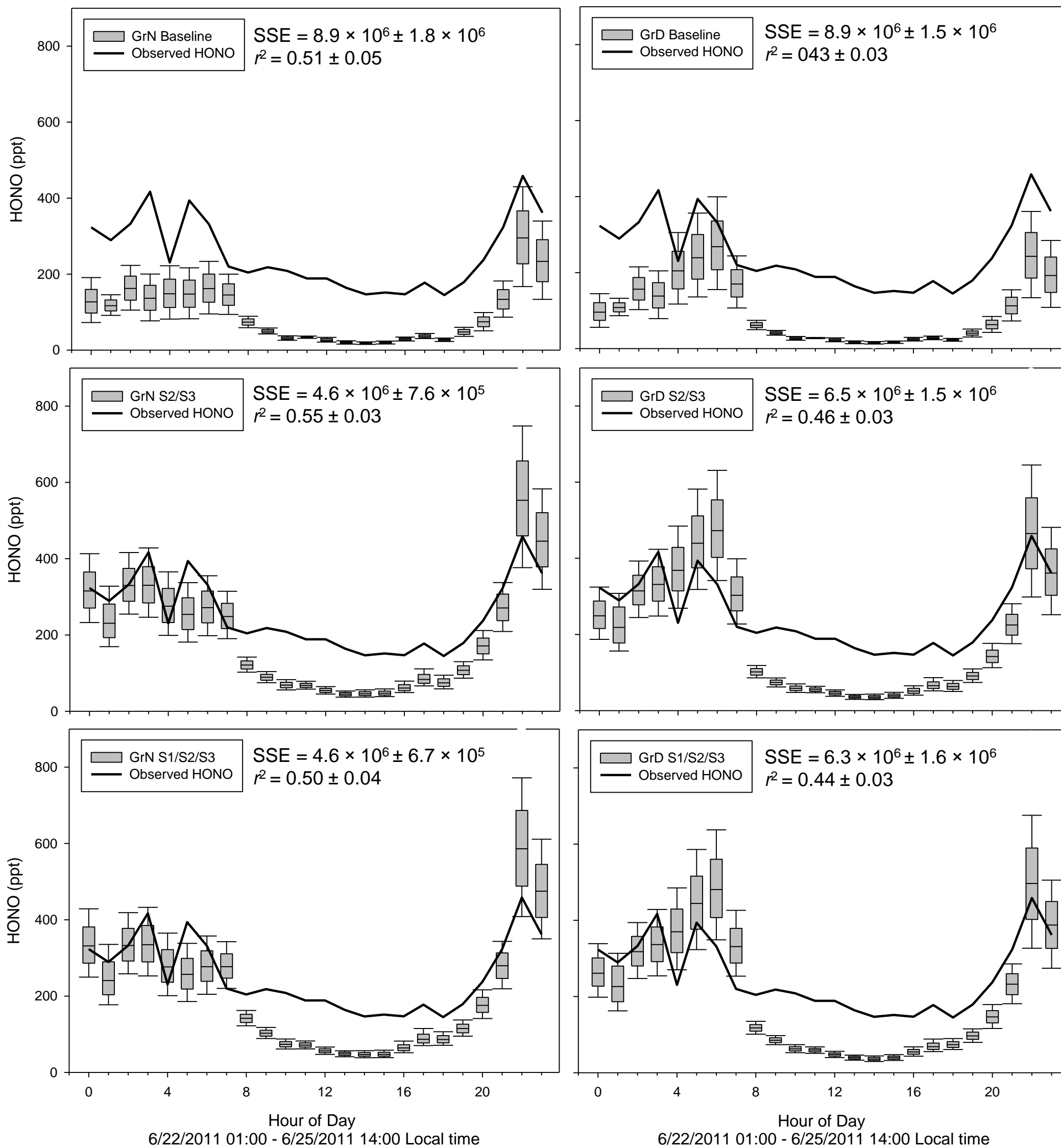


Figure 3. Summary of Monte Carlo simulation output for baseline scenarios, and scenarios with S2/S3 and S1/S2/S3 added to the baseline scenario.

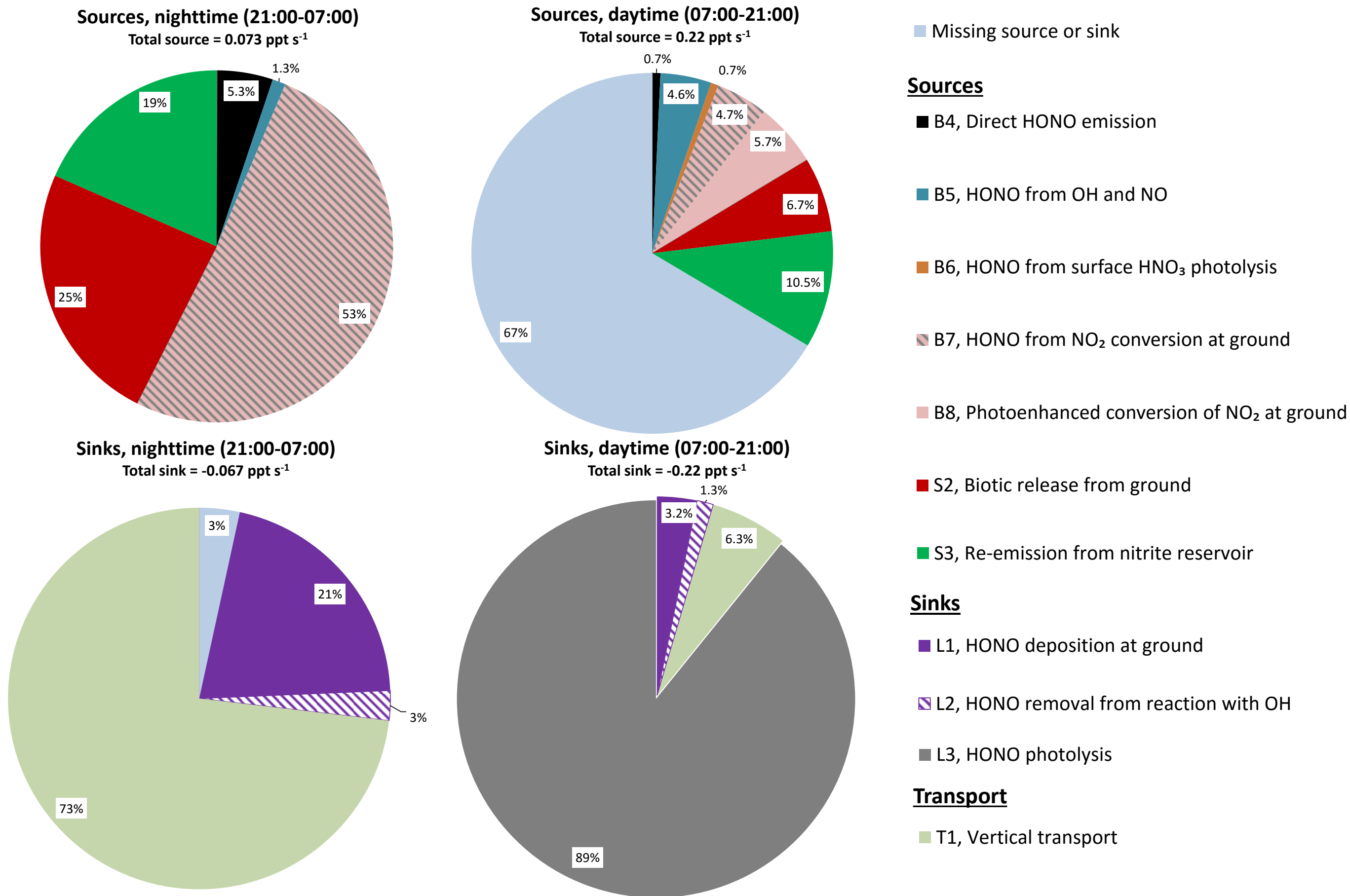


Figure 4. Relative contribution to HONO source or sink strength in GrN S2/S3 with 'likely' estimates of parameterizations. Contributions are averaged for the time period indicated above each pie chart across the modeling period (6/22/2011 01:00 – 6/25/2011 14:00 local time). Unknown source or sink is determined by stepwise addition of HONO source or sink such that modeled HONO equals measured HONO.



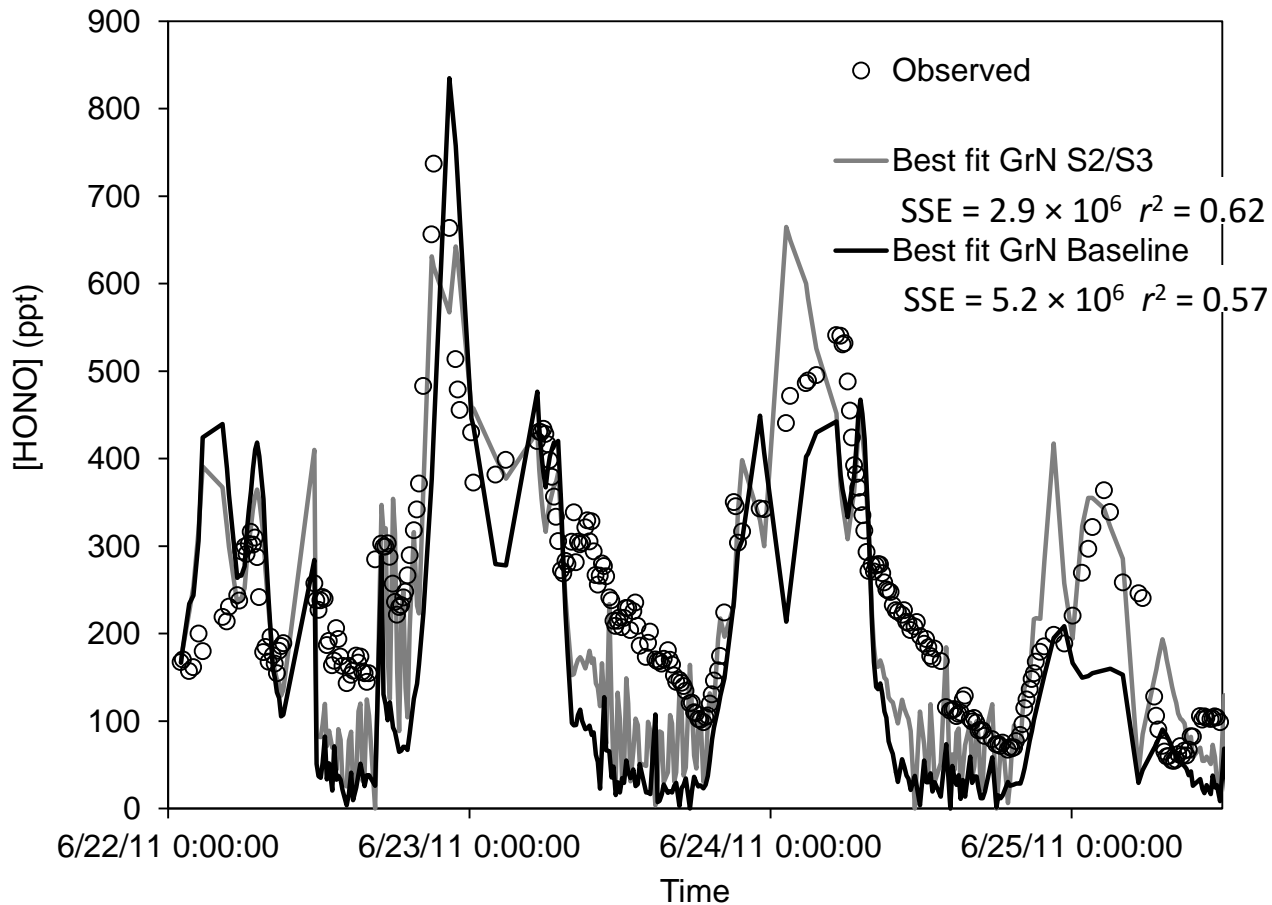


Figure 5. Model performance with best-fit parameters for the nighttime gradient (GrN) scenario with sources S2 and S3, compared to the nighttime gradient scenario with only baseline sources included.

- A two-layer box model evaluates HONO sources, sinks in outflow of Dallas-Fort Worth
- Monte Carlo simulation is applied to scenarios with 3 recently identified sources
- Improved model outcomes result from inclusion of 2 of 3 recently identified sources
- A substantial unknown source is still required for agreement with observation
- Missing HONO source is moderately correlated with  $j_{\text{NO}_2}$ , weakly correlated with  $\text{NO}_2$