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# Glacial Meltwater Modeling to Simulate Streamflow and Lake Levels in Taylor Valley, Antarctica

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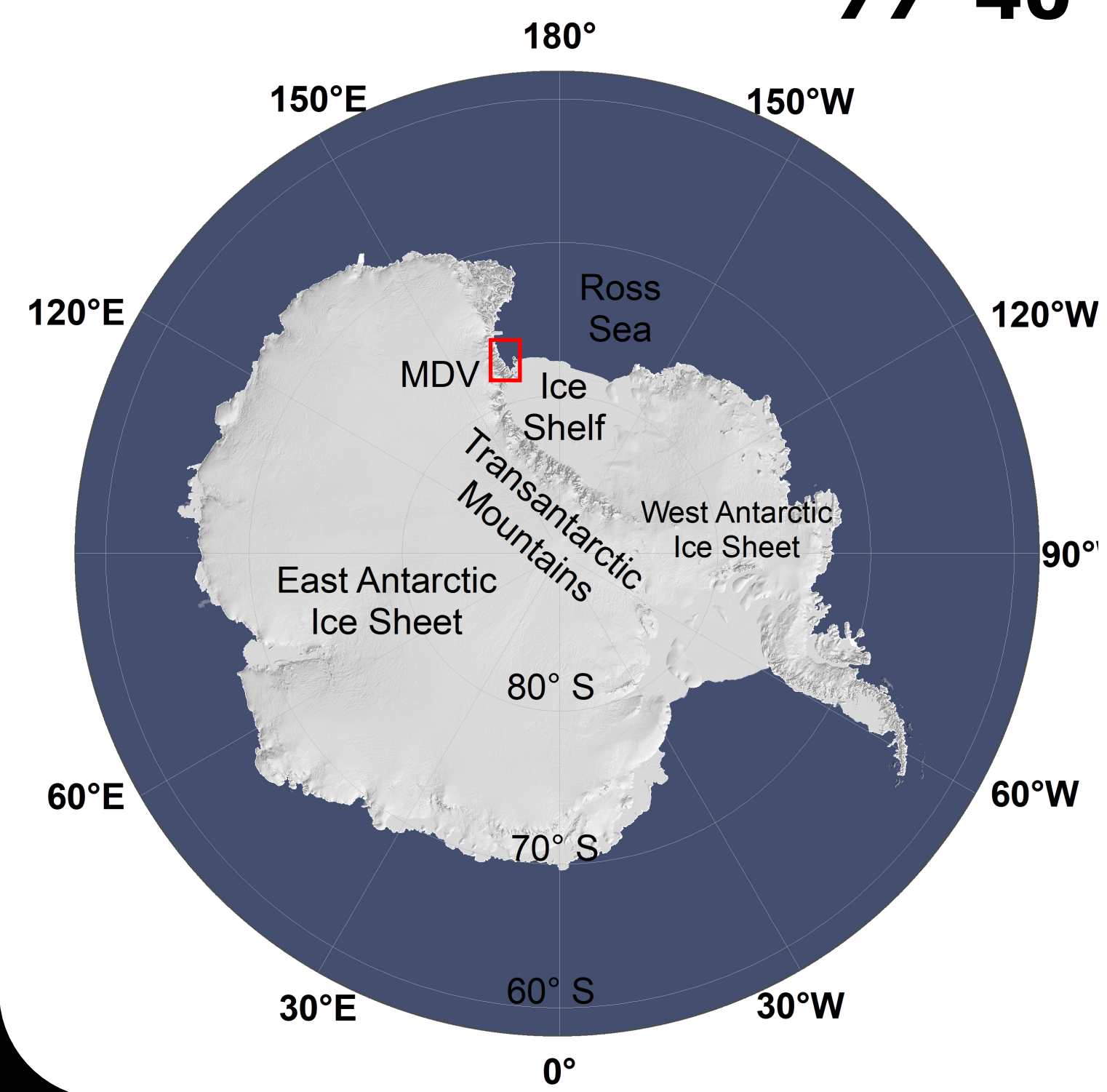
# Glacial meltwater modeling to simulate streamflow and lake levels in Taylor Valley, Antarctica

Julian Cross and Andrew G. Fountain

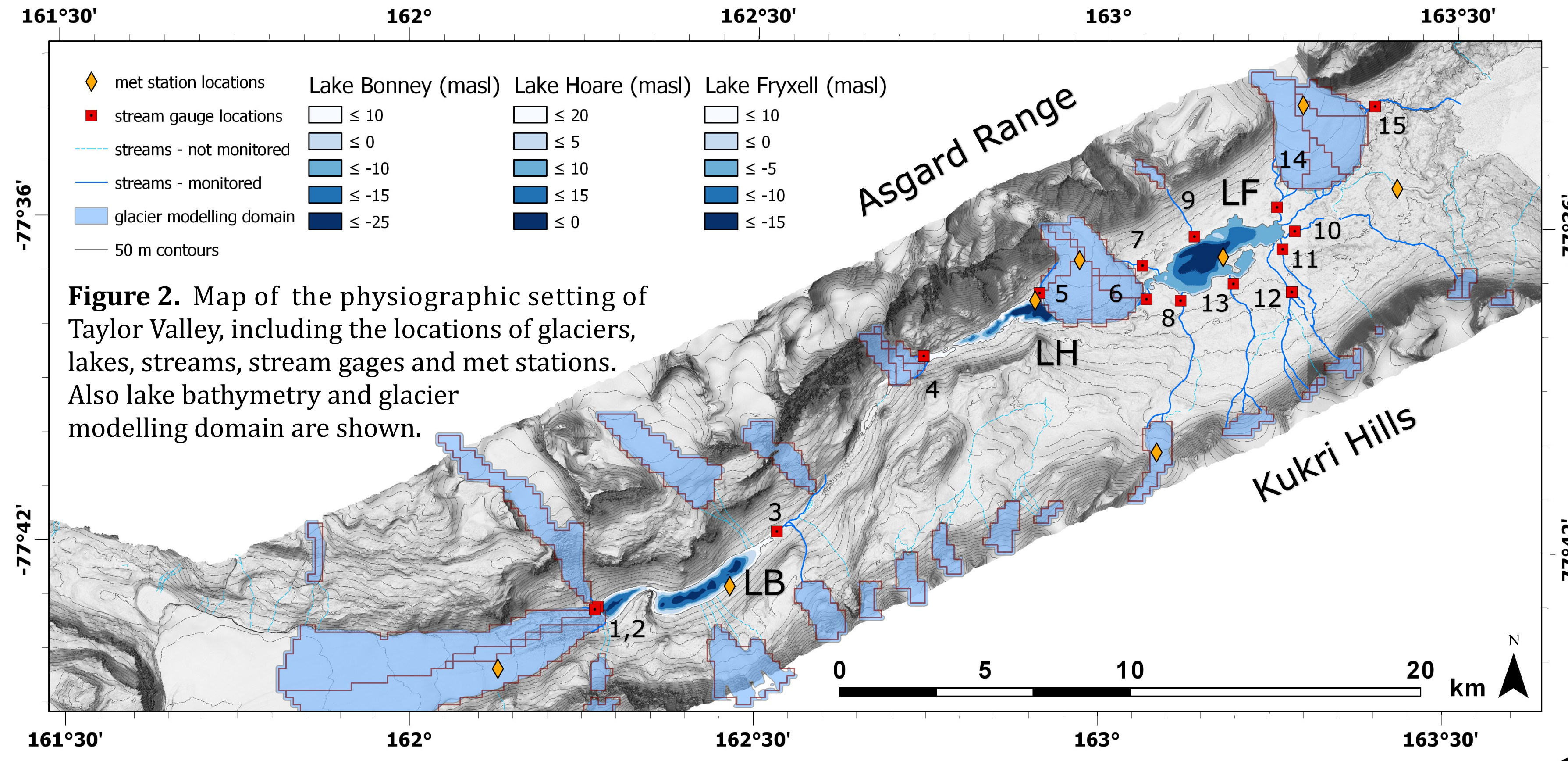
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## Study Site: Taylor Valley, Antarctica

77°40' S, 162°30' E



**Figure 1. (at left)** Map showing the location of the McMurdo Dry Valleys (MDV) region within Antarctica. The MDV, the largest ice-free region (4,500 km<sup>2</sup>) in Antarctica. The MDV are a polar desert: annual temperature of -18° C and precipitation <50 mm w.e. yr<sup>-1</sup>.



**Figure 2.** Map of the physiographic setting of Taylor Valley, including the locations of glaciers, lakes, streams, stream gages and met stations. Also lake bathymetry and glacier modelling domain are shown.

## I. Introduction

Three closed-basin, perennially ice-covered lakes occupy the valley floor: lakes Bonney, Hoare and Fryxell. Glacial meltwater accounts for nearly the total inflow to these lakes. Groundwater flux is essentially non-existent. Outflow is through sublimation of the frozen lake surface. Lake levels are highly sensitive to changes in climate and are mediated by the surface energy balance of the glaciers. With mean summer air temperatures are below 0°C, glacier ablation shows a complex sensitivity to solar radiation and wind speed. (Fountain et al. 1999)

### Objectives:

- Simulate modern streamflow and lake level change in Taylor Valley.
- Test meltwater and lake water balance model assumptions.
- Can these models be used to study paleo-lakes in the MDV?

## II. Methods

The distributed ICEMELT model (Hoffman et al. 2016) is applied to simulate streamflow and lake level from 1995 to 2013 at a grid resolution of 250 m. The meltwater model:

- is a distributed, physically-based energy balance model
- is driven by gridded local weather measurements
- was calibrated using ablation measurements at Taylor and Canada glaciers
- is tuned specifically to local ice and meteorological conditions
- assumes direct (same-day) meltwater routing

Additionally, the model accounts for solar radiation penetration into the ice, the spatial variability of albedo, and glacier topography that affects microclimate.

### Glacier Energy Balance Equation:

$$\chi(1 - \alpha) Q_{si} + Q_{li} + Q_{le} + Q_h + Q_l + Q_c = Q_M$$

A simple water balance method was used to estimate annual lake volume and level. The lake model relies on the following assumptions:

- simplifies inflows, treating glacier meltwater as the sole inflow
- accounts for sublimation from the lake ice surface and ignores groundwater and precipitation inflows
- basin geometry determined from 2 m lidar

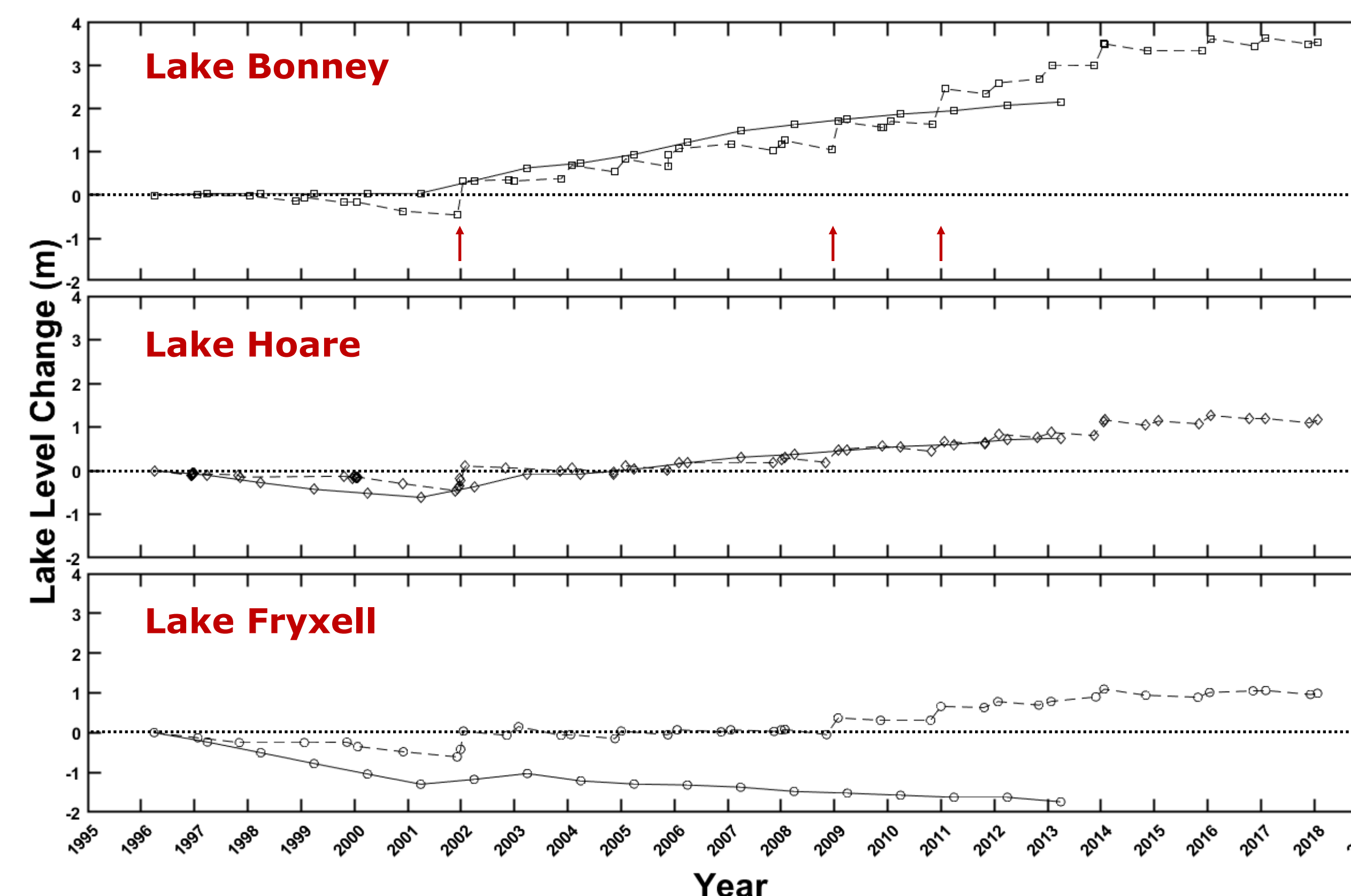
### Lake Water Balance Equation:

$$\frac{dV}{dt} = Q_{direct} + Q_{stream} + Q_{ground} + [P - S - E]A \rightarrow \frac{dV}{dt} = Q_{glacier} - S * A$$

### Simplified to:

Data Sources:  
Doran, P. 2016. McMurdo Dry Valleys Lake Levels. www.mcmiller.org.  
Gardner, C. 2007. McMurdo Dry Valleys Basic GIS Map Layers. www.mcmiller.org.  
Gooseff, M., and D. McKnight. 2016. McMurdo Dry Valleys Streamflow Daily Averages. www.mcmiller.org.  
McKnight, D., and M. Gooseff. 2017. McMurdo Dry Valleys Stream Descriptions. www.mcmiller.org.  
Prisco, J., and J. Schmok. 1995. McMurdo Dry Valleys Bathymetric Hypsographic Function Values. www.mcmiller.org.

## IV. Lake Level Results



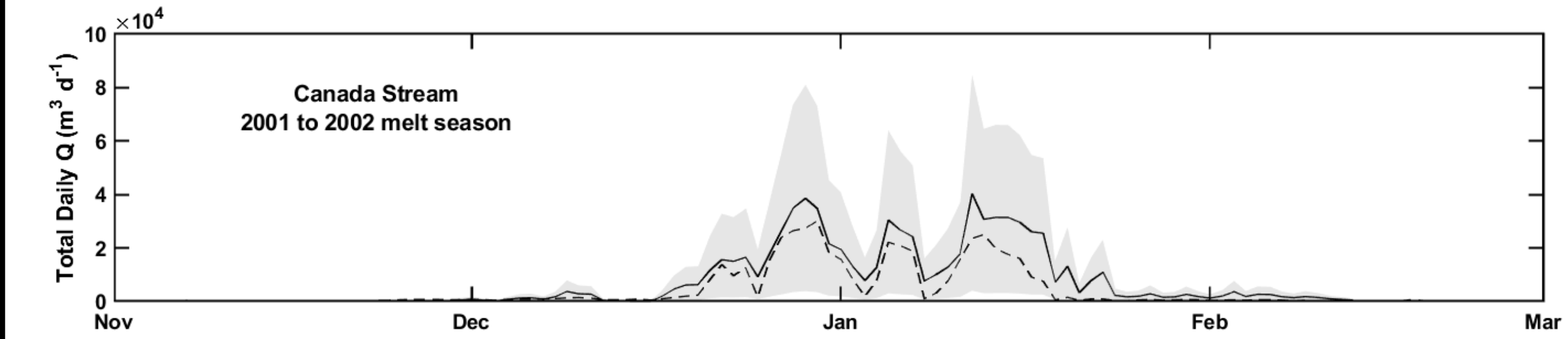
**Figure 7.** Simulated versus observed lake level change. Modeled using a sublimation rate of 0.35 m yr<sup>-1</sup> before 2004 and 0.27 m yr<sup>-1</sup> after 2003. Red arrows show 'flood years'.

## V. Conclusions

- With an adjustment of -0.05 to albedo, the meltwater model can predict streamflow moderately well.
- Predicted runoff results show higher efficiency in simulating dynamics but lower efficiency in simulating observed flow magnitudes.
  - Suggesting that routing assumptions hold (high dynamical efficiency)
  - Some meltwater sources are being neglected, e.g. ablation zone size or high elevation snow melt (low magnitude efficiency).
  - Meltwater model results are better for Canada and Taylor (Bonney basin) glaciers.
  - Increased model bias on Commonwealth and Kukri Hills glaciers suggest that unaccounted ice surface properties, e.g. surface roughness length, are present.
- Lake model results show that under-predicting melt inflow limits ability to model lake level change.
  - Particularly at Lake Fryxell and during flood years of 2001-02, 2008-09 and 2010-11.
- Lake surface sublimation rates required to fit lake level rise (0.27 to 0.35 m yr<sup>-1</sup>) are less than estimated, real-world values (0.35 to 1 m yr<sup>-1</sup>).

References:  
Fountain et al. 1999. Physical Controls on the Taylor Valley Ecosystem, Antarctica. *BioScience*.  
Hoffman et al. 2016. Distributed modeling of ablation (1996-2011) and climate sensitivity on the glaciers of Taylor Valley, Antarctica. *Journal of Glaciology*.

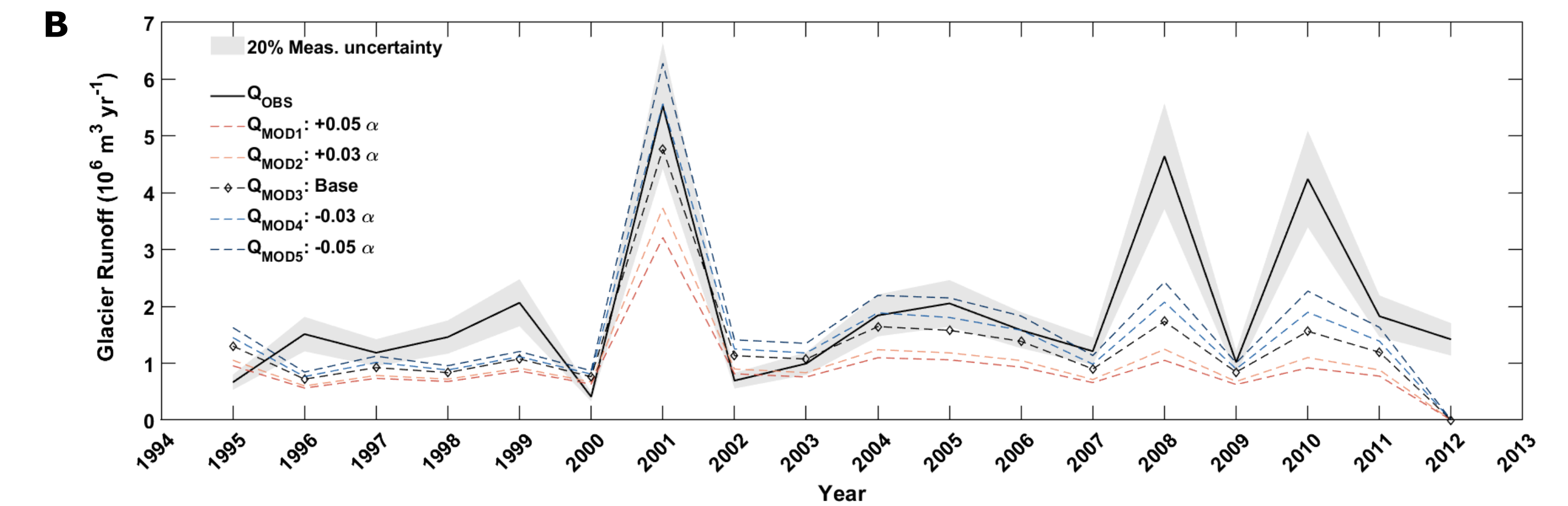
## III. Streamflow Results



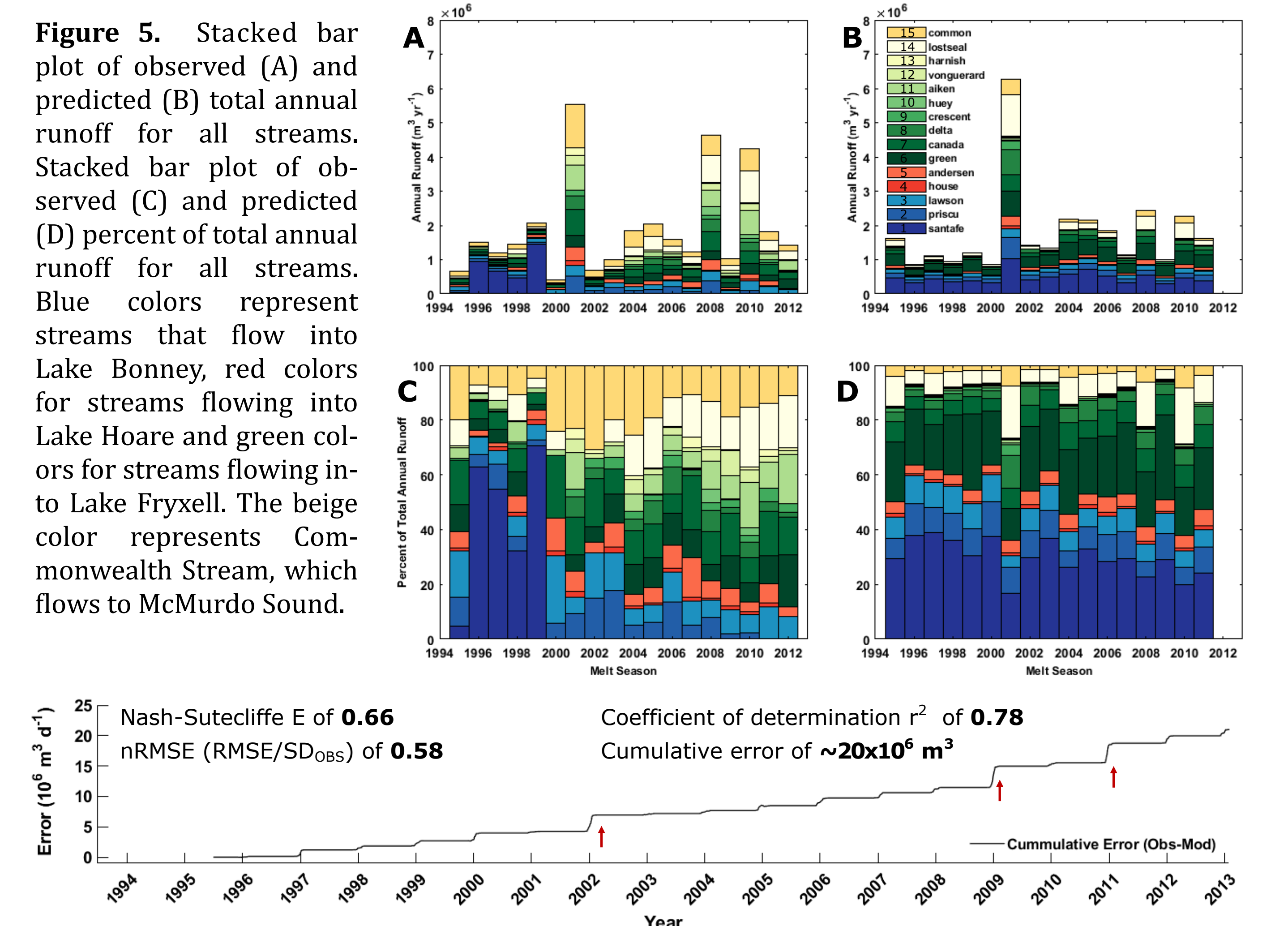
**Figure 3.** Simulated (dashed) versus observed (solid) daily runoff totals at Canada Stream (7) for the 2001-02 melt season. Associated daily streamflow measurement uncertainty (5-25%) is shown as the grey shaded area. This season was an anomalous 'flood year'.

**A**

	95-96	96-97	97-98	98-99	99-00	00-01	01-02	02-03	03-04	04-05	05-06	06-07	07-08	08-09	09-10	10-11	11-12	12-13	All
1 Santa Fe	0.37	0.07	0.09	0.20	-0.09	NAN	NAN	NAN	NAN	NAN	NAN	NAN	NAN	NAN	NAN	NAN	NAN	NAN	-0.05
2 Prisco	0.53	0.38	0.30	0.48	0.36	0.25	0.18	0.58	0.48	0.10	0.18	0.63	0.41	0.11	-0.53	0.70	NAN	NAN	0.32
3 Lawson	0.49	0.56	0.54	0.34	0.21	0.34	0.65	0.53	0.50	0.30	0.67	0.58	0.48	0.39	0.15	0.41	0.23	0.30	0.57
4 House	0.06	0.41	-0.18	0.09	0.07	0.08	0.75	NAN	0.33	0.33	0.28	0.02	0.68	0.74	-0.04	0.35	0.31	NAN	0.60
5 Anderson	0.04	0.50	0.20	0.17	0.14	0.14	0.24	0.76	0.61	-0.21	-0.21	0.70	0.78	0.07	0.57	0.65	0.55	0.46	0.68
6 Green	-0.55	0.01	-1.64	-0.01	-19.10	NAN	0.21	-4.92	-1.24	-1.01	0.36	0.19	0.52	0.71	-0.58	0.90	0.50	0.61	-0.13
7 Canada	-0.05	0.29	0.61	0.14	0.71	0.16	0.79	0.71	-0.60	0.05	0.55	0.77	0.39	0.53	0.30	0.55	0.62	0.31	0.62
8 Delta	NAN	NAN	-4.24	0.13	-33.93	0.30	0.20	-1.33	NAN	0.06	-0.12	0.31	0.23	0.58	-3.20	0.66	0.12	-1.21	0.41
9 Huey	NAN	0.25	0.07	-1.56	0.23	NAN	NAN	NAN	NAN	0.87	0.00	-0.08	0.06	-0.14	0.15	-0.06	NAN	NAN	-0.24
10 Alken	-0.04	-0.18	-0.13	-0.11	-0.11	0.09	-0.04	NAN	-0.05	0.00	-0.12	-0.06	-0.11	-0.21	-0.11	-0.16	-0.07	-0.00	-0.36
11 Von Guentard	0.00	0.03	0.13	-0.23	NAN	NAN	-0.09	-0.14	-0.20	-0.25	-0.24	-0.17	-0.15	-0.17	-0.26	-0.22	-0.17	-0.11	-0.52
12 Harnish	NAN	NAN	NAN	NAN	NAN	NAN	-0.06	-0.03	0.01	-0.11	-0.04	0.06	-0.24	-0.10	-0.04	-0.16	-0.58	-0.07	-0.34
13 Crescent	0.29	0.18	-0.14	0.10	0.42	NAN	-0.12	-0.02	-0.13	0.19	-0.05	-0.03	0.14	0.28	-0.19	-0.08	-0.09	-0.09	0.30
14 Lost Seal	0.73	0.27	0.29	0.30	0.32	NAN	NAN	0.31	0.45	0.25	0.74	0.24	0.29	0.04	0.58	0.65	0.08	0.49	0.49
15 Commonwealth	-1.47	-0.31	0.06	-0.26	-0.14	-0.44	0.44	-0.45	-0.45	0.09	-0.21	0.00	-0.13	0.28	0.48	0.37	0.27	-0.01	-0.11
All	0.27	0.23	0.47	0.33	0.15	0.39	0.82	0.59	0.64	0.61	0.79	0.65	0.47	0.46	0.49	0.71	0.37	0.66	0.66



**Figure 4.** Top panel (A) shows model efficiency in predicting total daily discharge using the Nash-Sutcliffe Efficiency (E) for all 15 streams by year. The left most column shows E values for each stream across all 18 modeled seasons and the bottom row shows E values for each season all seasons. The bottom panel (B) shows simulated (dashed) versus observed (solid) total annual runoff for all 18 modeled seasons based on 5 different models (MOD1 to MOD5). Model 3 is the base model with no global adjustment to albedo. MOD5 is the highest efficiency model with albedo lowered by 0.05, within the instrumental uncertainty of measured albedo values. All figures (including top panel) show results from MOD5.



**Figure 6.** Cumulative model error in (observations minus predictions) for total daily discharge across all model days. Red arrows show 'flood years'.