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# Ecohydrology of epiphytes: modeling water balance, CAM photosynthesis, and their climate impacts

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Submitted to Ecohydrology

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## Abstract

Epiphytes are aerial plants, often characterized by CAM (Crassulacean Acid Metabolism) photo-2 synthesis, which make up a significant portion of the biomass in some rainforests. Their unique з characteristics have not yet been included in ecohydrological models and their potential impact on 4 local hydrometeorology is largely unexplored. This work introduces a water balance model for epi-5 phytes, which adapts the soil-plant-atmosphere continuum model to represent a plant system with-6 out soil and couples it to the Photo3 photosynthesis model, which includes CAM photosynthesis. 7 The model, which is parameterized with field data of *Guzmania monostachia*, accurately captures 8 the observed hydraulic and photosynthetic behavior of the epiphytic species. The application of 9 vertical profiles of environmental inputs within the rainforest canopy shows increasing transpira-10 tion rates and decreasing water use efficiency with increasing canopy height, which corresponds to 11 observed distributions of epiphytes in rainforests. Given that vascular epiphytes constitute a maxi-12 mum of 35-50% of the foliar biomass in rainforests and contribute up to 13% of forest net primary 13 production, they may contribute up to 10-50% of total rainforest evapotranspiration, a significant 14 portion of the water cycle on the local ecosystem scale. The results of this work provide a missing 15 piece to current ecohydrological models, and can be integrated into Earth system models to help 16 improve the physical representation of transpiration and free-surface evaporation from canopy, in 17 current and future climates. 18

## <sup>19</sup> Keywords (8)

<sup>20</sup> epiphyte, Crassulacean acid metabolism (CAM) photosynthesis, evapotranspiration, plant hydraulics,

<sup>21</sup> plant water storage, modeling, rainforests

## 22 1 Introduction

Understanding and quantifying the mechanisms through which water moves through plants in the natural environment remains a central question in ecohydrology. Physiologically based plant models, which are used to quantify soil uptake and transpiration, typically model water fluxes through the soil-plant-atmosphere continuum. However, models for vascular epiphytes, which do not root in the soil, have not yet been developed or integrated into plant models. Accurately quantifying the water fluxes through epiphytes would greatly improve our understanding of their role in rainforests, arid ecosystems, and the larger global water cycle.

Epiphytes, by definition, grow non-parasitically on other plants, living on trunks and branches 30 of trees in the canopy. In contrast to plants that live in the soil, epiphytes obtain water and nutri-31 ents directly from the precipitation, air, and debris in their surrounding environment. Epiphytes 32 have been shown to exhibit marked differences in both leaf structure and water relations from 33 hemiepiphytes, their soil-rooting growth form (Holbrook & Putz, 1996). Globally, there are an es-34 timated 28,000 species of vascular epiphytes, which occur mostly in tropical regions (Zotz, 2016). 35 In rainforest ecosystems, vascular epiphytes are locally abundant and highly diverse, and they fill 36 important ecological niches, where they optimize space availability and often inhabit drier micro-37 areas within the canopy. Due to their drought tolerance and water use efficiency, epiphytes also 38 live in arid and semi-arid ecosystems. Epiphytes can affect the water budget of their environment 39 by changing the interception and water storage within the watershed, which indirectly affects soil 40 moisture (Pypker et al., 2006; Stanton et al., 2014). They can also affect the microclimate by 41 buffering temperature fluctuations and reducing the daytime vapor pressure deficit (Stanton et al., 42 2014). 43

Studies have shown that the most relevant abiotic constraint for the growth and vegetative function of epiphytes is water shortage (Zotz & Hietz, 2001). Epiphytes can obtain water from rain, dew, water vapor in the atmosphere, and stem runoff from their host. Water uptake and storage processes in epiphytes involve tank storage, absorptive structures such as trichomes and velamen, and large water storage capacitances in the plant tissue. Epiphyte tanks are water-impounding storage containers on the surface of the plants between the stem and the leaves that help to compensate

for access to soil moisture (Zotz & Thomas, 1999). The water balance of the tank depends on 50 the tank capacity, the catchment area of the plant, and aspects of the plant geometry that influence 51 evaporation (Zotz & Thomas, 1999). Trichomes are multicellular, hair-like structures located in 52 concavities on the leaf surface that absorb liquid water, water vapor, and nutrients from the atmo-53 sphere (Schmitt et al., 1989). Velamen, which are root-like structures, often serve to anchor the 54 epiphyte to its location, and can assist in absorbing water and nutrients from their environment as 55 well (Zotz, 2016). Epiphytes can also have a large water storage capacitance in their plant tissue, 56 allowing for long-term water storage between intermittent precipitation events (Martin & Schmitt, 57 1989). In addition, many species of epiphytes use CAM photosynthesis, a water-preserving photo-58 synthetic pathway in which they minimize water loss by closing their stomata during the day and 59 opening them at night when atmospheric vapor pressure deficits are lower (Benz & Martin, 2006; 60 Zotz, 2004). Facultative C3-CAM plants can reversibly induce CAM during dry periods to prolong 61 higher net carbon gain at lower water cost (Winter & Holtum, 2014). Silvera and Lasso (2016) es-62 timate that up to 50% of tropical epiphytic species exhibit some degree of CAM photosynthesis, 63 and Lüttge (2004) estimates that approximately 57% of vascular epiphytes are CAM plants. 64

According to Schlesinger and Jasechko (2014), tropical rainforests make up 16% of Earth's 65 land cover and receive an average of 1830 mm of precipitation per year, representing 35% of total 66 terrestrial precipitation. Based on data from MODIS and FAO models, tropical rainforests account 67 for 28.5 - 33.1% of terrestrial evapotranspiration (927-1076 mm/yr), and transpiration accounts 68 for 70  $\pm$ 14% of the total evapotranspiration (T/ET ratio) in tropical rainforests (Schlesinger & 69 Jasechko, 2014). On average, vascular epiphytes make up about 20% of the foliar biomass in rain-70 forests and montane forests of past studies (Zotz, 2016), but can make up to 35-50% of the foliar 71 biomass (Lüttge, 2004; Nadkarni, 1984), and contribute up to 13% of forest net primary production 72 (Richardson et al., 2000). Since epiphytes account for a significant portion of the foliar biomass of 73 some rainforests, they likely account for a large portion of the water budget in rainforests. Thus, 74 they may also make up a significant proportion of the total terrestrial evapotranspiration, and have 75 an impact on the global water fluxes modeled in climate and vegetation models. However, Earth 76 system models do not currently integrate epiphytes into the vegetation components of their mod-77 els, so the estimates of transpiration, free-surface evaporation from canopy, soil moisture, and T/ET 78

ratio may be less accurate. Since rainforests are important ecosystems for both the global water
 and carbon balances, a better physical representation of epiphyte behavior into the modeling of
 rainforest systems could impact the timing and magnitude of water and carbon fluxes.

Some developments in plant modeling have included limited epiphytic characteristics. For 82 example, Mu et al. (2011) improved the MODIS evapotranspiration algorithm by including night-83 time transpiration (i.e., cuticular transpiration) and evaporation from the water intercepted by the 84 surface of the canopy. Past studies on modeling the hydrological process of epiphytes have fo-85 cused on specific epiphytic species in certain locations and models of particular components of 86 epiphytes. For example, Jarvis (2000) integrated cloud deposition and interception from epiphytes 87 into a hydrological model of a montane cloud forest in Colombia. Zotz and Thomas (1999) devel-88 oped model calculations for tank water storage of bromeliads. Despite these advances, the unique 89 hydrological characteristics of epiphytes have not yet been developed into a plant hydraulic model. 90

As a first step towards a more accurate quantification of epiphyte water fluxes, this work in-91 troduces a comprehensive water balance model for vascular epiphytes. The hydrological model 92 is coupled with a photosynthesis model, Photo3, to represent both the C3 and CAM photosyn-93 thesis common in epiphytes (Hartzell et al., 2018). To model the water balance, the soil-plant-94 atmosphere continuum model from past studies is adapted to represent a plant system without soil. 95 The model incorporates epiphytic characteristics including water-impounding tanks, absorptive 96 structures such as trichomes and velamen, succulent properties, and the omission of root water 97 uptake. 98

<sup>99</sup> In this work, the model was parameterized for *Guzmania monostachia*, an epiphytic tank <sup>100</sup> bromeliad commonly known as a West Indian tufted airplant which performs C3 and facultative <sup>101</sup> CAM photosynthesis. The parameterized model is then verified with experimental data from two <sup>102</sup> past studies of transpiration and carbon assimilation in *G. monostachia* with primarily C3 photo-<sup>103</sup> synthetic behavior. Finally, the hydraulic behavior of a single tank CAM epiphyte in dry-down <sup>104</sup> conditions and the variation at different heights in a vertical canopy structure are explored. The <sup>105</sup> model accurately captures the hydraulic and photosynthetic behavior of epiphytes, while also high-

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lighting the need for more studies and modeling of epiphytes for future applications in climate
 modeling, ecological studies, and water resources engineering.

## **108 2** Methods

#### **109 2.1 Overview of Model**

The epiphyte model characterizes the main water fluxes and storages in epiphytes, as shown in Figure 1. In the water balance, there are two water storages: external tank storage and internal plant capacitance. The tank storage gains water from interception of precipitation, and loses water from free-surface evaporation and uptake from the plant. Water enters the epiphyte through uptake from trichomes, uptake from roots (e.g., velamen), and uptake from the tank. Water exits the epiphyte through transpiration from the stomata.

The epiphyte model is coupled to Photo3, which is a photosynthesis model that represents 116 the C3, C4, and CAM photosynthetic pathways in a consistent, physiologically based manner 117 (Hartzell et al., 2018). Photo3 includes the Farquhar et al. (1980) model for photosynthetic carbon 118 demand, an optimal control model for stomatal conductance (Buckley & Schymanski, 2014; Katul 119 et al., 2009; Medlyn et al., 2011), and a model of the soil-plant-atmosphere continuum for plant 120 hydraulics (Hartzell et al., 2018). For CAM photosynthesis, the photosynthetic C3 core is coupled 121 with a model for carbon fixation that is temporally separated from the Calvin cycle (Hartzell et al., 122 2018). 123

To capture the dynamics of the water fluxes through epiphytes, the epiphyte model follows an approach similar to the resistor-capacitor model of the soil-plant-atmosphere continuum (Hartzell et al., 2017, 2018; Hunt et al., 1991; Nobel & Jordan, 1983; Tyree & Ewers, 1991), without the soil component, and with the addition of schemes for external tank water storage and absorption of atmospheric water vapor through trichomes and velamen in humid conditions. The hydraulic model is a big-leaf plant model with an internal and external water storage capacity. Given inputs of air temperature, solar radiation, and specific humidity, the model estimates transpiration, carbon
 assimilation, and other hydraulic and photosynthetic variables.

#### **132 2.2 Model Implementation**

Environmental inputs to the epiphyte model include solar radiation ( $\phi$ ), specific humidity 133  $(q_a)$ , and air temperature  $(T_a)$ . Precipitation can either be provided as an input, set to 0 to simulate 134 a dry-down, or generated stochastically in the model. Outputs of the model include hydraulic 135 characteristics such as leaf water potential, and water balance processes such as transpiration, 136 free-surface tank evaporation, and changes in tank storage. The model operates on a 30-minute 137 time step to capture the strong dependence of CAM photosynthesis on variability in environmental 138 conditions (Hartzell et al., 2018). For all simulations, the model has a spin up period of at least one 139 day, in which it is forced with constant conditions (relative plant capacitance, relative tank water 140 content, etc.) that are equal to the initial conditions given as inputs to the model. This allows the 141 initial photosynthetic and hydraulic parameters to reach equilibrium before the model begins to 142 output data. 143

The epiphyte model is currently parameterized with hydraulic and photosynthetic properties 144 for Guzmania monostachia, a facultative C3-CAM tank epiphyte shown in Figure 1. G. monos-145 tachia typically exhibits C3 photosynthesis, but can exhibit CAM photosynthesis in response to 146 stresses such as drought and high light (Pierce et al., 2002). The epiphyte has a wide geographic 147 distribution in the tropics, from southern Florida through Peru and Bolivia, and is usually dis-148 tributed throughout the upper canopy in rainforests (Zotz & Andrade, 1997). This species was 149 selected because it is a well-studied epiphyte that lives in rainforest ecosystems, and it has impor-150 tant hydraulic characteristics (external tank storage) and photosynthetic characteristics (facultative 151 C3-CAM photosynthesis) that can be used to study plant hydraulic and photosynthetic behavior in 152 the rainforest canopy. The plant parameters for G. monostachia are meant to represent the epiphyte 153 at maturity, and are considered constant over the model duration. 154

#### 155 2.3 Plant Hydraulics

In the model, which is represented in Figure 2, water moves along a water potential gradient from the tank storage, through the plant, and into the atmosphere. The fluxes of water, including flux from tank  $(q_t)$ , flux from internal plant capacitance  $(q_w)$ , transpiration into atmosphere (T), and absorption of atmospheric water vapor  $(q_h)$ , are equivalent to the conductance (g) multiplied by the difference in water potential  $(\Psi)$  between each pair of nodes of the plant.

For consistency within the model, all water fluxes are expressed on a per unit ground area ba-161 sis. Fluxes commonly determined on a per unit leaf area basis, such as transpiration, are converted 162 to a per unit ground area basis with the leaf area index (LAI), which has units of  $m_{leaf}^2 / m_{ground}^2$ . 163 Fluxes determined on a per unit tank area basis, such as free-surface evaporation from the tank, are 164 converted to a per unit ground area basis with the tank area index (TAI), which has units of  $m_{tank}^2$ 165  $/ m_{ground}^2$ . The model represents an epiphyte at the plant scale, in which the area of the ground 166 equals the area of the plant. To scale the model to the ecosystem scale, the model results should be 167 multiplied by the ratio of epiphyte area to forest area. 168

The transpiration flux (T) is a function of the difference between the specific humidity internal to the leaf and the specific humidity of the atmosphere, i.e.,

$$T = g_{sa} LAI \frac{\rho_a}{\rho_w} (q_l(T_l, \psi_l) - q_a), \qquad (1)$$

where  $g_{sa} LAI$  is the series of the atmospheric and stomatal conductances normalized to a unit of ground area by the leaf area index,  $\rho_a$  is the density of air,  $\rho_w$  is the density of water,  $q_l$  is the specific humidity inside the stomata given as a function of the leaf temperature ( $T_l$ ) and leaf water potential ( $\psi_l$ ) (see Appendix A), and  $q_a$  is the specific humidity of the atmosphere.

When the humidity of the atmosphere is greater than the humidity of the leaf, water vapor from the atmosphere can be absorbed by the stomata or trichomes on the leaves of the epiphyte. The flux from the humid atmosphere into the leaf  $(q_h)$  is a function of the specific humidity internal to the leaf and the specific humidity of the atmosphere, i.e.,

$$q_h = g_h \, LAI \, \frac{\rho_a}{\rho_w} (q_a - q_l(T_l, \psi_l)), \tag{2}$$

where  $g_h LAI$  is the conductance of the absorption of atmospheric water vapor normalized to a unit of ground area by the leaf area index,  $\rho_a$  is the density of air,  $\rho_w$  is the density of water,  $q_a$  is the specific humidity of the atmosphere, and  $q_l$  is the specific humidity inside the stomata.

Many epiphytes are succulents with an internal plant capacitance. Following Bartlett et al. (2014) and Hartzell et al. (2017), the internal plant water storage is modeled as a concentrated storage, with adjustable parameter f representing the fraction of the plant resistance below the storage branch connecting to the xylem node of the plant (see Figure 2). In this model, it is assumed that the storage is located at mid-plant (f = 0.5). The flux between the internal plant capacitance and the xylem ( $q_w$ ) is a product of the storage conductance and water potential gradient, i.e.,

$$q_w = g_w \, LAI \, (\psi_w - \psi_x), \tag{3}$$

where  $g_w LAI$  is the storage conductance normalized to a unit of ground area by the leaf area index,  $\psi_w$  is the water potential of the internal storage, and  $\psi_x$  is water potential of the xylem node.

<sup>190</sup> The internal plant capacitance can be determined through the balance equation, i.e.,

$$LAI Z_w \frac{dx_w}{dt} = q_w = q_t + q_h - T,$$
(4)

where  $LAI Z_w$  is the total available water storage depth normalized to a unit of ground area by the leaf area index,  $x_w$  is the relative water content in the plant,  $q_w$  is the flux from the internal plant storage,  $q_t$  is the flux from the tank,  $q_h$  is the absorption of atmospheric water vapor, and T is transpiration.

The flux from the tank storage  $(q_t)$  is a product of the tank conductance and water potential gradient, i.e.,

$$q_t = g_{tpf} (\Psi_t - \Psi_x), \tag{5}$$

<sup>197</sup> where  $g_{tpf}$  is the series of storage and plant conductances to the relative plant height f,  $\psi_t$  is the <sup>198</sup> water potential of the tank, and  $\psi_x$  is the water potential of the xylem node. When the tank is <sup>199</sup> not empty, the water potential of the water in the tank ( $\psi_t$ ) is assumed to be 0 MPa since it is <sup>200</sup> free-standing water. The water potential of the internal storage ( $\psi_w$ ) is given as a function of the <sup>201</sup> pressure-volume curve (see Appendix C). The water potential at the xylem connection node ( $\psi_x$ ) <sup>202</sup> can be eliminated from Equations 3 and 5 by considering the water flux between the storage and <sup>203</sup> leaf nodes (solving Equation 8 for  $\psi_x$ ).

The free-surface evaporation flux from the epiphyte tank (E) is a function of the difference between the saturated specific humidity and the specific humidity of the atmosphere, i.e.,

$$E = g_a TAI \frac{\rho_a}{\rho_w} (q_{sat}^*(T_t) - q_a), \tag{6}$$

where  $g_a TAI$  is the atmospheric conductance normalized to a unit of ground area by the tank area index,  $\rho_a$  is the density of air,  $\rho_w$  is the density of water,  $q_{sat}^*$  is the saturated specific humidity at tank temperature ( $T_t$ ) (see Appendix A), and  $q_a$  is the specific humidity of the atmosphere. The temperature of the water in the tank was assumed to be an average of the leaf temperature and atmospheric temperature ( $T_t = (T_l + T_a)/2$ ).

#### The tank storage can be determined through the balance equation, i.e.,

$$TAI Z_t \frac{dx_t}{dt} = R(t) - E - q_t,$$
(7)

where *TAI Z<sub>t</sub>* is the maximum tank depth normalized to a unit of ground area by the tank area index,  $x_t$  is the relative fraction of tank storage with water, R(t) is the rainfall per unit ground area that falls on the epiphyte and enters the tank, *E* is the free-surface evaporation from the tank, and  $q_t$  is the uptake of tank water to the epiphyte. The rainfall (R(t)) can be entered as a model input, set to 0 to simulate a dry-down, or generated in the model as a stochastic marked Poisson process with a set mean rainfall frequency ( $\lambda$ ) and mean rainfall depth ( $\alpha$ ).

In addition, the fluxes through the leaf (transpiration and absorption of atmospheric water vapor) must be equal to the fluxes through the plant (flux from the tank and flux from the internal plant capacitance). Thus, the difference between the transpiration flux (*T*) and the absorption of atmospheric water vapor ( $q_h$ ) is a product of the plant conductance and water potential gradient, i.e.,

$$T - q_h = g_{pf} LAI (\Psi_x - \Psi_l), \tag{8}$$

where  $g_{pf} LAI$  is the plant conductance normalized to a unit of ground area by the leaf area index from the relative plant height f,  $\psi_x$  is the water potential of the xylem node, and  $\psi_l$  is the water potential of the leaf.

To solve the water balance of the epiphyte, the flux leaving the leaf node through transpiration is equal to the flux of water entering the leaf node, i.e.,

$$T = q_w + q_t + q_h, \tag{9}$$

where *T* is transpiration,  $q_w$  is the flux from the internal plant storage,  $q_t$  is the flux from the tank, and  $q_h$  is the absorption of atmospheric water vapor.

The system of water flux equations are connected with the equation for the energy balance of the plant system, which equates the incoming solar radiation to the outgoing sensible heat and latent heat fluxes, i.e.,

$$\phi = g_a \rho_a c_p (T_l - T_a) + \lambda_w \rho_w (T - q_h), \qquad (10)$$

where  $\phi$  is the net incoming solar radiation,  $g_a$  is the atmospheric conductance,  $\rho_a$  is the density of air,  $c_p$  is the specific heat of air,  $T_l$  is the leaf temperature,  $T_a$  is the atmospheric temperature,  $\lambda_w$  is the latent heat of vaporization,  $\rho_w$  is the density of water, T is the transpiration from the leaf, and  $q_h$  is the absorption of atmospheric water vapor. To simplify the model, the plant was considered as its own thermodynamic system. The tank was not considered to be a part of the plant system.

<sup>238</sup> Combining Equations 1, 2, 5, and 3 into Equation 9 gives the water balance for the epiphyte <sup>239</sup> as a function of the unknowns,  $\psi_l$  and  $T_l$ . Given the water balance and the energy balance from <sup>240</sup> Equation 10, the system of equations can be solved for  $\psi_l$  and  $T_l$ .

#### 241 2.4 Model Parameterization

The model was parameterized for *G. monostachia*, a facultative epiphytic tank bromeliad. The plant hydraulic parameters for *G. monostachia* are shown in Table 1.

The tank area index (*TAI*) was estimated from empirical equations from Zotz and Thomas (1999). In the study, Zotz and Thomas (1999) give both the projected area of the plant onto the ground ( $A_{proj}$ ) and the projected area of the tank ( $A_{tank}$ ) as a function of the dry weight of the plant for *G. monostachia*, i.e.,

$$A_{proj} = \frac{34.67 \, W}{1 + 0.086 \, W},\tag{11}$$

248 and

$$A_{tank} = \frac{5.57 \, W}{1 + 0.072 \, W},\tag{12}$$

where the areas  $(A_{proj} \text{ and } A_{tank})$  are in cm<sup>2</sup>, and the dry weight of the plant (*W*) is in g. Assuming a mean dry weight of 10 g for a mature *G. monostachia* plant, the ratio of  $A_{tank}$  to  $A_{proj}$ , or *TAI*, is 0.17 m<sup>2</sup><sub>tank</sub> / m<sup>2</sup><sub>ground</sub>.

The maximum tank depth  $(Z_t)$  was also estimated from empirical equations from Zotz and Thomas (1999). In the study, Zotz and Thomas (1999) give the maximum tank water content  $(C_{tank})$  as a function of the dry weight of the plant for *G. monostachia*, i.e.,

$$C_{tank} = \frac{5.48 \, W}{1 + 0.013 \, W},\tag{13}$$

where the maximum tank water content ( $C_{tank}$ ) and the dry weight of the plant (W) are in g. Assuming a mean dry weight of 10 g for a mature *G. monostachia* plant, and given the maximum tank water content ( $C_{tank}$ ) from Equation 13 and the area of the tank ( $A_{tank}$ ) from Equation 12, the maximum tank depth is 0.015 m.

The maximum plant capacitance depth  $(Z_w)$  was estimated from empirical equations from Zotz and Andrade (1997). In the study, Zotz and Andrade (1997) give the maximum plant water <sup>261</sup> content (*PWC*) as a function of the dry weight of the plant for *G. monostachia*, i.e.,

$$PWC = 1.50 + 4.99 W, \tag{14}$$

where the maximum plant water content (*PWC*) and the dry weight of the plant (*W*) are in g. Assuming a mean dry weight of 10 g for a mature *G. monostachia* plant, and given the maximum plant water content (*PWC*) from Equation 14 and the area of the plant ( $A_{plant}$ ) from Equation 11, the maximum plant capacitance depth is 0.0027 m.

The remaining plant hydraulic parameters in Table 1 are based on values given in past exper-266 imental studies or estimates given in existing models. The tank conductance  $(g_t)$  is not given in 267 past literature, but was estimated to be similar to soil conductances for rooting plants. The shape 268 parameters  $(d_1 \text{ and } d_2)$  and the total moles of solute  $(n_s)$  were determined by fitting a pressure-269 volume curve to observed data with a nonlinear plant capacitance equation (see Figure 3). The 270 plant capacitance is the sum of the osmotic pressure ( $\Omega$ ) and turgor pressure ( $\Pi$ ) (see Appendix 271 C). Several of the plant photosynthetic parameters for the CAM species in Photo3 (Hartzell et al., 272 2018) were adjusted to fit observed data from Zotz and Andrade (1997) and Pierce et al. (2002) for 273 G. monostachia. The adjusted photosynthetic parameters are given in Table 2. 274

Model results were compared to a study by Zotz and Andrade (1997) measuring water loss 275 in natural conditions during drought. In the study, 16 G. monostachia plants of various sizes were 276 used. The plants began well-watered ( $x_w = 1.0$ ), but without an initial tank water content ( $x_t = 0$ ). 277 Then the plants were allowed to dry under natural conditions, and were weighed daily at noon to 278 measure water loss, which was assumed to be equal to daily transpiration rates. Because the plants 279 were under well-watered conditions before the start of the experiment, and because G. monostachia 280 only exhibits some, but not full CAM behavior in response to stresses (Pierce et al., 2002), the plant 281 was modeled as a C3 plant. 282

The study by Zotz and Andrade (1997) was conducted in a clearing outside on Barro Colorado Island, Panama in February 1995. Since specific environmental conditions were not provided by Zotz and Andrade (1997), the environmental inputs for the model were obtained from the February 1995 weather data from the Lutz meteorological tower on Barro Colorado Island, from the Physical
 Monitoring Program of the Smithsonian Tropical Research Institute.

The model was then compared to results from a study by Pierce et al. (2002) measuring car-288 bon assimilation and C3-CAM behavior in a lab setting during drought. In this study, well-watered 289 G. monostachia plants were placed in a controlled environment, with temperatures of 28/22°C 290 (light/dark period), relative humidity of 60/90% (light/dark period), a photosynthetic photon flu-291 ence rate (PPFR) of 300  $\mu$ mol/(m<sup>2</sup> s) at plant height for a photoperiod of 12 hours, and no pre-292 cipitation input (Pierce et al., 2002). Similar to the experimental set-up, the model simulated the 293 same environmental conditions, and began with a well-watered plant ( $x_w = 1.0$ ) and an empty tank 294  $(x_t=0)$ . Because only 10% of carbon assimilation occurred at night, the plants were modeled as 295 C3 plants. 296

#### 297 2.5 Dry-down Simulation

A simulation of the epiphyte model was run to demonstrate the water use of the plant under 298 drought conditions. For the dry-down of the epiphyte, with results given in Section 3.2, the model 299 was run for one week with no precipitation input. The epiphyte began with a full tank ( $x_t = 1.0$ ) 300 and full internal plant capacitance ( $x_w = 1.0$ ), and was simulated as a C3 plant with hydraulic and 301 photosynthetic parameters for G. monostachia (see Tables 1 and 2). Although the epiphyte model 302 was parameterized to C3 photosynthesis, the model was also run with CAM photosynthesis (using 303 the existing C3 parameters) to demonstrate the qualitative CAM behavior from the model. To show 304 the long-term behavior without the impacts of variations in daily weather, the solar radiation and 305 atmospheric temperature were the repeated daily averages of the month of February of 1995 from 306 the Lutz meteorological tower on Barro Colorado Island, Panama (weather data obtained from 307 the Physical Monitoring Program of the Smithsonian Tropical Research Institute), and the relative 308 humidity was kept constant at 90%. 309

#### **310 2.6** Application to Rainforest Canopy

As an example of an application of the epiphyte model, and to understand how epiphytes may 311 respond to varying environmental conditions within a rainforest canopy microclimate, the model 312 was run with varying environmental inputs that simulated the vertical weather profiles within the 313 canopy. The model was run as a C3 plant in a dry-down simulation with a full initial external 314 tank ( $x_t = 1.0$ ) and internal plant capacitance ( $x_w = 1.0$ ). The environmental inputs, shown in Figure 315 4, were measured in February 2010 at the Lutz meteorological tower on Barro Colorado Island, 316 Panama (obtained from the Physical Monitoring Program of the Smithsonian Tropical Research 317 Institute) at canopy heights of 1 m (ground-level), 20 m, 42 m, and 48 m (top of canopy). Since the 318 solar radiation was only measured at 48 m, the solar radiation at 48 m was adjusted by a factor to 319 estimate the radiation at lower heights based on a similar observed canopy profile from Kumagai 320 et al. (2001). 321

## **322 3 Results and Discussion**

#### 323 3.1 Parameterization with G. monostachia

Figure 5 shows the observed changes in transpirational water stress of G. monostachia for 324 a 12-day rainless period during the 1995 dry season in Barro Colorado Island, Panama (Zotz & 325 Andrade, 1997) compared to the model transpiration results using the parameterizations from Ta-326 bles 1 and 2. Figure 5 also shows the percent error between the daily observations and model 327 results. In comparison to the observed results, the modeled results show a similar decrease in tran-328 spiration rates, with a progressive decrease within 4-5 days, and similar minimum transpiration 329 rates. The model generally shows similar behavior to the water loss curve from Zotz and Andrade 330 (1997). The model overestimation on the first day may be due to the assumption that the plants 331 started with a completely full internal plant capacitance ( $x_w = 1.0$ ), even though the actual capac-332 itance may have been slightly less than 1.0. After the initial decrease, the modeled transpiration 333 rates responded to daily fluctuations in environmental inputs reasonably well. However, due to 334

the uncertain weather data (the specific start date of the experiment was not specified in Zotz and Andrade (1997)), the model match to the observed data is uncertain. In addition, if the plants in the experiment exhibited some CAM behavior, the results may be different, since they were modeled with C3 photosynthesis.

Figure 6 shows the observed changes in carbon assimilation of G. monostachia for a 6-day 339 dry period under lab conditions (Pierce et al., 2002) compared to the modeled carbon assimilation 340 results using C3 photosynthesis and the parameterizations from Tables 1 and 2. Figure 6 also shows 341 the percent error between the maximum daily observed assimilation and the maximum daily model 342 assimilation. Both the model and observed data show carbon assimilation rates of 3-4  $\mu$ mol/(m<sup>2</sup> 343 s), which is below the maximum carbon assimilation rate for G. monostachia of 4.67  $\mu$ mol/(m<sup>2</sup> 344 s) determined by Males and Griffiths (2018). In comparison to the observed results, the modeled 345 results show a slower decrease in peak carbon assimilation rates, and lower carbon assimilation 346 rates integrated over each day. In the original study, the plants exhibited approximately 10% CAM 347 behavior, which is shown by the small peaks in assimilation during the night in Figure 6, which is 348 not shown in the modeled results, since the plants were modeled as C3 plants. 349

#### **350 3.2 Dry-down of an Epiphyte**

The results from the simulation show temporal estimates of carbon assimilation, water fluxes, tank free-surface evaporation, and plant transpiration during a 7-day dry-down of an epiphyte for both C3 photosynthesis (Figure 7) and CAM photosynthesis (Figure 8). The results of this simulation characterize the general hydrologic behavior of epiphytes from the model.

For both photosynthetic pathways, tank storage  $(x_t)$  decreases as evaporation (E) and uptake from the tank into the plant  $(q_t)$  occur. Due to increased temperature and solar radiation, more free-surface evaporation occurs during the day, until it stops when the tank is empty. For CAM photosynthesis, some uptake from the tank into the plant occurs at night when carbon assimilation is occurring.

The tank is completely empty after 1 day for C3 photosynthesis and 2 days for CAM photo-360 synthesis. The number of days it takes for the full tank to empty is within the range determined 361 experimentally by Zotz and Thomas (1999), but is lower than the average value from the study. 362 The time it takes for a full tank to dry out is lower in this model than in the results of Zotz and 363 Thomas (1999) likely because this model does not apply an empirical adjustment accounting for 364 the reduction in evaporation due to the protection from the epiphyte leaves. Figures 7 and 8 also 365 show internal plant capacitance storage  $(x_w)$  and flux  $(q_w)$  for the dry-down. When there is water 366 in the tank during the first few days, there is minimal flux from the internal plant capacitance, and 367 then flux increases as water is withdrawn from the plant capacitance under increasing water stress. 368

Figures 7 and 8 show transpiration (T, mm/d per unit ground area) for the 7-day dry-down 369 for C3 and CAM photosynthesis, respectively. For the C3 pathway, transpiration occurs during the 370 daytime, and for the CAM pathway, transpiration also occurs at night when carbon assimilation 371 occurs. The transpiration rates are significantly higher when the epiphyte tank has water. The leaf 372 water potential remains high when there is water in the tank since the tank water has a water poten-373 tial of 0 MPa, and then begins to decrease as transpiration occurs with an empty tank. The results 374 from Figures 7 and 8 agree with typical behavior in epiphytes characterized by rapid water uptake 375 during and immediately after precipitation events and low rates of water loss between events, as 376 described by Zotz (2016). For this specific simulation, there was no absorption of atmospheric 377 water vapor  $(q_h)$ . When the atmospheric specific humidity is higher than the leaf specific humidity, 378 which usually occurs under higher water stresses and more humid environments, the model results 379 show an absorption of atmospheric water vapor. 380

Figures 7 and 8 also show carbon assimilation ( $A_n$ ,  $\mu$ mol/(m<sup>2</sup> s) per unit leaf area) for the 381 C3 and CAM pathways. The C3 carbon assimilation results in Figure 7 show carbon assimilation 382 during the daytime. The CAM carbon assimilation results in Figure 8 reflect the typical daily 383 CAM cycle, which is characterized by the majority of the carbon assimilation occurring at night, 384 followed by a peak in carbon assimilation during the early morning, when the stomata remain open 385 for the continued net uptake of CO<sub>2</sub> (Bartlett et al., 2014; Winter & Smith, 1996). A small amount 386 of carbon assimilation also occurs in the evening, when the stomata open to uptake atmospheric 387 CO<sub>2</sub> that is immediately utilized in the Calvin cycle (Bartlett et al., 2014). Since the model was 388

parameterized for a C3 plant, the magnitude of carbon assimilation in the CAM simulation is higher than expected, but can be adjusted in future studies when more CAM experimental data is available. Similar to transpiration, the carbon assimilation rates are higher when the epiphyte tank has water. The low stomatal conductances in Figures 7 and 8 agree with Zotz (2016), who observes that epiphytes tend to have a low stomatal conductance, and thus low transpiration rates.

Previous studies are lacking experimental observations of photosynthetic and hydraulic be-394 havior of epiphytes in the transition between when there is water in the external tank and when 395 the tank is empty. The epiphyte model presented in this study shows distinct behavior in the two 396 phases, but the behavior with water in the tank and the transition to an empty tank has not yet 397 been verified with experimental data. Nonetheless, the model still shows the expected behavior of 398 a plant under stressed and non-stressed conditions. Detailed studies of epiphytes that exhibit full 399 CAM photosynthesis are also missing in current research. In future studies, the epiphyte model can 400 be used in tandem with experiments to formulate new hypotheses and better understand epiphyte 401 tank water use behavior. 402

#### **3.3** Application to Vertical Canopy Profiles

Figure 9 shows the simulation results of daily transpiration rates (T, mm/d per unit ground)404 area) and water use efficiency (mmol<sub>CO<sub>2</sub></sub> assimilated per mol<sub>H<sub>2</sub>O</sub> transpired) at canopy heights of 405 1, 20, 42, and 48 m when the rainforest canopy environmental inputs from Figure 4 were applied 406 to the model. The application of the vertical profiles of canopy microclimatic inputs show that 407 transpiration rates generally increase with increasing canopy height when the plant begins the 408 simulation well-watered, which is a result of the higher solar radiation and temperature higher in 409 the canopy. As the plant dries out, the epiphytes located lower in the canopy show more stable 410 transpiration rates, while the epiphytes in the upper canopy show rapidly decreasing transpiration 411 rates. Water use efficiency is higher for plants at lower heights throughout the simulation. The 412 varying rates of transpiration are consistent with the observed vertical distribution of epiphytes in 413 the canopy. Zotz (2016) observes that most epiphytes grow at heights of 5-20 m in a 40 m tall 414

rainforest, which corresponds to the results from this simulation that show that the plants at lower
heights have more stable transpiration rates and higher water use efficiencies.

#### 417 3.4 Implications for Microclimate and Earth System Modeling

The unique hydraulic processes of epiphytes have been shown to have effects on rainforest 418 canopy microclimate. The presence of epiphytes as shade-providers and water reservoirs impact 419 the temperature and humidity in the canopy. Within the rainforest canopy, epiphytes behave as 420 a capacitor, dampening the daily fluctuations in atmospheric humidity and temperature (Freiberg, 421 2001; Stanton et al., 2014). In a study of a tropical forest during the late dry season, Stuntz et al. 422 (2002) found that the total evapotranspiration from tree crowns with epiphytes, including epiphyte 423 evapotranspiration, is less than the total evapotranspiration from tree crowns without epiphytes. 424 The study also showed that within a single tree crown, vascular epiphytes provide microsites 425 with lower temperatures and less evapotranspiration than do areas without epiphytes (Stuntz et 426 al., 2002). The model presented in this paper provides a first step in integrating these epiphytic 427 impacts on microclimate into a plant model. Through future epiphyte modeling advancements, 428 the complex feedbacks between epiphytes, their host vegetation, the atmosphere can be further 429 explored. 430

As vegetation components of Earth system models become more advanced, integrating epi-431 phytic behavior, as outlined in this study, may help to improve estimates of water fluxes on both 432 a global and local scale. Although the total biomass of epiphytes in rainforests is unknown, vas-433 cular epiphytes constitute an average of 20% and a maximum of 35-50% of the foliar biomass in 434 rainforests of past studies (Lüttge, 2004; Nadkarni, 1984; Zotz, 2016), and contribute up to 13% 435 of forest net primary production (Richardson et al., 2000). Epiphytes may account for a simi-436 lar portion of the water budget in rainforests. Given rainforests provide approximately 30% of 437 global evapotranspiration (Schlesinger & Jasechko, 2014), epiphytes could contribute up to ap-438 proximately 3-6% of total terrestrial evapotranspiration, a significant water flux for Earth system 439 modeling. 440

Further development of Earth system models could benefit from the integration of specific epi-441 phytic characteristics. Liu et al. (2020) highlighted the importance of integrating plant hydraulics 442 into Earth system models to more accurately estimate transpiration in response to atmospheric 443 moisture stress, but the integration of epiphyte hydraulics has not yet been explored. A study that 444 compares current Earth system model performance with and without the integration of an epiphyte 445 model is needed to quantify the impacts that epiphyte modeling could have on the accuracy of wa-446 ter fluxes. Or and Lehmann (2019) integrated a canopy interception component to estimate global 447 surface evaporation based on the leaf area index and a maximum water layer thickness on leaves. 448 Mu et al. (2011) modified the MODIS evapotranspiration algorithm by including nighttime tran-449 spiration (i.e., cuticular transpiration), and evaporation from the water intercepted by the surface of 450 the canopy. More detailed models, such as the one presented in this work, which models nighttime 451 transpiration (i.e., CAM photosynthesis) and both free-surface evaporation and plant uptake from 452 canopy interception, could help to produce a better physical representation of evapotranspiration in 453 hydrological and climate models. Both the magnitude and timing of transpiration from the canopy 454 could be different in this model than in traditional plant models. In addition, the volume of inter-455 cepted precipitation stored in the canopy may experience quicker uptake in the epiphyte model, 456 in which the water can be used for transpiration, unlike traditional plant models. In addition, the 457 integration of a plant system without soil into the Earth system models could impact microclimatic 458 atmospheric humidity due to uptake from epiphytes, and may have impacts on the estimated soil 459 moisture as well. 460

### **461 4 Conclusions**

This work introduces a new model that represents epiphyte water use behavior. The epiphyte model, for the first time, presents a physiologically based vascular plant system model without soil. Using the Photo3 model that represents both C3 and CAM photosynthesis, the model applies environmental inputs to a big-leaf, numerical water balance model, with the Farquhar et al. (1980) model for photosynthetic carbon demand, an optimal control model for stomatal conductance, and a resistor-capacitor model of the plant-atmosphere continuum for plant hydraulics. Hydraulic

characteristics unique to epiphytes, including external tank storage, absorption of atmospheric 468 humidity, succulent properties, and omission of root uptake from soil, are integrated into the model. 469 The parameterized model results show some agreement with simulations of G. monostachia under 470 drought conditions, but further validation is needed to characterize epiphyte water use when under 471 well-watered conditions, when transitioning between conditions with full and empty external tank 472 storage, and when exhibiting full CAM photosynthetic behavior. When forced with vertical profiles 473 of environmental inputs measured within the rainforest canopy, the results of the model show 474 increasing transpiration rates and decreasing water use efficiency with increasing canopy height. 475 The modeled optimal water use efficiency at low to mid canopy heights agrees with the higher 476 observed frequency of epiphytes at low to mid heights in rainforests. 477

The results of this work demonstrate the first step towards a more accurate quantification of epiphyte water fluxes. In the future, the epiphyte model can be integrated into Earth system models to help produce more reliable estimates of transpiration and free-surface evaporation from canopy. As climate change and sustainable water use become more pressing issues in the future, ecohydrological models, including this epiphyte model, will be helpful in quantifying the hydraulic behavior of plants and global water fluxes.

## **Appendix A: Specific Humidity**

Following Jones (1992), the specific humidity of the leaf  $(q_l)$  can be calculated as

$$q_l(T_l, \Psi_l) = q_{l \, sat}(T_l) \, \exp\left[\frac{V_w \, \Psi_l}{R \, T_l}\right],\tag{15}$$

where  $q_{l sat}$  is the saturated specific humidity at leaf temperature  $(T_l)$ ,  $V_w$  is the partial molar volume of water,  $\psi_l$  is the water potential of the leaf, and *R* is the universal gas constant.

The saturated humidity  $(q_{sat})$  is related to the temperature (T) by (Jones, 1992)

$$q_{sat}(T) = \frac{0.622}{p_a} a_{sat} \exp\left[\frac{b_{sat}(T-273)}{c_{sat}+T-273}\right],$$
(16)

where  $q_{sat}$  is the saturated specific humidity at temperature *T* in Kelvin,  $p_a$  is atmospheric pressure (Pa), and  $a_{sat}$ ,  $b_{sat}$ , and  $c_{sat}$  are empirical constants given in Table 3.

## **Appendix B: Conductances**

Following Daly et al. (2004), the plant conductance  $(g_p)$  is modeled by a vulnerability curve, i.e.,

$$g_p = g_{p max} \exp\left[-\left(\frac{-\Psi_l}{j}\right)^h\right],$$
 (17)

so that  $g_p$  is near  $g_{p max}$  for high  $\psi_l$  and is close to 0 for low  $\psi_l$ , since  $g_p$  drops when the water potential is too low because of xylem cavitation. The shape parameters *j* and *h* are both equal to 2 (Hartzell et al., 2018).

<sup>497</sup> The internal storage conductance  $(g_w)$ , is modeled in a similar manner to the plant conduc-<sup>498</sup> tance  $(g_p)$ , i.e.,

$$g_w = g_{w max} \exp\left[-\left(\frac{-\Psi_w}{j}\right)^h\right],\tag{18}$$

To simplify the model, and because the water potential of the water in the tank is always 0, the tank conductance  $(g_t)$  is set to a constant value when there is water in the tank (see Table 1), and 0 when the tank is empty.

Following Jones (1992), the stomatal conductance for water  $(g_s)$  is closely related to the stomatal conductance for CO<sub>2</sub>, i.e.,

$$g_s = 1.6 \, g_{s \, CO_2} + g_{cut}, \tag{19}$$

where  $g_{sCO_2}$  is the stomatal conductance for CO<sub>2</sub> (see Photo3 model (Hartzell et al., 2018) for details on calculations) and  $g_{cut}$  is the cuticular conductance, which accounts for the small amount of water vapor lost when carbon assimilation is not occurring. The conductance of the absorption of atmospheric water vapor  $(g_h)$  was assumed to be similar in magnitude to the cuticular conductance  $(g_{cut})$ .

The atmospheric conductance  $(g_a)$  is a function of wind speed and canopy height (Jones, 1992), i.e.,

$$g_a = \frac{u_z \, k^2}{(\ln[(z-d)/z_o])^2},\tag{20}$$

where  $u_z$  is the wind speed, k is the Von Karman constant, z is the height of the wind speed measurement, d is 0.64 \* canopy height, and  $z_o$  is 0.13 \* canopy height.

## **Appendix C: Water Potentials**

Since the water in the tank is free-standing water, it is assumed to have a water potential ( $\psi_t$ ) of 0 MPa.

The plant water potential ( $\psi_w$ ) is the sum of the osmotic and turgor pressure, following the method given by Bartlett et al. (2014). The plant osmotic pressure ( $\Omega$ ) is given by (Hem, 1985)

$$\Omega = \frac{RT_l}{V_w} \ln\left(\frac{n_w}{n_w + n_s}\right),\tag{21}$$

where *R* is the ideal gas constant,  $T_l$  is the leaf temperature (K),  $V_w$  is the molar volume of liquid water (m<sup>3</sup>/mol),  $n_s$  is the total moles of solute, and  $n_w$  is the total moles of water, on a total leaf area basis (mol/m<sup>2</sup>). The total moles of water is given by

$$n_w = \frac{x_w Z_w}{V_w},\tag{22}$$

where  $x_w$  is the relative plant capacitance fraction,  $Z_w$  is the maximum water storage capacitance, and  $V_w$  is the molar volume of liquid water. The osmotic pressure when the plant is well-watered  $(x_w = 1, \text{ full turgor})$  is determined by the total moles of solute  $(n_s)$ , which is assumed to be a constant for simplicity. The osmotic potential of epiphytes is usually at a lower magnitude (-1 MPa) than water stressed terrestrial plants because in the absence of soil, a more negative osmotic <sup>526</sup> potential is not needed to drive water uptake during short pulses of water availability (Zotz & Hietz, <sup>527</sup> 2001). When the plant is well-watered, the osmotic pressure is balanced by the turgor pressure. <sup>528</sup> The plant turgor pressure ( $\Pi$ ) is given by a power law equation (Ranney et al., 1990), i.e.,

$$\Pi = (x_w - d_1)^{d_2},\tag{23}$$

where  $x_w$  is the relative plant capacitance fraction, and  $d_1$  and  $d_2$  are fitted values based on observed pressure-volume data from Zotz and Andrade (1997) (see Figure 3).

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## 538 Conflict of Interest

<sup>539</sup> No conflict of interest was declared.

## **Data Availability Statement**

The epiphyte model software can be accessed at https://github.com/gretamiller/Epiphyte. It was developed in Python 3.7 by Greta Miller, Samantha Hartzell, and Amilcare Porporato, and first made available in 2020. It was built off of the Photo3 photosynthesis model for rooting plants,

- <sup>544</sup> which can be accessed at https://samhartz.github.io/Photo3/. Additional data that support the find-
- <sup>545</sup> ings of this study are available upon request.

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Parameter	Value	Units	Description	
LAI	3.0 <sup>a</sup>	$m_{leaf}^2$ / $m_{ground}^2$	Leaf area index	
TAI	0.17 <sup>b</sup>	$m_{tank}^2$ / $m_{ground}^2$	Tank area index	
$Z_t$	0.015 <sup>b</sup>	m	Maximum tank storage depth	
$Z_w$	0.0027 <sup>b, c</sup>	m	Maximum plant capacitance depth	
a	12 <sup>d</sup>	mm/s	Atmospheric conductance	
8a	12		per unit ground area, Eq. 20	
g p max	0.076 <sup>e</sup>	µm/(s MPa)	Maximum plant conductance per unit leaf area	
a	0.0045 <sup>f</sup>	µm/(s MPa)	Maximum internal storage conductance	
8w max			per unit leaf area	
8cut	0.01 <sup>g</sup>	mm/s	Cuticular conductance per unit leaf area	
<b>a</b> .	0.01 <sup>g</sup>	mm/s	Conductance of the absorption of atmospheric	
8 <i>h</i>			water vapor per unit leaf area	
$g_t$	0.5 <sup>h</sup>	µm/(s MPa)	Tank conductance	
f	0.5 <sup>f</sup>	_	Relative height of plant capacitance	
h	$2^{\mathrm{f}}$	_	Shape parameter for Eq. 17 and 18	
j	$2^{\mathrm{f}}$	_	Shape parameter for Eq. 17 and 18	
$d_1$	$0.0^{i}$	_	Shape parameter for Eq. 22	
$d_2$	2.0 <sup>i</sup>	_	Shape parameter for Eq. 22	
$n_s$	1.1 <sup>i</sup>	mol/m <sup>2</sup>	Shape parameter for Eq. 23	

Table 1: Plant hydraulic parameters for G. monostachia

<sup>a</sup> Based on reasonable estimate of similar species
<sup>b</sup> Based on Zotz and Thomas (1999)

<sup>c</sup> Based on Zotz and Andrade (1997)

<sup>d</sup> Based on Jones (1992) for a wind speed of 0.75 m/s at 2 m altitude, with a plant height of 0.5 m

<sup>e</sup> Based on North et al. (2016)

<sup>f</sup> Based on estimate from Hartzell et al. (2018)

<sup>g</sup> Based on Zotz (2016)

<sup>h</sup> Estimate within range of soil conductances for rooting plants

<sup>i</sup> Fitted to pressure-volume curve from Zotz and Andrade (1997)

Parameter	Value	Units	Description
$\Psi_{LA0}$	-1.0 <sup>a</sup>	MPa	Point of maximum plant water stress
$\Psi_{LA1}$	-0.3 <sup>a</sup>	MPa	Onset of plant water stress
$V_{c,max0}$	10.4 <sup>b</sup>	$\mu$ mol/(m <sup>2</sup> s)	Maximum carboxylation rate
$J_{c,max0}$	20.8 <sup>b</sup>	$\mu$ mol/(m <sup>2</sup> s)	Electron transport rate
M <sub>max</sub>	152 <sup>b</sup>	mol/m <sup>3</sup>	Maximum malic acid concentration

Table 2: Photosynthetic parameters for *G. monostachia* 

<sup>a</sup> Parameters for CAM species in Hartzell et al. (2018), adjusted to fit observed data
<sup>b</sup> Parameters for CAM species in Hartzell et al. (2018), reduced by factor of 0.8 to fit observed data

Parameter	Value	Units	Description
a <sub>sat</sub>	613.75 <sup>a</sup>	_	Constant for Eq. 16
$b_{sat}$	17.502 <sup>a</sup>	_	Constant for Eq. 16
c <sub>sat</sub>	240.97 <sup>a</sup>	_	Constant for Eq. 16
$p_a$	101,325	Pa	Atmospheric pressure
$\lambda_w$	$2.5^{*}10^{6}$	J/kg	Latent heat of vaporization
$\rho_a$	1.27	kg/m <sup>3</sup>	Air density
$ ho_w$	998	kg/m <sup>3</sup>	Water density
$c_p$	1012	J/(kg K)	Specific heat of air
R	8.314	J/(mol K)	Universal gas constant
$V_w$	1.81*10 <sup>5</sup>	m <sup>3</sup> /mol	Molar volume of water
k	0.41	_	Von Karman constant

Table 3: Fixed Constants

<sup>a</sup> Based on Jones (1992)



Figure 1: (A) Photograph of *Guzmania monostachia* in Maui, Hawaii (photo by Forest and Kim Starr, under Creative Commons License: https://creativecommons.org/licenses/by/2.0/deed.en). (B) Water balance of an epiphyte, where the water fluxes are labeled in black and the water storages are labeled in blue.



Figure 2: Model schematic of water fluxes and storages for epiphytes. Water from the tank storage, with a maximum height of  $Z_t$ , and water from internal plant capacitance, with a maximum height of  $Z_w$ , moves along a water potential gradient, through the plant, and into the atmosphere with humidity  $q_a$ . The fluxes of water, including flux from tank  $(q_t)$ , flux from internal plant capacitance  $(q_w)$ , are equivalent to the conductance (g) multiplied by the difference in water potential  $(\Psi)$  between each pair of nodes of the plant. The transpiration from the epiphyte (T), absorption of atmospheric water vapor  $(q_h)$ , and evaporation from the tank storage (E) are equivalent to the conductance (g) multiplied by the difference in specific humidity between the surface and the atmosphere.



Figure 3: Pressure-volume curve for *G. monostachia*. The solid line is a fit to the data from Zotz and Andrade (1997) using a nonlinear plant capacitance equation (See Appendix C). The water potential  $(\Psi_w)$  is the sum of the osmotic pressure ( $\Omega$ ) and turgor pressure ( $\Pi$ ).



Figure 4: Environmental inputs at different heights within the canopy. Atmospheric temperature  $(T_a)$ , solar radiation ( $\phi$ ), and atmospheric humidity  $(q_a)$  vary throughout time with canopy height. Data are from the first week of February 2010 from the Lutz meteorological tower on Barro Colorado Island, Panama (Physical Monitoring Program of the Smithsonian Tropical Research Institute).



Figure 5: (A) Observed changes in transpirational water stress of *G. monostachia* for a 12-day rainless period during the 1995 dry season in Barro Colorado Island, Panama (Zotz & Andrade, 1997) in comparison to the model transpiration results. The observed data are daily means  $\pm$  SD for 16 plants, in mm/d per unit leaf area. (B) Percent error in daily transpiration between model and experimental data.



Figure 6: (A) Observed changes in carbon assimilation of *G. monostachia* for a 6-day dry period under lab conditions (Pierce et al., 2002) in comparison to the model carbon assimilation results with C3 photosynthesis. The carbon assimilation rates are measured per unit leaf area. (B) Percent error in peak daily assimilation between model and experimental data.



Figure 7: Model results for a 7-day dry-down of an epiphyte with C3 photosynthesis, starting with a full tank and full plant capacitance. (A) Free-surface evaporation from tank (*E*, mm/d per unit ground area). (B) Uptake by the epiphyte from the tank storage ( $q_t$ , mm/d per unit ground area). (C) Relative tank storage ( $x_t$ ). (D) Flux between internal plant storage and xylem ( $q_w$ , mm/d per unit ground area). (E) Relative plant capacitance ( $x_w$ ). (F) Transpiration (*T*, mm/d per unit ground area) for a 7-day dry-down of an epiphyte, starting with a full tank and full plant capacitance. (G) Leaf water potential ( $\psi_l$ , MPa). (H) Carbon assimilation ( $A_n$ ,  $\mu$ mol/(m<sup>2</sup> s) per unit leaf area). (I) Stomatal conductance ( $g_s$ , mol/(m<sup>2</sup> s)). The vertical gray bars indicate nighttime, the vertical white bars indicate daytime, and the red line indicates the time when the tank fully emptied.



Figure 8: Model results for a 7-day dry-down of an epiphyte with CAM photosynthesis, starting with a full tank and full plant capacitance. (A) Free-surface evaporation from tank (*E*, mm/d per unit ground area). (B) Uptake by the epiphyte from the tank storage ( $q_t$ , mm/d per unit ground area). (C) Relative tank storage ( $x_t$ ). (D) Flux between internal plant storage and xylem ( $q_w$ , mm/d per unit ground area). (E) Relative plant capacitance ( $x_w$ ). (F) Transpiration (*T*, mm/d per unit ground area) for a 7-day dry-down of an epiphyte, starting with a full tank and full plant capacitance. (G) Leaf water potential ( $\psi_t$ , MPa). (H) Carbon assimilation ( $A_n$ ,  $\mu$ mol/(m<sup>2</sup> s) per unit leaf area). (I) Stomatal conductance ( $g_s$ , mol/(m<sup>2</sup> s)). The vertical gray bars indicate nighttime, the vertical white bars indicate daytime, and the red line indicates the time when the tank fully emptied.



Figure 9: (A) Average daily transpiration rates (T, mm/d per unit ground area) and (B) average daily water use efficiency (WUE, mmol<sub>CO<sub>2</sub></sub> assimilated per mol<sub>H<sub>2</sub>O</sub> transpired) at different heights within the rainforest canopy.